



Dissolved Organic Matter Influences N₂ Fixation in the New Caledonian Lagoon (Western Tropical South Pacific)

Mar Benavides, Mar Martias, Hila Elifantz, Ilana Berman-Frank, Cecile Dupouy, Sophie Bonnet

► To cite this version:

Mar Benavides, Mar Martias, Hila Elifantz, Ilana Berman-Frank, Cecile Dupouy, et al.. Dissolved Organic Matter Influences N₂ Fixation in the New Caledonian Lagoon (Western Tropical South Pacific). *Frontiers in Marine Science*, 2018, 5, 10.3389/fmars.2018.00089 . hal-02024170

HAL Id: hal-02024170

<https://amu.hal.science/hal-02024170>

Submitted on 7 Jan 2021

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial 4.0 International License



Dissolved Organic Matter Influences N₂ Fixation in the New Caledonian Lagoon (Western Tropical South Pacific)

Mar Benavides^{1,2*}, Chloé Martias², Hila Elifantz³, Ilana Berman-Frank³, Cécile Dupouy² and Sophie Bonnet²

¹ Marine Biology Section, Department of Biology, University of Copenhagen, Helsingør, Denmark, ² Aix-Marseille Université, Université de Toulon, CNRS, IRD, MIO UM 110, Marseille, France, ³ Mina and Everard Goodman Faculty of Life Sciences, Bar-Ilan University, Ramat Gan, Israel

OPEN ACCESS

Edited by:

Beatriz Mouriño-Carballido,
University of Vigo, Spain

Reviewed by:

Ana Fernandez-Carrera,
University of Vigo, Spain
Arvind Singh,

Physical Research Laboratory, India

*Correspondence:

Mar Benavides
mar.benavides@mio.osupytheas.fr

Specialty section:

This article was submitted to
Aquatic Microbiology,
a section of the journal
Frontiers in Marine Science

Received: 30 November 2017

Accepted: 05 March 2018

Published: 22 March 2018

Citation:

Benavides M, Martias C, Elifantz H,
Berman-Frank I, Dupouy C and
Bonnet S (2018) Dissolved Organic
Matter Influences N₂ Fixation in the
New Caledonian Lagoon (Western
Tropical South Pacific).
Front. Mar. Sci. 5:89.
doi: 10.3389/fmars.2018.00089

Specialized prokaryotes performing biological dinitrogen (N₂) fixation (“diazotrophs”) provide an important source of fixed nitrogen in oligotrophic marine ecosystems such as tropical and subtropical oceans. In these waters, cyanobacterial photosynthetic diazotrophs are well known to be abundant and active, yet the role and contribution of non-cyanobacterial diazotrophs are currently unclear. The latter are not photosynthetic (here called “heterotrophic”) and hence require external sources of organic matter to sustain N₂ fixation. Here we added the photosynthesis inhibitor 3-(3,4-dichlorophenyl)-1,1-dimethylurea (DCMU) to estimate the N₂ fixation potential of heterotrophic diazotrophs as compared to autotrophic ones. Additionally, we explored the influence of dissolved organic matter (DOM) on these diazotrophs along a coast to open ocean gradient in the surface waters of a subtropical coral lagoon (New Caledonia). Total N₂ fixation (samples not amended with DCMU) ranged from 0.66 to 1.32 nmol N L⁻¹ d⁻¹. The addition of DCMU reduced N₂ fixation by >90%, suggesting that the contribution of heterotrophic diazotrophs to overall N₂ fixation activity was minor in this environment. Higher contribution of heterotrophic diazotrophs occurred in stations closer to the shore and coincided with the decreasing lability of DOM, as shown by various colored DOM and fluorescent DOM (CDOM and FDOM) indices. We tested the response of diazotrophs (in terms of *nifH* gene expression and bulk N₂ fixation rates) upon the addition of a mix of carbohydrates (“DOC” treatment), amino acids (“DON” treatment), and phosphonates and phosphomonomers (“DOP” treatment). While *nifH* expression increased significantly in *Trichodesmium* exposed to the DOC treatment, bulk N₂ fixation rates increased significantly only in the DOP treatment. The lack of *nifH* expression by gammaproteobacteria, in any of the DOM addition treatments applied, questions the contribution of non-cyanobacterial diazotrophs to fixed nitrogen inputs in the New Caledonian lagoon. While the metabolism and ecology of heterotrophic diazotrophs is currently elusive, a deeper understanding of their ecology and relationship with DOM is needed in the light of increased DOM inputs in coastal zones due to anthropogenic pressure.

Keywords: non-cyanobacterial diazotrophs, *nifH*, DCMU, DOM, CDOM, FDOM

INTRODUCTION

Biological dinitrogen (N_2) fixation provides an important source of fixed nitrogen to fuel primary production in the oceans (Karl et al., 2002), being especially critical in areas devoid of other significant fixed nitrogen sources such as the oligotrophic subtropical gyres. N_2 fixation is performed by specialized prokaryotic microbes called “diazotrophs.” These include cyanobacterial and non-cyanobacterial (bacteria and archaea) groups, that are commonly detected and quantified via the presence of the *nifH* gene, which encodes for a subunit of the nitrogenase enzyme system (Zehr, 2011). Traditionally, marine diazotrophs were hypothesized to be constrained by warm, low inorganic nitrogen waters (Sohm et al., 2011). Yet, several filamentous cyanobacteria are able to obtain energy from dissolved organic matter-DOM- (Rippka et al., 1979), and several reports have shown that the most abundant photoautotrophs in the world’s oceans (*Prochlorococcus* and *Synechococcus*) harbor the genes necessary to use amino acids, peptides and sugars (Yelton et al., 2016). As evidence accumulates, it seems that the ability of using DOM as an alternative source of nutrients under limiting situations is common in various planktonic groups (Stoecker et al., 2017). Among marine diazotrophs, unicellular cyanobacteria such as *Cyanothece* have been reported to grow on glycerol (Feng et al., 2010), and filamentous diazotrophs such as *Trichodesmium* may obtain as much carbon from amino acids assimilation as from CO_2 fixation in the field (Benavides et al., 2017). Non-cyanobacterial and non-photosynthetic diazotrophs require external sources of DOM as energy sources (Riemann et al., 2010). The globally widespread non-cyanobacterial diazotroph Gamma A (Langlois et al., 2015) upregulates *nifH* gene expression upon glucose + mannitol additions (Moisander et al., 2011). Previous work has also shown that N_2 fixation activity is stimulated by DOM additions (Bonnet et al., 2013; Rahav et al., 2013, 2015), and correlates with labile DOM compounds in environments where non-cyanobacterial (non-photosynthetic) diazotrophs predominate, such as the aphotic mesopelagic layer (Benavides et al., 2015), or the surface ultraoligotrophic eastern Mediterranean Sea (Rahav et al., 2016).

N_2 fixation and diazotroph community studies have mostly focused on populations in the North Atlantic (Benavides and Voss, 2015) and North Pacific Oceans, while other areas such as the Indian and the South Atlantic and Pacific Oceans have been chronically under-sampled (Luo et al., 2012). The western tropical Southwest Pacific (WTSP) Ocean comprises warm, oligotrophic, trace metal rich waters that harbor among the high N_2 fixation rates measured at sea ($>500 \mu\text{mol N m}^{-2} \text{ d}^{-1}$; Bonnet et al., 2017). The activity and distribution of diazotrophs in the WTSP has been related to temperature, iron and inorganic nutrient availability thresholds (e.g., Bonnet et al., 2009; Moisander et al., 2010). In austral summer conditions, the relatively warmer and productive waters like the Arafura, Timor and Solomon Seas ($\sim 26.5\text{--}30^\circ\text{C}$) are dominated by *Trichodesmium* and UCYN-B (Messer et al., 2016), while UCYN-A remains confined to the colder and more oligotrophic Coral Sea ($\sim 22\text{--}26.5^\circ\text{C}$; Moisander et al., 2010; Bonnet et al., 2015; Messer et al., 2016). Together with the high abundances

of diazotroph phylotypes (Moisander et al., 2010; Bonnet et al., 2015), the WTSP has been recently acknowledged as a hotspot of global importance for N_2 fixation activity (Bonnet et al., 2017). Thus, understanding how environmental factors (besides temperature and inorganic nutrients) shape the distribution and activity of diazotrophs in the WTSP is of utmost importance to assess their potential to fuel primary production in these oligotrophic waters and their contribution to the oceanic nitrogen budget.

