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20,000 years of societal vulnerability and adaptation to climate change in southwest Asia

Matthew D. Jones¹  | Nizar Abu-Jaber² | Ahmad AlShdaifat¹ | Douglas Baird³ | Benjamin I. Cook⁴ | Mark O. Cuthbert⁵  | Jonathan R. Dean⁶  | Morteza Djamali⁷ | Warren Eastwood⁸ | Dominik Fleitmann⁹ | Alan Haywood¹⁰ | Ola Kwiecien¹¹ | Joshua Larsen⁸  | Lisa A. Maher¹² | Sarah E. Metcalfe¹ | Adrian Parker¹³ | Cameron A. Petrie¹⁴  | Nick Primmer¹ | Tobias Richter¹⁵ | Neil Roberts¹⁶ | Joe Roe¹⁷  | Julia C. Tindall¹⁰ | Ezgi Ünal-İmer¹⁸ | Lloyd Weeks¹⁹ 

¹School of Geography, University of Nottingham, Nottingham, UK

²Center for the Study of Natural and Cultural Heritage, German Jordanian University, Amman, Jordan

³Department of Archaeology, Classics and Egyptology, University of Liverpool, Liverpool, UK

⁴NASA Goddard Institute for Space Studies, New York, New York

⁵School of Earth and Ocean Sciences, Cardiff University, Cardiff, UK

⁶School of Environmental Sciences, University of Hull, Hull, UK

⁷Institut Méditerranéen de Biodiversité et d'Ecologie (UMR 7263—CNRS/Aix-Marseille Université/IRD/Université d'Avignon), Aix-en Provence, France

⁸School of Geography, Earth and Environmental Sciences, University of Birmingham, Birmingham, UK

⁹Department of Archaeology and Centre for Past Climate Change, University of Reading, Reading, UK

¹⁰School of Earth and Environment, University of Leeds, Leeds, UK

¹¹Institute for Geology, Mineralogy and Geophysics, Ruhr-Universität Bochum, Bochum, Germany

¹²Department of Anthropology, University of California, Berkeley, California

¹³Department of Social Sciences, Human Origins and Palaeoenvironments Research Group, Oxford Brookes University, Oxford, UK

¹⁴Department of Archaeology and Anthropology, University of Cambridge, Cambridge, UK

¹⁵Center for the Study of Early Agricultural Societies, University of Copenhagen, Copenhagen, Denmark

¹⁶School of Geography, Earth and Environmental Sciences, Plymouth University, Plymouth, UK

¹⁷Institute of Archaeology, University College London, London, UK

¹⁸Department of Geological Engineering, Hacettepe University, Ankara, Turkey

¹⁹School of Humanities, Arts and Social Sciences, University of New England, Armidale, New South Wales, Australia

Correspondence

Matthew D. Jones, School of Geography,
University of Nottingham, Nottingham, UK.
Email: matthew.jones@nottingham.ac.uk

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The Fertile Crescent, its hilly flanks and surrounding drylands has been a critical region for studying how climate has influenced societal change, and this review focuses on the region over the last 20,000 years. The complex social, economic, and environmental landscapes in the region today are not new phenomena and understanding their interactions requires a nuanced, multidisciplinary understanding of the past. This review builds on a history of collaboration between the social and natural palaeoscience disciplines. We provide a multidisciplinary, multiscale perspective on the relevance of past climate, environmental, and archaeological research in assessing present day vulnerabilities and risks for the populations of southwest Asia. We discuss the complexity of palaeoclimatic data interpretation, particularly in relation to hydrology, and provide an overview of key time periods

of palaeoclimatic interest. We discuss the critical role that vegetation plays in the human–climate–environment nexus and discuss the implications of the available palaeoclimate and archaeological data, and their interpretation, for palaeonarratives of the region, both climatically and socially. We also provide an overview of how modelling can improve our understanding of past climate impacts and associated change in risk to societies. We conclude by looking to future work, and identify themes of “scale” and “seasonality” as still requiring further focus. We suggest that by appreciating a given locale's place in the regional hydroscape, be it an archaeological site or palaeoenvironmental archive, more robust links to climate can be made where appropriate and interpretations drawn will demand the resolution of factors acting across multiple scales.

This article is categorized under:

Human Water > Water as Imagined and Represented

Science of Water > Methods

Water and Life > Nature of Freshwater Ecosystems

KEYWORDS

archaeology, Holocene, hydrology, Iran, Levant, palaeoclimate, Turkey

1 | INTRODUCTION

Climate, including drought, has influenced societal change in southwest (SW) Asia, not just in the last decades (e.g., Kelley, Mohtadi, Cane, Seager, & Kushnir, 2015), but for millennia (e.g., Kaniewski, Van Campo, & Weiss, 2012). The Fertile Crescent, its hilly flanks and surrounding drylands have long been a critical region for studying human societal change, first, as being an initial stepping point out of Africa for the first anatomically modern humans (Bae, Douka, & Petraglia, 2017; Hershkovitz et al., 2018), and then as a center for some of the earliest agricultural villages (Barker, 2009; Willcox, Buxo, & Herveux, 2009) and cities (Lawrence, Philip, Hunt, Snape-Kennedy, & Wilkinson, 2016; Ur, 2017). Since these early developments, the region has been the scene of many further social, technological and economic changes and exchanges. While climate has often been discussed as one potential driver for these developments (e.g., Büntgen et al., 2016; H. Weiss, 2016), the modern complex social, economic, and environmental landscapes of the region emphasize the importance of a nuanced, multidisciplinary understanding of past climate change and its relationship to human behaviors (e.g., Jones, Maher, Richter, Macdonald, & Martin, 2016; Ur, 2015). Understanding the vulnerabilities of social and natural systems to change requires high-resolution reconstructions and modelling of the co-evolution of climate and human communities through time, and SW Asia provides a uniquely long record to explore these dynamics. With climate model projections for the region indicating rising temperatures and reduced rainfall in the coming decades (Pachauri et al., 2014), it is important to consider prehistoric and historical datasets regarding the relationships between climate, water availability and people to frame the impact of these changes on populations in the area today.

This study focuses on research published over the last decade since the last comprehensive reviews of the region's palaeoclimate of the late Pleistocene and Holocene by Robinson, Black, Sellwood, and Valdes (2006) and Enzel et al. (2008). Recent progress in modelling capabilities and improvements in data quality across a range of disciplines now allows us to better test hypotheses of human–climate–environment interactions in the past at a variety of scales, both in time and space. With current geopolitical unrest in the region, a renewed focus on topics of migration and conflict, linked directly or indirectly to drought (e.g., Flohr et al., 2017; Gleick, 2014; Kelley et al., 2015), a long-term perspective on these issues is especially timely. In this study we outline our current understanding of climatic changes in SW Asia over the last 20,000 years and integrate this information with the latest archaeological and historical evidence. This combined approach provides a multiscale perspective on the relevance of past climate, environmental, and archaeological research in assessing present day vulnerabilities and risks for human populations in the region. We build on a long history of collaborations between the social and natural palaeoscience disciplines (c.f. Roberts et al., 2018).

The review focuses on the region from present day Turkey, south along the eastern Mediterranean coast, southeast to the Arabian Peninsula and east to the Iranian Plateau (Figure 1). This region marks a meeting point between continents and weather systems that adds to the complexity of reconstructing palaeoclimate and the related, or not, trajectory of past human histories. The last 20,000 years witnessed one of the most dramatic global climatic changes (glacial/interglacial transition), but also includes the full scope of Holocene climatic variability. This time period also includes the development of agriculture and