New Caledonia, situated at the eastern edge of the WTSP, is surrounded by a reef enclosing one of the largest coral lagoons in the world. Here, primary production is nitrogen-limited throughout the year (Torréton et al., 2010), prompting diazotrophic activity (Garcia et al., 2007; Biegala and Raimbault, 2008) and the presence of various cyanobacterial diazotroph phylotypes including *Trichodesmium* (often occurring as large blooms; Rodier and Le Borgne, 2008), diatom-diazotroph symbioses (DDAs) as well the unicellular cyanobacteria UCYN-A1, UCYN-A2, UCYN-B, and UCYN-C (Turk-Kubo et al., 2015). However, non-cyanobacterial diazotrophs have been far less studied, and despite representatives such as the putative gammaproteobacterium γ -24774A11 that has been detected at abundances of $10^2\text{--}10^3 \text{ nifH copies L}^{-1}$ within the lagoon (Turk-Kubo et al., 2015), their contribution to N_2 fixation in this environment remains unquantified. Moreover, the importance of DOM as a source of energy and nutrition for both cyanobacterial and non-cyanobacterial diazotrophs has never been explored in this environment. In this study we assess if the inhibition of photosynthesis triggers DOM nutrition in pelagic diazotrophs, and whether the N_2 fixation activity and/or *nifH* expression of these pelagic diazotrophs is enhanced by different model DOM molecules.

MATERIALS AND METHODS

Hydrography, Oxygen, Inorganic Nutrients, Chl α and DOC Measurements

Sampling took place on 13 January 2016 approximately from 8:00 a.m. to noon at four stations along an anthropogenic pressure gradient in the New Caledonian lagoon onboard the R/V *Archamia*. Station 1 was close to the rim of the coral reef barrier and Station 4 near the harbor of the city of Nouméa (Figure 1, Table 1). This transect connects the passage of Dumbéa (the main canal for the entrance of ships into the lagoon) with the harbor. Sampling started at Station 1 at high tide and proceeded toward the coast as the tide ebbed.

At each station, temperature and salinity were measured by casting a SeaBird SB19 plus CTD probe from the surface to the bottom. Water samples for the measurement of oxygen, nutrients, chlorophyll *a* (Chl *a*) and DOC concentrations, and colored DOM (CDOM) and fluorescent DOM (FDOM) parameters (see below) were collected from the subsurface ($\sim 3 \text{ m}$), using a 5 L Niskin bottle and transported to the lab for analysis in coolers equipped with icepacks within 2 h. Samples for the measurement of dissolved oxygen were taken in triplicate and analyzed ashore using an automated

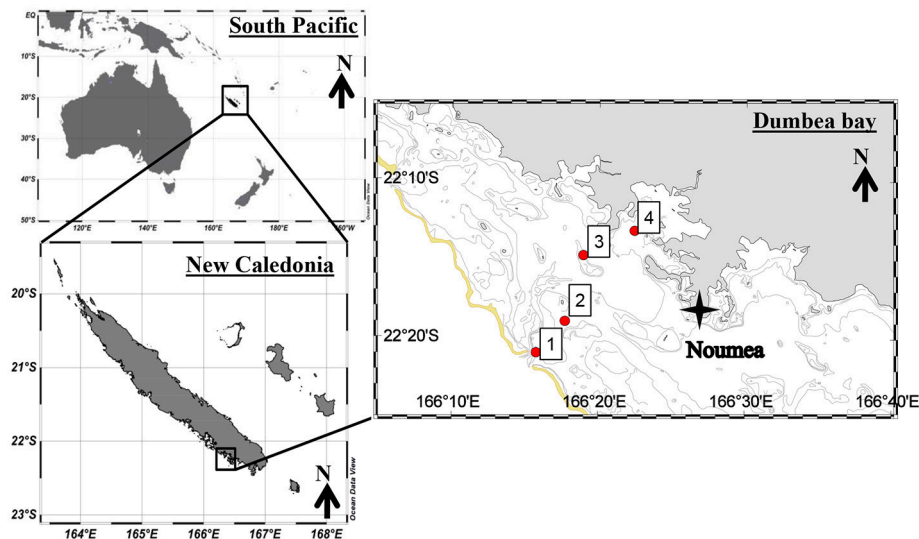


FIGURE 1 | Map of stations sampled in the New Caledonian lagoon. The yellow line represents the coral reef barrier.

titration system with a photometric endpoint (Langdon, 2010). Samples for the determination of nitrate plus nitrite (NO_x) and phosphate (PO_4^{3-}) concentrations were taken on acid-washed 20 mL polyethylene tubes, poisoned with 1% HgCl_2 , and stored at 4°C until analysis. The samples were analyzed on a AA3 Bran+Luebbe autoanalyzer following JGOFS recommendations (JGOFS, 1994). The detection limits were 50 and 10 nM for NO_x and PO_4^{3-} , respectively. Chl *a* concentrations were estimated from seawater samples (500 mL) filtered through 25 mm GF/F filters (Whatman) and stored at −80°C until analysis, extracted in methanol on a AU Turner Designs bench fluorometer, previously calibrated with pure Chl *a* (Herbland et al., 1985). Samples for DOC were collected on triplicate precombusted (450°C, 6 h) glass ampoules and acidified to pH 2 with 50 μL of 50% phosphoric acid. The samples were analyzed by high temperature catalytic oxidation on a TOC-V analyzer (Sohrin and Sempéré, 2005).

FDOM and CDOM Analyses

Samples for the analysis of FDOM and CDOM were filtered through 0.2 μm polycarbonate filters in a custom-made filtration system pre-cleaned with 1 M HCl and Milli-Q water. Samples for FDOM analyses were stored at 4°C until analysis (within 1 month), while samples for CDOM absorption analyses were directly measured on a liquid waveguide capillary flow cell (WPI) according to the method of Kowalczyk et al. (2013). Excitation-emission matrix (EEM) spectra for analysis of FDOM were determined using a Hitachi F-7000 spectrofluorometer (Tedetti et al., 2011). FDOM components were identified by parallel factor analysis (PARAFAC), a statistical decomposition technique used to extract the most representative fluorescent components within EEMs dataset. To increase the validity of the PARAFAC analysis, data from the present experiment were pooled with 39 other samples obtained from the same waters, 3 months before and

after the experiment. Indeed, to avoid systematic biases, remove signals unrelated to fluorescence and to validate the number of FDOM components, a large dataset was needed (Stedmon et al., 2003). For example, in the case where two of the components had very similar fluorescence characteristics and/or were highly correlated in their appearance among samples, the PARAFAC approach would have difficulties to discern between the two. We then determined several spectral indices: the absorption at 442 nm as a proxy of phytoplankton production and the absorption at 350 nm as an indicator of terrestrial inputs in the bay (Passow and Alldredge, 1995), a biological index (BIX) of recent autochthonous FDOM contribution (Huguet et al., 2009), a specific ultraviolet absorbance at 254 nm of DOC (SUVA_{254}), which is an indicator of aromatic DOM (Weishaar et al., 2003), and the ratio of spectral slopes (S_r) at 275–295 nm ($S_{275-295}$) and 350–400 nm ($S_{350-400}$) ranges, related to the molecular weight shift of DOM (Helms et al., 2008; Table 2).

N_2 and Carbon Fixation Rates

Seawater for the quantification of N_2 and carbon fixation rates was sampled using a Teflon pump connected to polyethylene tubing. Bulk N_2 and carbon fixation rates were measured at each transect station on triplicate 2.3 L transparent polycarbonate bottles filled to overflow and closed air-free with septum screwcaps. Once in the lab (within ~2 h) the bottles were amended with stable isotopes to assay N_2 and carbon fixation simultaneously. Each bottle was amended with $\text{NaH}^{13}\text{CO}_3$ 99 atom% ^{13}C to a final ~10 atom% enrichment, and 2.5 mL 98.9 atom% $^{15}\text{N}_2$ gas (both isotopes acquired from Cambridge Isotope Laboratories) injected through the septum using a gas-tight syringe (Montoya et al., 1996). Commercial $^{15}\text{N}_2$ gas stocks have been recently reported to be contaminated with varying amounts of other nitrogenated forms such as nitrous oxide and nitrite (Dabundo et al., 2014). The $^{15}\text{N}_2$ gas batch used here was

TABLE 1 | Position, core parameters, CDOM and FDOM information obtained from each transect station.