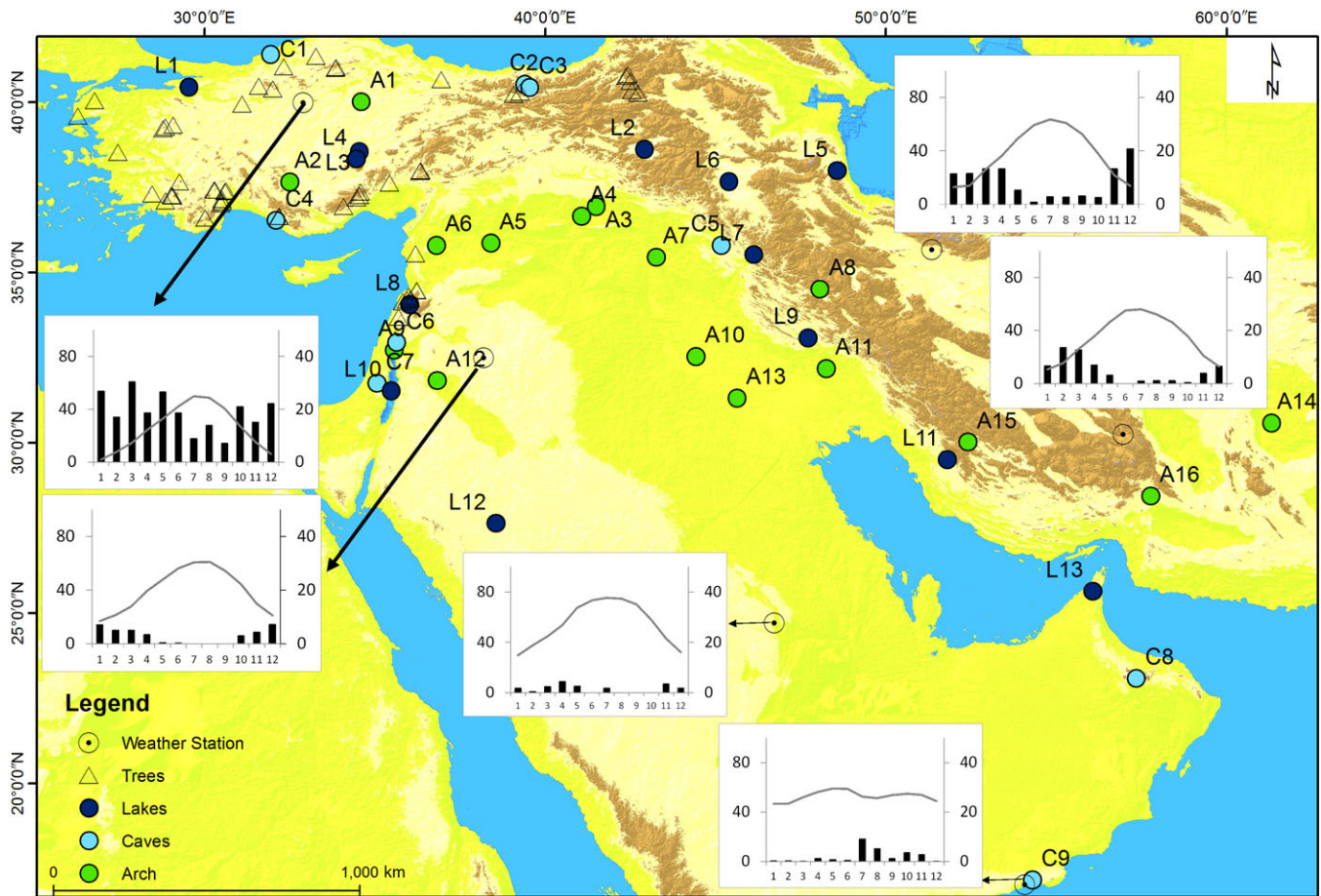


FIGURE 1 Map of the region showing key palaeoenvironmental archives (Table 1; tree rings from Touchan et al., 2014), archaeological sites, and climate seasonality across the region (data from KNMI Climate Explorer). Climate plots show mean monthly (January–December) precipitation (mm; left hand axis) and average air temperature (°C; right hand axis). Archaeological sites shown are A1: Hattusas, A2: Çatal Hoyuk, A3: Tell Leilan, A4: Tell Brak, A5: Abu Hureyra, A6: Ebla, A7: Assur, A8: Godin Tepe, A9: Ohalo II, A10: Babylon, A11: Susa, A12: Azraq, A13: Uruk, A14: Shahr-I Sokta, A15: Tall-e Malyan and A16: Konar Sandal

the first urban societies, the first example of which can be found in this region. There are important records of environmental change through our time period of focus from the seas surrounding our study region (e.g., Heyvaert & Baeteman, 2007; Leroy et al., 2011), including insight into sea-level change (Benjamin et al., 2017; Goldberg, Lau, Mitrovica, & Lamychev, 2016) and how this may have impacted the movement of people, but we do not focus on these directly here. We focus on the environments where people were living, that is, the archaeological sites that they inhabited and terrestrial archives of palaeoenvironmental change, from trees, caves, and lakes and wetlands.