Station	Bottom depth (m)	Temperature (°C)	Salinity	Oxygen (μmol kg ⁻¹)	NO _x (μM)	PO ₄ ³⁻ (μM)	Chl <i>a</i> (μg L ⁻¹)	DOC (μM)	<i>a</i> ₄₄₂ (m ⁻¹)	<i>a</i> ₃₅₀ (m ⁻¹)	<i>S</i> _{275–295} (nm ⁻¹)	Sr	SUVA ₂₅₄	BIX	Humic-like (QSU)	Tyrosine-like (QSU)	Tryptophan-like (QSU)	UCYN-A (<i>nifH</i> gene copies L ⁻¹)	UCYN-B (<i>nifH</i> gene copies L ⁻¹)	<i>Trichodesmium</i> (<i>nifH</i> gene copies L ⁻¹)	γ-24774A11 (<i>nifH</i> gene copies L ⁻¹)
1	19	26.97	35.59	201.6 ± 2.5	0.1 ± 0.0	0.1 ± 0.0	0.22 ± 0.0	80.3 ± 2.5	0.41	0.22	0.03	1.90	1.54	1.01	5.0	26.40	0.19	7.6 ± 1.1 × 10 ⁴	3.2 ± 1.3 × 10 ⁴	1.2 ± 0.2 × 10 ⁶	1.8 ± 0.4 × 10 ⁶
2	10	27.31	35.61	200.2 ± 0.3	0.1 ± 0.0	0.1 ± 0.0	0.27 ± 0.1	97.7 ± 0.1	0.07	0.13	0.04	2.78	1.21	1.01	4.83	9.33	0.82	2.9 ± 0.6 × 10 ⁴	7.8 ± 2.5 × 10 ³	8.4 ± 1.6 × 10 ⁴	7.5 ± 3.6 × 10 ⁵
3	8	28.08	35.52	197.9 ± 0.4	0.1 ± 0.0	0.1 ± 0.0	0.16 ± 0.1	81.6 ± 3.8	0.06	0.13	0.04	3.30	1.17	1.03	4.84	5.46	n/d	n/d	n/d	9.4 ± 2.5 × 10 ⁵	8.3 × 10 ⁵
4	18	28.23	35.57	194.4 ± 0.9	0.0 ± 0.0	0.1 ± 0.0	0.28 ± 0.0	87.9 ± 5.6	0.40	0.24	0.03	1.99	0.48	1.03	7.14	5.16	0.17	1.6 ± 1.5 × 10 ⁵	1.3 ± 0.5 × 10 ⁴	2.0 ± 0.6 × 10 ⁶	n/d

Standard deviation values (± SD) are provided where available.

TABLE 2 | Spectral indices investigated in this study.

Abbreviation	Name	Formula
HIX	Humification Index	Ex: 255 nm Em: $\frac{\int 480-434}{\int 344-300}$
BIX	Biological Index	Ex: 255 nm Em: $\frac{310-380}{310-430}$
SUVA ₂₅₄	Specific UV Absorbance (254 nm)	$\frac{ABS(254)}{DOC}$
Sr	Spectral ratio	$\frac{S_{275-295}}{S_{350-400}}$

tested for contamination levels at J. Granger's lab (University of Connecticut, CT, USA), which were found to be negligible and hence did not affect our N₂ fixation rates (Benavides et al., 2015). After isotope additions, the incubation bottles were inverted 20 times, transferred to running seawater incubators covered with blue light screening reproducing the light intensity at the corresponding sampling depth, and incubated for 24 h. To test for non-cyanobacterial (non-photoautotrophic) N₂ fixation activity, a second set of incubation bottles was amended with DCMU to a final concentration of 50 μM (Rahav et al., 2015). These bottles were covered with black fabric bags to keep them in the dark at the same conditions as described above.

After incubation, the samples were filtered through precombusted GF/F filters (Whatman), and the filters were stored at −20°C until analysis. Background (time zero δ¹⁵N and δ¹³C) samples were taken at every station to allow an accurate calculation of N₂ and carbon fixation rates. The samples were analyzed by continuous-flow isotope ratio mass spectrometry (IRMS) using an Integra2 Analyser (Integra CN), calibrated with International Atomic Energy Agency standards IAEA 310-A for nitrogen (47.2‰) and IAEA 303-B for carbon (466‰). The minimum quantifiable N₂ fixation rate was 0.0041 nmol N L⁻¹ d⁻¹, as determined by error propagation analysis (Gradoville et al., 2017). We note that the use of the traditional ¹⁵N₂ “bubble method” (Montoya et al., 1996) instead of the “dissolved method” (Mohr et al., 2010) may have resulted in an underestimation of the actual rates (Grokopf et al., 2012), although the long incubation time (24 h) likely reduces this possibility.

DOM Addition Experiments

To test for the enhancement of N₂ fixation rates and/or *nifH* gene expression as a response to DOM enrichments, we sampled seawater in 25 extra 2.3 L polycarbonate bottles at Station 2. Five bottles were amended with a mix of carbohydrates —“DOC treatment”- (sodium pyruvate, sodium acetate and glucose), another five with a mix of amino acids —“DON treatment”- (alanine, leucine, and glutamic acid), and another five were amended with a mix of phosphomonoesters and phosphonates —“DOP treatment”- (methylphosphonic acid, 2-aminoethylphosphonic acid and fructose 1,6-biphosphate). All DOM mixtures were prepared to be equimolar in carbon (all contained 4 μM carbon; Benavides et al., 2015). Five bottles were used as a control (no DOM additions), and another five bottles were used as a “time zero” (i.e., sacrificed at the start of the experiment to know the baseline conditions). All bottles were spiked with 2.5 mL ¹⁵N₂ and incubated for 24 h as described above. At the end of the incubation period, samples for NO_x,

PO_4^{3-} and DOC analyses were taken from three bottles of each treatment (i.e., DOC, DON, DOP) as described above. The fourth and fifth bottles from each treatment were kept for RNA analyses, respectively (see below).

DNA and RNA Sampling, Extraction and qPCR Assays

Samples for DNA extractions were obtained at each transect station in 2.3 L polycarbonate bottles. Once ashore, the samples were filtered through 0.2 μm Supor filters (PALL) using a peristaltic pump. Upon filtration the filters were transferred to bead beating tubes including ~ 50 μL of a mixture of 0.1 mm and 0.5 mm diameter glass beads (BioSpec Products), and stored at -80°C until analysis. DNA was extracted via the phenol-chloroform method (Massana et al., 1997). RNA samples were taken in duplicate from each treatment (see “DOM addition experiments” above), filtered through 0.2 μm Supor filters and stored in sterile bead beating tubes containing 350 μL RLT buffer (Invitrogen) and 3.5 μL β -mercaptoethanol, and finally stored at -80°C until analysis. The RNA was extracted by TRI Reagent (Invitrogen). 1 mL of Tri reagent was added to each sample that was disrupted and homogenized by pipetting and vortexing in order to break the cells. The cell lysate was incubated at room temperature for 3 h with occasional vortexing. The samples were then centrifuge at 12,000 g for 20 min at 4°C . The supernatant was transferred to a new tube and 10% of Phase Separation Reagent, BCP (Molecular Research Center, Inc., Cincinnati, OH, USA) was added. After additional incubation (30 min) at room temperature and centrifugation (12,000 g, 15 min, 4°C) the supernatant was transferred to a new tube and a cleanup and DNase treatment was performed using the RNA Clean and Concentrator-5 kit (Zymo Research, Irvin, CA, USA). Complementary DNA (cDNA) was created using the qScriptTM cDNA Synthesis Kit (Quantabio, Beverly, MA, USA) using the reverse primers of the *nifH* (*nifH2* and *nifH3*; Zani et al., 2000). Quantitative PCR (qPCR) assays were performed on a CFX-96 Touch Real-Time PCR Detection System (Bio-Rad) using the TaqMan Fast Advanced Master Mix (Applied Biosystems) and primers for *nifH* of UCYN-A, UCYN-B, *Trichodesmium* and γ -24774A11 (here referred to as gammaproteobacteria) (Turk et al., 2011).

RESULTS

Hydrography, Nutrients and *in Situ* DOM

The temperature of surface waters increased from 26.97 to 28.23 $^\circ\text{C}$ moving onshore from Station 1 toward Station 4, while salinity remained close to 35.6 at all stations (Table 1). The concentration of oxygen decreased onshore from 201 to 194 $\mu\text{mol kg}^{-1}$, while inorganic nutrient concentrations (NO_x and PO_4^{3-}) did not vary significantly along the transect, and neither did Chl *a* concentrations, with the exception of Station 3 where Chl *a* was slightly lower (Table 1). DOC concentrations remained between 80 and 90 μM along the transect, showing a peak of ~ 98 μM at Station 2 (Table 1).

CDOM absorption at 442 nm and 350 nm (a_{442} and a_{350}) presented lower values at station 2 and 3 as compared to the

values recorded at Stations 1 and 4 (Table 1) which indicated a decrease in CDOM concentrations. Moreover, the spectral slope $S_{275-295}$ and the spectral ratio S_r were high with values >1 and >0.03 nm^{-1} , respectively, which indicated the presence of photodegraded DOM with low molecular weight. In addition, the SUVA_{254} indices were <4 (this reference value is indicative of DOM with low aromaticity) with an average value of 1.09 ± 0.44 . Then CDOM at 442 nm and 350 nm is strongly photodegraded with a low aromaticity. Three fluorescent DOM compounds were identified by the PARAFAC model: terrestrial humic-like (compound 1), tyrosine-like (compound 2) and tryptophan-like (compound 3; Table 3). Humic-like compound signatures were higher closer to the coast (Station 4) and decreased to values around ~ 5 QSU at the other stations. This pattern is opposite to that of tyrosine-like compounds signals, which increased toward the barrier reef. Tryptophane-like compound concentrations did not follow any coast-open ocean pattern but presented the highest concentrations at Station 2. The BIX indices were homogenous and >1 (this value is indicative of an enrichment in DOM from bacterial origin) along the transect (Table 1). Overall, the optical analyses of DOM suggest that while river-derived refractory DOM was more abundant close to the coast, biological-derived labile DOM compounds were present homogeneously at all stations along the transect.

N_2 and Carbon Fixation Activity and Diazotroph Abundance

No clear spatial trends were observed in either N_2 or carbon fixation rates along the transect. N_2 fixation rates in samples not amended with DCMU ranged from 0.66 to 1.32 $\text{nmol N L}^{-1} \text{d}^{-1}$, with the highest rates measured at Station 3 (Figure 2A). Carbon fixation rates in samples without DCMU ranged from 1.50 to 2.14 $\mu\text{g C L}^{-1} \text{d}^{-1}$ (Figure 2B). Both N_2 and carbon fixation were effectively inhibited in DCMU-amended samples ($>90\%$; Figures 2A,B).

Among the diazotroph phylotypes quantified by qPCR, *Trichodesmium* and gammaproteobacteria showed the highest abundance, ranging between 8.4×10^4 and 2×10^6 , and 7.5×10^5 and 1.8×10^6 *nifH* copies L^{-1} , respectively (Table 1). UCYN-A and UCYN-B phylotypes were one to two orders of magnitude lower, and undetectable at station 3 (Table 1).

N_2 Fixation, Diazotroph Abundance and Optical Response to Model DOM Compound Additions

At Station 2 we tested whether the addition of DOC, DON or DOP upregulated *nifH* expression and/or enhanced N_2 fixation rates. N_2 fixation rates were similar in the control, DOC and DON treatments (~ 0.8 $\text{nmol N L}^{-1} \text{d}^{-1}$), yet they doubled to ~ 2 $\text{nmol N L}^{-1} \text{d}^{-1}$ when DOP was added (Figure 3). However, due to the variability among replicates, the enhancement of either N_2 fixation or expression of *nifH* genes was not statistically significant in any treatment (Kruskal-Wallis, $p > 0.05$).

The DOC, DON, and DOP treatments yielded slightly higher abundances of UCYN-A *nifH* transcripts compared to transcripts measured in the controls, yet similar transcript

TABLE 3 | Excitation and Emission maxima (Ex/Em) of three CDOM fluorescent components validated by PARAFAC model and identification by comparison with the literature data.

	This study			Previous study				FDOM sources
	Comp.	Ex(nm)	Em(nm)	Comp.	Ex(nm)	Em(nm)	Reference	
COMPOUND 1	Humic like	<220	454(322)	Humic-like (A)	230–260	400–480	Coble, 1996	Allochthonous: terrestrial inputs from river
				Terrestrial humic-like	<260	454	Yamashita et al., 2010	
				HumiC-like (A)	230	476	Tedetti et al., 2016	
COMPOUND 2	Tyrosine like	225(275)	304	C7	240(300)	338	Murphy et al., 2007	Autochthonous: heterotrophic bacterial degradation of DOM
				Component 8	275	304	Stedmon and Markager, 2005	
				Tyrosine-like (B)	225(275)	304	Tedetti et al., 2016	
COMPOUND 3	Tryptophane like	270(335)	340	Tryptophane (T1)	220–235	334–360	Coble, 1996	Autochthonous: coral production, heterotrophic bacterial degradation of DOM
					230(300)	352	Martias et al., 2018	

numbers were with obtained in each of the three treatments ($\sim 5\text{--}6 \times 10^3$ *nifH* transcripts L^{-1} ; **Figure 3**). UCYN-B did not show differential expression among treatments, while *Trichodesmium* increased by one order of magnitude when exposed to the DOC treatment (**Figure 3**). Despite being the most abundant ambient group as determined by DNA qPCR counts, gammaproteobacteria did not show detectable *nifH* gene expression in our DOM enrichment experiments (including control bottles). The prolonged incubation (24 h) in polycarbonate bottles could have affected *nifH* expression in this group (Moisander et al., 2014).

Regarding optical parameters, the DOC treatment resulted in a peak of tyrosine-like compounds concentrations (+51%) whereas no changes were recorded in humic-like and tryptophan-like compounds concentrations (**Table 4**). The effect of DON and DOP treatments on concentrations of fluorescent compounds are not shown as FDOM results were considered outliers in the validation of PARAFAC model.

DISCUSSION

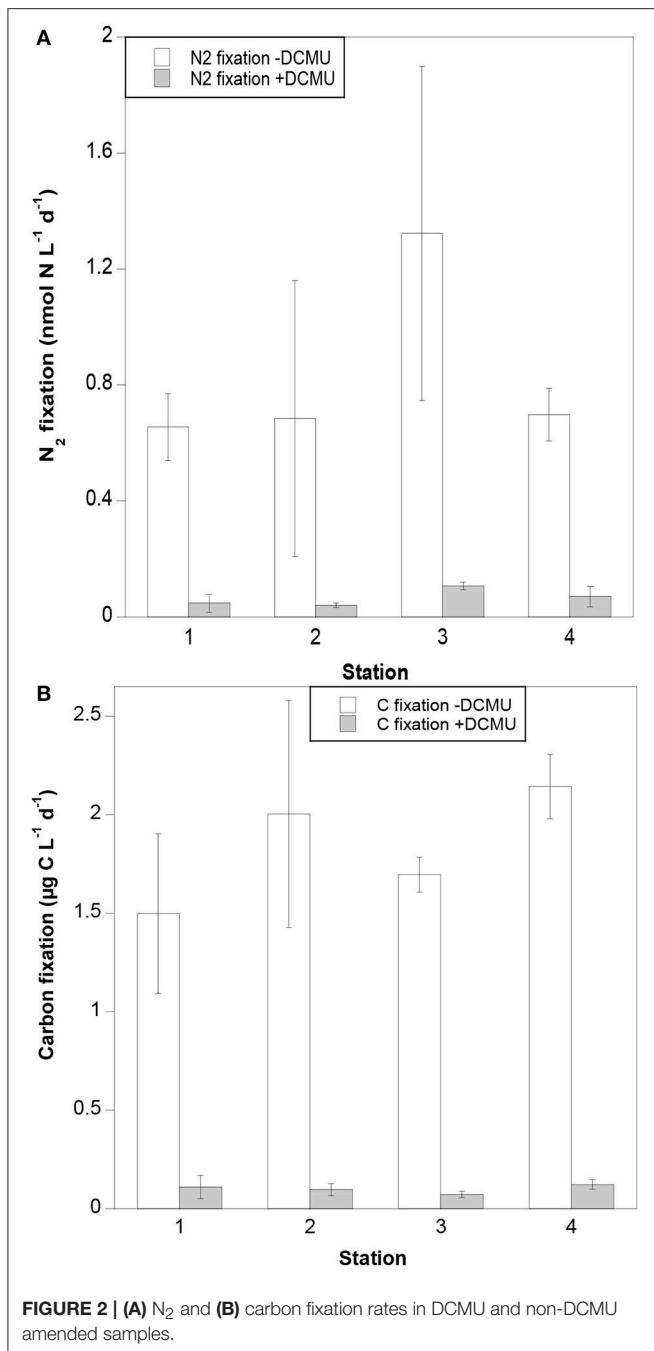
Spatial Variability of N_2 Fixation Rates and Diazotrophic Phylotypes