We discuss the complexity of palaeoclimatic data interpretation and its associated uncertainties, particularly in relation to hydrology, before providing an overview of key time periods of palaeoclimatic interest and relating them to archaeological narratives. Vegetation plays a critical role in the human–climate–environment nexus acting as both a control on, and being affected by, natural and anthropogenic environmental change, and we outline some of the outstanding questions regarding vegetation change over the last 20,000 years. We discuss the implications of available palaeoclimate and archaeological data, and their interpretation, on climatic and social palaeonarratives of the region. We also provide an overview of how modelling of climate, palaeoenvironmental archives and people can improve our understanding of past climate impacts and associated change in risk to societies. We end this review by discussing potential future directions for research in order for multiple disciplines to develop more sophisticated, evidence-based hypotheses of the impact of environmental change on societies in the past, integrating complementary, yet distinct, types of palaeoenvironmental and archaeological datasets.

1.1 | Regional climate

As a framework for this review, we very briefly outline here the modern climatology of the region. In general, SW Asia is characterized by wet winters and springs and dry, warm, summers, although the rainfall source areas and circulation patterns that drive these are varied (Enzel, Kushnir, & Quade, 2015) and complicate attempts to reconstruct palaeoclimate beyond local, site-specific, conditions (Stevens et al., 2001). Winter and spring storms from the Atlantic and Mediterranean extend

over the Anatolian Plateau and into northern and west Iran, and across the Levant into northern Arabia. Additional rains, particularly in the south of the region come from the Red Sea, tropical Africa and the Gulf of Oman (Enzel et al., 2015).

Summers are characterized by dry and warm conditions, except in the most southerly parts of the Arabian Peninsula where Africa and Indian Monsoon systems bring precipitation to the mountains of Yemen and Oman. These weather systems also impact the dry conditions of the rest of the region indirectly, due to circulation patterns that bring either descending air, linked to Hadley Cell circulation, or warm and dry north, north-easterly winds across the region (e.g., Lionello et al., 2006).

The balance between winter and summer conditions is the key to water availability and a broad regional picture of local palaeoenvironmental conditions is needed to pull apart any of the subtleties evident in the modern climate regime in the past, or how this might have changed through time.

1.2 | Chronology

Archaeological and palaeoclimate discussions, and particularly comparisons between different records, require robust chronological control, and much has been written on this in general and for our study region (e.g., Blockley & Pinhasi, 2011; Maher, Banning, & Chazan, 2011). In Section 2, we discuss the common dating methods for each of our climate archives of focus. Issues relating to dating archaeological sites are discussed in the later sections (e.g., Section 6). Different disciplines, and dating methods, traditionally use particular chronological conventions (e.g., AD/BC, CE/BCE, BP) such that direct comparisons can be confusing. Here, we use a common chronological notation of ka, thousands of years ago, noting the type of age estimate used where appropriate, more recent events are described in years AD. All radiocarbon age estimates are calibrated unless otherwise stated.

2 | ARCHIVES AND PROXIES: CLIMATE AND HYDROLOGY

Some of the continued uncertainty around proxy interpretation comes from the need to understand the nature of climate experienced by different archives and how these signals are recorded by different proxies within the archives before being incorporated into the geological record. This is an important issue when interrogating any proxy record, but is especially so when considering that record in terms of change that may impact the populations using a given resource, in this case water.

The concept of drought propagation (Box 1) provides a useful starting point for understanding the translation of meteorological anomalies or climate changes through to hydrological changes (in fluxes or hydrochemistry) and then onto the proxy system. This concept has largely been derived from studies of humid catchments and it should be noted that the semiarid to arid climate experienced by much of SW Asia may show important differences; for instance, the sensitivity of the relative partitioning of precipitation into available soil moisture, runoff and groundwater recharge may vary markedly for different types of climate change or anomaly (Van Loon & Van Lanen, 2012). Lakes and speleothems, for example, may be more sensitive to meteorological changes that affect runoff and groundwater recharge generation such as rainfall intensity changes (Cuthbert, Baker, et al., 2014; Cuthbert, Rau, et al., 2014; Markowska et al., 2016). Vegetation, including trees, dependent on soil moisture may have a higher sensitivity to hydrological changes influencing the long-term balance between infiltration and evapotranspiration and are more obviously susceptible to direct human impact, such as through forest clearance or grazing.