Gammaproteobacteria and *Trichodesmium* were the most abundant diazotrophs among the four phylotypes quantified (**Table 1**). The *nifH* copies of these two phylotypes were higher than in a previous study performed in the lagoon during the austral summer. The abundances, reported by Turk-Kubo et al. in *nifH* copies, ranged from 2.8×10^3 to 6.4×10^4 (UCYN-A), 3.5×10^2 to 2.7×10^3 (UCYN-B), 3.4×10^2 and 1.4×10^5 (*Trichodesmium*), and 4×10^2 to 1.9×10^3 (gammaproteobacteria) *nifH* copies L^{-1} . Typically, the *nifH* copy numbers reported here are similar to published numbers from Berthelot et al. (2017) who reported abundances up to 1.3×10^3 (UCYN-A), 7.5×10^4 (UCYN-B), 1.6×10^5 (*Trichodesmium*),

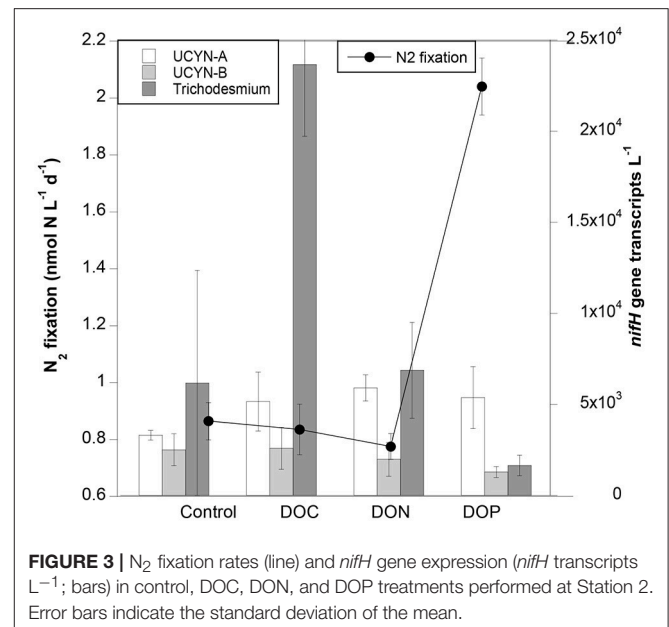
and 2.7×10^4 (gammaproteobacteria) *nifH* copies L^{-1} from the Solomon Sea.

DCMU blocks photosynthesis and thus the energy source required by autotrophic cyanobacterial diazotrophs to fix N_2 . Hence, measured N_2 fixation rates from DCMU-amended samples should be representative of non-cyanobacterial N_2 fixation. The lack of *nifH* expression in gammaproteobacteria in our DOM addition experiments (see below) questions whether this specific group was active. Nevertheless, other non-cyanobacterial diazotrophs such as *Bradyrhizobium* and *Mesorhizobium*, previously detected in the lagoon (Pfreundt et al., 2016), could be responsible for the measured N_2 fixation rates in DCMU-amended bottles. N_2 fixation rates in DCMU-amended incubations represented 4–7% of N_2 fixation measured without amendments. The New Caledonian lagoon is dominated by autotrophic cyanobacterial diazotrophs (Rodier and Le Borgne, 2008; Turk-Kubo et al., 2015), where non-cyanobacterial phylotypes have been reported to represent <2% of the diazotrophic community (Turk-Kubo et al., 2015; Pfreundt et al., 2016). Hence, such a low contribution to total N_2 fixation rates may not be surprising, but deserves further study in other parts of the lagoon and other times of the year. Still, N_2 fixation rates in DCMU-amended bottles ranged between 0.04 and 0.11 $\text{nmol N L}^{-1} \text{d}^{-1}$, which is in the range of N_2 fixation rates measured in environments where non-cyanobacterial diazotrophs prevail (Moisander et al., 2017).

It has to be noted that putative non-cyanobacterial N_2 fixation was higher at stations closer to the coast (Stations 3 and 4), which may be linked to the Dumbéa river inputs importing bioavailable CDOM at 350 nm and terrestrial humic-like DOM compounds into the bay. Spectral indices (S_r and $S_{275-295}$) revealed low molecular weight and photodegraded terrestrial CDOM/FDOM at Station 4 (the most impacted by the river), which has been seen to enhance its availability



to heterotrophic bacteria in oligotrophic waters (Su et al., 2017). At Station 1, DOM sources come from autochthonous phytoplankton production), which is also available for bacterial growth (based on a_{442} values >0.1 as compared to Stations 2 and 3; Steinberg et al., 2004). The observed relationship between DOM lability proxies and non-cyanobacterial N₂ fixation activity in the present study support the findings of previous studies in the WTSP and the Mediterranean Sea (Benavides et al., 2015, 2016; Rahav et al., 2016). Nevertheless, a mechanistic understanding of how non-cyanobacterial diazotrophs benefit from different DOM molecules (e.g., metabolic pathways,



affinity for different substrates) is still lacking (Bombar et al., 2016).

Diazotrophic Responses to DOM Additions

The influence of the three different DOM treatments applied (DOC, DON, and DOP) differentially influenced *nifH* gene expression of different phylotypes detected, with transcript abundance increasing similarly in UCYN-A for all amendments as compared to the controls (from ~ 3 to $5\text{--}6 \times 10^3$ *nifH* transcripts L⁻¹), while no differences were observed between controls and amendments for UCYN-B transcripts (Figure 3). UCYN-B fixes carbon during the day and N₂ at night to avoid nitrogenase enzyme inhibition by photosynthetically derived oxygen. Thus, *nifH* expression in UCYN-B is typically higher at night (Church et al., 2005; Zehr et al., 2007). Hence, the absence of *nifH* expression in UCYN-B during the day (when our RNA samples were taken) is not surprising, even in the presence of external DOM sources. Similarly, Moisander et al. (2011) observed that *nifH* expression was not upregulated in UCYN-B when incubated with a mixture of glucose and mannitol. UCYN-A do not have a tricarboxylic acid pathway (Tripp et al., 2010), and thus require reduced organic carbon sources provided by their hosts (Thompson et al., 2012). It is therefore plausible that DOM additions in our experiments enhanced the diazotrophic activity of UCYN-A, similar to the experiments of Moisander et al. (2011), who demonstrated upregulated *nifH* gene expression in UCYN-A upon the addition of a glucose and mannitol mixture. In our experiments, however, a preference for DOC, DON or DOP was not observed and the *nifH* expression response was similar across treatments (Figure 3).

Active assimilation of both DOC and DON was measured in natural *Trichodesmium* colonies from the WTSP open ocean waters in similar experiments (Benavides et al., 2017). However,

TABLE 4 | NO_x , PO_4^{3-} , DOC concentrations and values of FDOM-related ratios for the different DOM additions (DOC, DON, and DOP treatments) performed at Station 2.

Treatment	NO_x (μM)	PO_4^{3-} (μM)	DOC (μM)	Humic-like (QSU)	Tyrosine-like (QSU)	Tryptophan-like (QSU)
Control	0.1 ± 0.0	0.1 ± 0.0	90.6 ± 15.5	6.54	44.79	11.71
DOC	0.2 ± 0.1	0.1 ± 0.0	98.6 ± 3.3	6.57	67.45	11.53
DON	0.2 ± 0.0	0.1 ± 0.0	104.1 ± 30.6	n/d	n/d	n/d
DOP	0.2 ± 0.0	0.3 ± 0.0	107.3 ± 7.5	n/d	n/d	n/d

Standard deviation values ($\pm$ SD) are provided where available.

in that previous study only the addition of DON resulted in an increase of filament-specific N_2 fixation rates (Benavides et al., 2017). The bulk water N_2 fixation measurements presented here do not allow us to discern the contribution of each diazotroph phylotype. Nevertheless, while bulk N_2 fixation rates were highest in the DOP treatment, *nifH* gene expression by *Trichodesmium* was highest in the DOC treatment (Figure 3). The peak of tyrosine-like compounds coinciding with the highest *nifH* expression of *Trichodesmium* observed in the DOC treatment may be related to the production of CDOM by this cyanobacterium (Steinberg et al., 2004). Surprisingly, DOM additions did not result in any *nifH* expression in the group of non-cyanobacterial diazotrophs tested in this study. In marine waters, maximum *nifH* expression of gammaproteobacteria typically reaches $\times 10^4$ *nifH* transcripts L^{-1} (Luo et al., 2012; Langlois et al., 2015). For example, Church et al. (2005) reported an average of 4×10^2 *nifH* transcripts L^{-1} for gammaproteobacteria in a diel cycle study at station ALOHA in the North Pacific Ocean, which is in the same order of magnitude as in a study in the WTSP by Moisander et al. (2014). Nevertheless, a study including samples from a wide range of oceanographic cruises across the North and South Atlantic and Pacific Oceans reported up to $\times 10^5$ gammaproteobacterial *nifH* transcripts per liter in sunlit surface waters (Langlois et al., 2015). In offshore waters of the WTSP, the addition of glucose and mannitol upregulated *nifH* expression in gammaproteobacteria (Moisander et al., 2011).

The lack of significant responses to DOM additions in this study could reflect sufficient DOM availability (based on $\text{Sr}_{5275-295}$ and SUVA_{254} indexes) potentially provided by corals in the lagoon which are known for producing tryptophane-like compounds (available DOM sources for bacteria; Weishaar et al., 2003). However, the absence of detected gammaproteobacterial *nifH* transcripts, also in the control bottles, suggests that bottle confinement effects may have inhibited this group (Moisander et al., 2014). In addition to the diel changes in *nifH* gene expression between phylotypes (Church et al., 2005; Turk et al., 2011), it is possible that the *nifH* expression measured in our samples does not reflect the immediate response of diazotrophs to DOM additions. It must be noted that while RNA samples represent a snapshot of *nifH* expression at the time of sampling, N_2 fixation rates integrate the nitrogenase activity that has taken place over a period of 24 h. Such differences probably also account for the uncoupling observed between N_2 fixation rates and *nifH* gene expression.

N_2 Fixation Rates Enhanced by DOP Additions

Diazotrophs differ in how they access DOP. For example, *Crocospaera watsonii* (here represented by the UCYN-B phylotype) obtains inorganic phosphorus using high affinity phosphate binding systems, and organic phosphorus sources such as phosphomonoesters utilizing alkaline phosphatases (Dyhrman and Haley, 2006). In addition to these phosphorus sources, *Trichodesmium* is also able to grow on phosphonates and use phosphite (Dyhrman et al., 2006; Polyviou et al., 2015). It is intriguing why the *nifH* gene expression of *Trichodesmium* or of any of the other diazotroph phylotypes tested was not enhanced when DOP was enriched, and deserves further study. However, an enhancement of N_2 fixation activity was observed in the DOP treatment (Figure 3), and could be due to inorganic phosphorus limitation alleviation. Unfortunately, we do not have alkaline phosphatase activity data to confirm whether the DOP added to our incubations was ultimately transformed to inorganic phosphorus compounds, but PO_4^{3-} concentrations nearly doubled in DOP-amended samples with respect to the control (data not shown), suggesting that active alkaline phosphatases were functioning. The fact that N_2 fixation rates enhanced by DOP additions was not reflected in a significant enhancement of *nifH* expression of any of the diazotroph phylotypes tested suggests that the N_2 fixation activity measured may be provided by a diazotroph not accounted for in our *nifH* gene q-RT-PCR assays. Alternatively, transcription may be interrupted when the number of transcripts is enough to activate the nitrogenase protein, resulting in an uncoupling between expression and nitrogenase activity, as previously reported in *Trichodesmium* cultures (Levitan et al., 2010).

CONCLUSIONS

Although cyanobacterial diazotrophs predominate in the New Caledonian lagoon, here we demonstrate that non-cyanobacterial diazotrophs contribute $\sim 10\%$ of total N_2 fixation rates. Comparable to previous findings (see Moisander et al., 2017) gammaproteobacteria were the most abundant based on DNA-derived qPCR *nifH* gene counts, yet we did not detect any *nifH* gene expression for this group throughout the experiments. This lack of detectable expression together with the low N_2 fixation rates measured in DCMU-treated samples questions the contribution of non-cyanobacterial diazotrophs to biologically fixed nitrogen inputs in the lagoon. Yet, the response of the

cyanobacterial diazotroph phylotypes to the amended DOM treatments suggests that mixotrophic nutrition may be important in these organisms and should be investigated further.

AUTHOR CONTRIBUTIONS

Conceived and designed the experiments: MB, CM, SB; Performed the experiments: MB, CM; Analyzed the data: MB, CM, HE; Contributed reagents, materials, analysis tools: CM, SB, HE, CD, IB-F; Wrote the paper: MB, CM, CD, HE, IB-F, SB.

FUNDING

MB was funded by the People Programme (Marie Skłodowska-Curie Actions) of the European Union's Seventh Framework

Programme (FP7/2007-2013) under REA grant agreement number 625185. Funding was obtained for IB-F through a collaborative grant from Ministry of Science and Technology (MOST) Israel and the High Council for Science and Technology (HCST)-France (2013). CM and SB were funded by the Institut de Recherche pour le Développement (IRD, France).

ACKNOWLEDGMENTS

We thank the captain of R/V *Archamia* for his invaluable help at sea. We would like to thank O. Grosso for IRMS analyses. We are also grateful to P. Gérard and L. Jamet from the Laboratoire des Moyens Analytiques (LAMA) at IRD in Nouméa (New Caledonia) for oxygen, nutrient and DOC analyses.

REFERENCES

- Benavides, M., Berthelot, H., Duhamel, S., Raimbault, P., and Bonnet, S. (2017). Dissolved organic matter uptake by *Trichodesmium* in the Southwest Pacific. *Sci. Rep.* 7, 1–6. doi: 10.1038/srep41315
- Benavides, M., Bonnet, S., Hernández, N., Martínez-Pérez, A. M., Nieto-Cid, M., Álvarez-Salgado, X. A., et al. (2016). Basin-wide N₂ fixation in the deep waters of the Mediterranean Sea. *Glob. Biogeochem. Cycles* 30, 1–19. doi: 10.1002/2015GB005326
- Benavides, M. H., Moisaner, P., Berthelot, H., Dittmar, T., Grosso, O., and Bonnet, S. (2015). Mesopelagic N₂ fixation related to organic matter composition in the Solomon and Bismarck Seas (Southwest Pacific). *PLoS ONE* 10:e0143775. doi: 10.1371/journal.pone.0143775
- Benavides, M., and Voss, M. (2015). Five decades of N₂ fixation research in the North Atlantic Ocean. *Front. Mar. Sci.* 2, 1–20. doi: 10.3389/fmars.2015.00040
- Berthelot, H., Benavides, M., Moisaner, P. H., Grosso, O., and Bonnet, S. (2017). High-nitrogen fixation rates in the particulate and dissolved pools in the Western Tropical Pacific (Solomon and Bismarck Seas). *Geophys. Res. Lett.* 44, 8414–8423. doi: 10.1002/2017GL073856
- Biegala, I. C., and Raimbault, P. (2008). High abundance of diazotrophic picocyanobacteria (<3 μm) in a Southwest Pacific coral lagoon. *Aquat. Microb. Ecol.* 51, 45–53. doi: 10.3354/ame01185
- Bombar, D., Paerl, R. W., and Riemann, L. (2016). Marine non-cyanobacterial diazotrophs: moving beyond molecular detection. *Trends Microbiol.* 24, 916–927. doi: 10.1016/j.tim.2016.07.002
- Bonnet, S., Biegala, I. C., Dutrieux, P., Slemmons, L. O., and Capone, D. G. (2009). Nitrogen fixation in the western equatorial Pacific: rates, diazotrophic cyanobacterial size class distribution, and biogeochemical significance. *Global Biogeochem. Cycles* 23, 1–13. doi: 10.1029/2008GB003439
- Bonnet, S., Caffin, M., Berthelot, H., and Moutin, T. (2017). Hot spot of N₂ fixation in the western tropical South Pacific pleads for a spatial decoupling between N₂ fixation and denitrification. *Proc. Natl. Acad. Sci. U.S.A.* 114, E2800–E2801. doi: 10.1073/pnas.1619514114
- Bonnet, S., Dekaezemaeker, J., Turk-Kubo, K. A., Moutin, T., Hamersley, R. M., Grosso, O., et al. (2013). Aphotic N₂ fixation in the eastern tropical South Pacific Ocean. *PLoS ONE* 8:e81265. doi: 10.1371/journal.pone.0081265
- Bonnet, S., Rodier, M., Turk-Kubo, K. A., Germineaud, C., Menkes, C., Ganachaud, A., et al. (2015). Contrasted geographical distribution of N₂ fixation rates and *nifH* phylotypes in the Coral and Solomon Seas (southwestern Pacific) during austral winter conditions. *Global Biogeochem. Cycles* 29, 1874–1892. doi: 10.1002/2015GB005117
- Church, M. J., Short, C. M., Jenkins, B. D., Karl, D. M., and Zehr, J. P. (2005). Temporal patterns of nitrogenase gene (*nifH*) expression in the oligotrophic North Pacific Ocean. *Appl. Environ. Microbiol.* 71, 5362–5370. doi: 10.1128/AEM.71.9.5362-5370.2005
- Coble, P. G. (1996). Characterization of marine and terrestrial DOM in seawater using excitation-emission matrix spectroscopy. *Mar. Chem.* 51, 325–346.
- Dabundo, R., Lehmann, M. F., Treibergs, L., Tobias, C. R., Altabet, M. A., Moisaner, P. H., et al. (2014). The contamination of commercial ¹⁵N₂ gas stocks with ¹⁵N-labeled nitrate and ammonium and consequences for nitrogen fixation measurements. *PLoS ONE* 9:e110335. doi: 10.1371/journal.pone.0110335
- Dyhrman, S. T., Chappell, P. D., Haley, S. T., Moffett, J. W., Orchard, E. D., Waterbury, J. B., et al. (2006). Phosphonate utilization by the globally important marine diazotroph *Trichodesmium*. *Nature* 439, 68–71. doi: 10.1038/nature04203
- Dyhrman, S. T., and Haley, S. T. (2006). Phosphorus scavenging in the unicellular marine diazotroph *Crocospaera watsonii*. *Appl. Environ. Microbiol.* 72, 1452–1458. doi: 10.1128/AEM.72.2.1452-1458.2006
- Feng, X., Bandyopadhyay, A., Berla, B., Page, L., Wu, B., Pakrasi, H. B., et al. (2010). Mixotrophic and photoheterotrophic metabolism in *Cyanothece* sp. ATCC 51142 under continuous light. *Microbiology* 156, 2566–2574. doi: 10.1099/mic.0.038232-0
- Garcia, N., Raimbault, P., and Sandroni, V. (2007). Seasonal nitrogen fixation and primary production in the Southwest Pacific: nanoplankton diazotrophy and transfer of nitrogen to picoplankton organisms. *Mar. Ecol. Prog. Ser.* 343, 25–33. doi: 10.3354/meps06882
- Gradoville, M. R., Bombar, D., Crump, B. C., Letelier, R. M., Zehr, J. P., and White, A. E. (2017). Diversity and activity of nitrogen-fixing communities across ocean basins. *Limnol. Oceanogr.* 62, 1895–1909. doi: 10.1002/lno.10542
- Grokovf, T., Mohr, W., Baustian, T., Schunck, H., Gill, D., Kuypers, M. M. M., et al. (2012). Doubling of marine dinitrogen-fixation rates based on direct measurements. *Nature* 488, 361–364. doi: 10.1038/nature11338
- Helms, J. R., Stubbins, A., Ritchie, J. D., Minor, E. C., Kieber, D. J., and Mopper, K. (2008). Absorption spectral slopes and slope ratios as indicators of molecular weight, source, and photobleaching of chromophoric dissolved organic matter. *Limnol. Oceanogr.* 53, 955–969. doi: 10.2307/40058211
- Herbland, A., Le Bouteiller, A., and Raimbault, P. (1985). Size structure of phytoplankton biomass in the equatorial Atlantic Ocean. *Deep. Res.* 32, 819–836. doi: 10.1016/0198-0149(85)90118-9
- Huguet, A., Vacher, L., Relexans, S., Saubusse, S., Froidefond, J. M., and Parlanti, E. (2009). Properties of fluorescent dissolved organic matter in the Gironde estuary. *Org. Geochem.* 40, 706–719. doi: 10.1016/j.orggeochem.2009.03.002
- JGOFS (1994). Protocols For The Joint Global Ocean Flux Study (JGOFS) core measurements, *IOC Manuals Guid.* 29, 180.
- Karl, D., Michaels, A., Bergman, B., Capone, D., Carpenter, E., Letelier, R., et al. (2002). Dinitrogen fixation in the world's oceans. *Biogeochemistry* 57/58, 47–98. doi: 10.1023/A:1015798105851
- Kowalczyk, P., Tilstone, G. H., Zabłocka, M., Röttgers, R., and Thomas, R. (2013). Composition of dissolved organic matter along an Atlantic meridional transect from fluorescence spectroscopy and parallel factor analysis. *Mar. Chem.* 157, 170–184. doi: 10.1016/j.marchem.2013.10.004

- Langdon, C. (2010). "Determination of dissolved oxygen in seawater by Winkler titration using the amperometric technique," in *Go-sh. Repeat Hydrogr. Man. A Collect. Expert Reports Guidel*, 18.
- Langlois, R., Großkopf, T., Mills, M., Takeda, S., and LaRoche, J. (2015). Widespread distribution and expression of Gamma A (UMB), an uncultured, diazotrophic, γ -proteobacterial *nifH* phylotype. *PLoS ONE* 10:e0128912. doi: 10.1371/journal.pone.0128912
- Leviton, O., Brown, C. M., Sudhaus, S., Campbell, D., LaRoche, J., and Berman-Frank, I. (2010). Regulation of nitrogen metabolism in the marine diazotroph *Trichodesmium* IMS101 under varying temperatures and atmospheric CO₂ concentrations. *Environ. Microbiol.* 12, 1899–1912. doi: 10.1111/j.1462-2920.2010.02195.x
- Luo, Y.-W., Doney, S. C., Anderson, L. A., Benavides, M., Berman-Frank, I., Bode, A., et al. (2012). Database of diazotrophs in global ocean: abundance, biomass and nitrogen fixation rates. *Earth Syst. Sci. Data* 4, 47–73. doi: 10.5194/essd-4-47-2012
- Martias, C., Tedetti, M., Lantoine, F., Jamet, L., and Dupouy, C. (2018). Characterization and sources of colored dissolved organic matter in a coral reef ecosystem subject to ultramafic erosion pressure (New Caledonia, Southwest Pacific). *Sci. Total Environ.* 616–617, 438–452. doi: 10.1016/j.scitotenv.2017.10.261
- Massana, R., Murray, A. E., Preston, C. M., and DeLong, E. F. (1997). Vertical distribution and phylogenetic characterization of marine planktonic Archaea in the Santa Barbara channel. *Appl. Environ. Microbiol.* 63, 50–56.
- Messer, L. F., Mahaffey, C., Robinson, C. M., Jeffries, T. C., Baker, K. G., Isaksson, J. B., et al. (2016). High levels of heterogeneity in diazotroph diversity and activity within a putative hotspot for marine nitrogen fixation. *ISME J.* 10, 1499–1513. doi: 10.1038/ismej.2015.205
- Mohr, W., Großkopf, T., Wallace, D. W. R., and LaRoche, J. (2010). Methodological underestimation of oceanic nitrogen fixation rates. *PLoS ONE* 5:e0012583. doi: 10.1371/journal.pone.0012583
- Moisander, P. H., Beinart, R. A., Hewson, I., White, A. E., Johnson, K. S., Carlson, C. A., et al. (2010). Unicellular cyanobacterial distributions broaden the oceanic N₂ fixation domain. *Science* 327, 1512–1514. doi: 10.1126/science.1185468
- Moisander, P. H., Benavides, M., Bonnet, S., Berman-Frank, I., White, A. E., and Riemann, L. (2017). Chasing after non-cyanobacterial nitrogen fixation in marine pelagic environments. *Front. Microbiol.* 8:1736. doi: 10.3389/fmicb.2017.01736
- Moisander, P. H., Serros, T., Paerl, R. W., Beinart, R. A., and Zehr, J. P. (2014). Gammaproteobacterial diazotrophs and *nifH* gene expression in surface waters of the South Pacific Ocean. *ISME J.* 8, 1962–1973. doi: 10.1038/ismej.2014.49
- Moisander, P. H., Zhang, R., Boyle, E. A., Hewson, I., Montoya, J. P., and Zehr, J. P. (2011). Analogous nutrient limitations in unicellular diazotrophs and *Prochlorococcus* in the South Pacific Ocean. *ISME J.* 6, 733–744. doi: 10.1038/ismej.2011.152
- Montoya, J. P., Voss, M., Kahler, P., and Capone, D. G. (1996). A simple, high-precision, high-sensitivity tracer assay for N₂ fixation. *Appl. Environ. Microbiol.* 62, 986–993.
- Murphy, K. R., Stedmon, C. A., Waite, T. D., and Ruiz, G. M. (2007). Distinguishing between terrestrial and autochthonous organic matter sources in marine environments using fluorescence spectroscopy. *Mar. Chem.* 108, 40–58. doi: 10.1016/j.marchem.2007.10.003
- Passow, U., and Alldredge, A. L. (1995). A dye-binding assay for the spectrophotometric measurement of transparent exopolymer particles (TEP) in the ocean. *Limnol. Oceanogr.* 40, 1326–1335. doi: 10.4319/lo.1995.40.7.1326
- Pfreundt, U., Van Wambeke, F., Caffin, M., Bonnet, S., and Hess, W. R. (2016). Succession within the prokaryotic communities during the VAHINE mesocosms experiment in the New Caledonia lagoon. *Biogeosciences* 13, 2319–2337. doi: 10.5194/bg-13-2319-2016
- Polyviou, D., Hitchcock, A., Baylay, A. J., Moore, C. M., and Bibby, T. S. (2015). Phosphite utilization by the globally important marine diazotroph *Trichodesmium*. *Environ. Microbiol. Rep.* 7, 824–830. doi: 10.1111/1758-2229.12308
- Rahav, E., Bar-Zeev, E., Ohayon, S., Elifantz, H., Belkin, N., Herut, B., et al. (2013). Dinitrogen fixation in aphotic oxygenated marine environments. *Front. Microbiol.* 4:227. doi: 10.3389/fmicb.2013.00227
- Rahav, E., Giannetto, M. J., and Bar-Zeev, E. (2016). Contribution of mono and polysaccharides to heterotrophic N₂ fixation at the eastern Mediterranean coastline. *Sci. Rep.* 6:27858. doi: 10.1038/srep27858
- Rahav, E., Herut, B., Mulholland, M. R., Belkin, N., Elifantz, H., and Berman-Frank, I. (2015). Heterotrophic and autotrophic contribution to dinitrogen fixation in the Gulf of Aqaba. *Mar. Ecol. Prog. Ser.* 522, 67–77. doi: 10.3354/meps11143
- Riemann, L., Farnelid, H., and Steward, G. F. (2010). Nitrogenase genes in non-cyanobacterial plankton: prevalence, diversity and regulation in marine waters. *Aquat. Microb. Ecol.* 61, 235–247. doi: 10.3354/ame01431
- Rippka, R., Deruelles, J., Waterbury, J. B., Herdman, M., and Stanier, R. Y. (1979). Generic assignments, strain histories and properties of pure cultures of cyanobacteria. *Microbiology* 111, 1–61. doi: 10.1099/00221287-111-1-1
- Rodier, M., and Le Borgne, R. (2008). Population dynamics and environmental conditions affecting *Trichodesmium* spp. (filamentous cyanobacteria) blooms in the south-west lagoon of New Caledonia. *J. Exp. Mar. Bio. Ecol.* 358, 20–32. doi: 10.1016/j.jembe.2008.01.016
- Sohm, J. A., Webb, E. A., and Capone, D. G. (2011). Emerging patterns of marine nitrogen fixation. *Nat. Rev. Microbiol.* 9, 499–508. doi: 10.1038/nrmicr02594
- Sohrni, R., and Sempéré, R. (2005). Seasonal variation in total organic carbon in the northeast Atlantic in 2000–2001. *J. Geophys. Res.* 110:C10S90. doi: 10.1029/2004JC002731
- Stedmon, C. A., and Markager, S. (2005). Resolving the variability of dissolved organic matter fluorescence in a temperate estuary and its catchment using PARAFAC analysis. *Limnol. Oceanogr.* 50, 686–697. doi: 10.4319/lo.2005.50.2.0686
- Stedmon, C. A., Markager, S., and Bro, R. (2003). Tracing dissolved organic matter in aquatic environments using a new approach to fluorescence spectroscopy. *Mar. Chem.* 82, 239–254. doi: 10.1016/S0304-4203(03)00072-0
- Steinberg, D. K., Nelson, N. B., Carlson, C. A., and Prusak, A. C. (2004). Production of chromophoric dissolved organic matter (CDOM) in the open ocean by zooplankton and the colonial cyanobacterium *Trichodesmium* spp. *Mar. Ecol. Prog. Ser.* 267, 45–56. doi: 10.3354/meps267045
- Stoecker, D. K., Hansen, P. J., Caron, D. A., and Mitra, A. (2017). Mixotrophy in the marine plankton. *Ann. Rev. Mar. Sci.* 9, 311–335. doi: 10.1146/annurev-marine-010816-060617
- Su, Y., Hu, E., Feng, M., Zhang, Y., Chen, F., and Liu, Z. (2017). Comparison of bacterial growth in response to photodegraded terrestrial chromophoric dissolved organic matter in two lakes. *Sci. Total Environ.* 579, 1203–1214. doi: 10.1016/j.scitotenv.2016.11.104
- Tedetti, M., Cuét, P., Guigue, C., and Goutx, M. (2011). Characterization of dissolved organic matter in a coral reef ecosystem subjected to anthropogenic pressures (La Réunion Island, Indian Ocean) using multi-dimensional fluorescence spectroscopy. *Sci. Total Environ.* 409, 2198–2210. doi: 10.1016/j.scitotenv.2011.01.058
- Tedetti, M., Marie, L., Röttgers, R., Rodier, M., Van Wambeke, F., Helias, S., et al. (2016). Evolution of dissolved and particulate chromophoric materials during the VAHINE mesocosm experiment in the New Caledonian coral lagoon (south-west Pacific). *Biogeosciences* 13, 3283–3303. doi: 10.5194/bg-13-3283-2016
- Thompson, A. W., Foster, R. A., Krupke, A., Carter, B. J., Musat, N., Vault, D., et al. (2012). Unicellular cyanobacterium symbiotic with a single-celled eukaryotic alga. *Science* 337, 1546–1550. doi: 10.1126/science.1222700
- Torréton, J. P., Rochelle-Newall, E., Pringault, O., Jacquet, S., Faure, V., and Briand, E. (2010). Variability of primary and bacterial production in a coral reef lagoon (New Caledonia). *Mar. Pollut. Bull.* 61, 335–348. doi: 10.1016/j.marpolbul.2010.06.019
- Tripp, H. J., Bench, S. R., Turk, K. A., Foster, R. A., Desany, B. A., Niazi, F., et al. (2010). Metabolic streamlining in an open-ocean nitrogen-fixing cyanobacterium. *Nature* 464, 90–94. doi: 10.1038/nature08786

- Turk, K. A., Rees, A. P., Zehr, J. P., Pereira, N., Swift, P., Shelley, R., et al. (2011). Nitrogen fixation and nitrogenase (*nifH*) expression in tropical waters of the eastern North Atlantic. *ISME J.* 5, 1201–1212. doi: 10.1038/ismej.2010.205
- Turk-Kubo, K. A., Frank, I. E., Hogan, M. E., Desnues, A., Bonnet, S., and Zehr, J. P. (2015). Diazotroph community succession during the VAHINE mesocosm experiment (New Caledonia lagoon). *Biogeosciences* 12, 7435–7452. doi: 10.5194/bg-12-7435-2015
- Weishaar, J. L., Aiken, G. R., Bergamaschi, B. A., Fram, M. S., Fujii, R., and Mopper, K. (2003). Evaluation of specific ultraviolet absorbance as an indicator of the chemical composition and reactivity of dissolved organic carbon. *Environ. Sci. Technol.* 37, 4702–4708. doi: 10.1021/es030360x
- Yamashita, Y., Maie, N., Briceño, H., and Jaffé, R. (2010). Optical characterization of dissolved organic matter in tropical rivers of the Guayana Shield, Venezuela. *J. Geophys. Res.* 115:G00F10. doi: 10.1029/2009JG000987
- Yelton, A. P., Acinas, S. G., Sunagawa, S., Bork, P., Pedrós-Alió, C., and Chisholm, S. W. (2016). Global genetic capacity for mixotrophy in marine picocyanobacteria. *ISME J.* 10, 2946–2957. doi: 10.1038/ismej.2016.64
- Zani, S., Mellon, M. T., Collier, J. L., and Zehr, J. P. (2000). Expression of *nifH* genes in natural microbial assemblages in lake George, New York, detected by reverse transcriptase PCR. *Appl. Environ. Microbiol.* 66, 3119–3124.
- Zehr, J. P. (2011). Nitrogen fixation by marine cyanobacteria. *Trends Microbiol.* 19, 162–173. doi: 10.1016/j.tim.2010.12.004
- Zehr, J. P., Montoya, J. P., Jenkins, B. D., Hewson, I., Mondragon, E., Short, C. M., et al. (2007). Experiments linking nitrogenase gene expression to nitrogen fixation in the North Pacific subtropical gyre. *Limnol. Oceanogr.* 52, 169–183. doi: 10.4319/lo.2007.52.1.0169

Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

The reviewer AF-C and handling Editor declared their shared affiliation.

Copyright © 2018 Benavides, Martias, Elifantz, Berman-Frank, Dupouy and Bonnet. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) and the copyright owner are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.