Due to the large lag time and attenuation of hydrological change between many groundwater and some surface water bodies (Cuthbert et al., 2017; Cuthbert et al., 2019), certain climate proxy archives may be relatively insensitive to meteorological droughts recorded by others (Box 1). In addition, it is worth noting that modern drought, as defined here, is mostly relevant in the wetter parts of the region, as some areas are always dry with rare rainfall events the exception to mean conditions. Understanding the likely governing hydrological processes and the position of the archive and proxy in the landscape (Figure 2) are, therefore, key to interpreting these differences, and linking proxy records back to climate.

We briefly review here the key terrestrial regional palaeoenvironmental archives and consider their place in the regional hydroscape. Given the strong precipitation gradients across the region (Figure 1), the presence of a tree, speleothem or wetland in the landscape is itself a measure of hydrological state. The presence, or growth, of the archive can therefore be used as a first-order proxy for a threshold in water availability (Vaks et al., 2003).

2.1 | Trees

In low latitudes, including over SW Asia (St. George & Ault, 2014), tree growth is typically more sensitive to moisture availability than temperature. Trees are large-scale integrators of both changes in water supply and demand, they themselves are key parts of the hydrological cycle and are typically better indicators of soil moisture and/or hydrological drought (Box 1). Outside of the humid tropics, trees typically put on a single ring of growth every year, allowing precise annual dating of records using cross-dating, as first described by Douglass (1941). Due to the precisely dated nature of these records, they can

BOX 1

DROUGHT AND ITS PROPAGATION

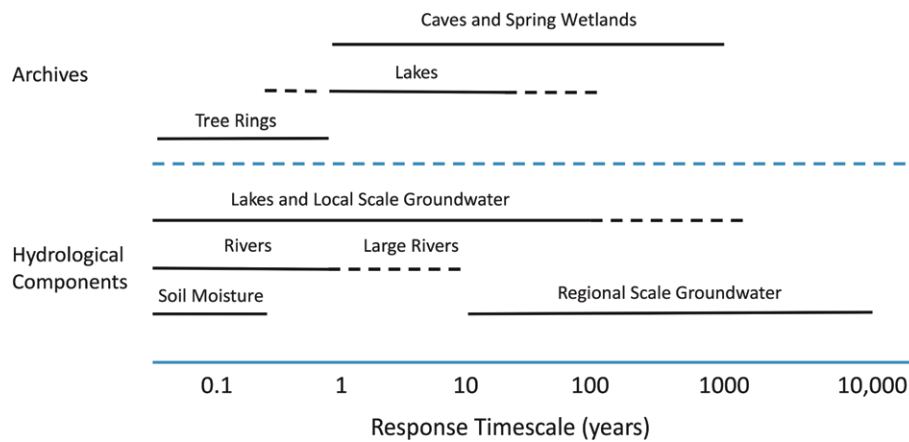
Drought, defined as a deficit of water compared to normal conditions (Van Loon, 2015), is hypothesized to play a significant role in societal behaviour in SW Asia today (Kelley et al., 2015) as well as in the past (e.g., Weiss, 2016). Drought can be defined in multiple ways (Van Loon, 2015), and these different definitions can help us to conceptualize the impacts of hydrological change, broadly, on human populations at different points through time (Rohling, 2016).

Meteorological drought: A precipitation deficiency, possibly combined with increased evaporation.

Soil-moisture (agricultural) drought: Soil moisture deficit, often linked to crop failure.

Hydrological drought: Lower than average availability of surface and subsurface water.

Meteorological drought is filtered by catchments and their biophysical characteristics such as geology, soils and vegetation, so that agricultural and hydrological drought, those more likely to be captured by geological archives and impact human populations, can be damped relative to, or lag behind, the climatic forcing.



Typical response times to a climatic perturbation (e.g., meteorological drought) of critical components of the hydrological cycle for SW Asia, and the geological archives that are the focus of this study. A simple conceptual framework is complicated by the large scale of some SW Asian catchments, such as the Tigris and Euphrates Rivers, where response time may be further delayed due to the significant potential distance from the source of the perturbation, as well as smoothing of the signal in longer residence time systems.

In modern hydrology, droughts are generally at the annual scale or greater, but less than decadal, and therefore reflect movement away from the longer-term mean which is itself derived from approximately three decades of data (Van Loon, 2015). There are few archives, beyond tree rings and varved lake records, which could record such drought at this resolution in the past. Care must be taken, therefore, in discussing drought using relatively low-resolution proxy records which may more likely record changes in mean climate state, and thus lead to different societal impacts (Ur, 2015).

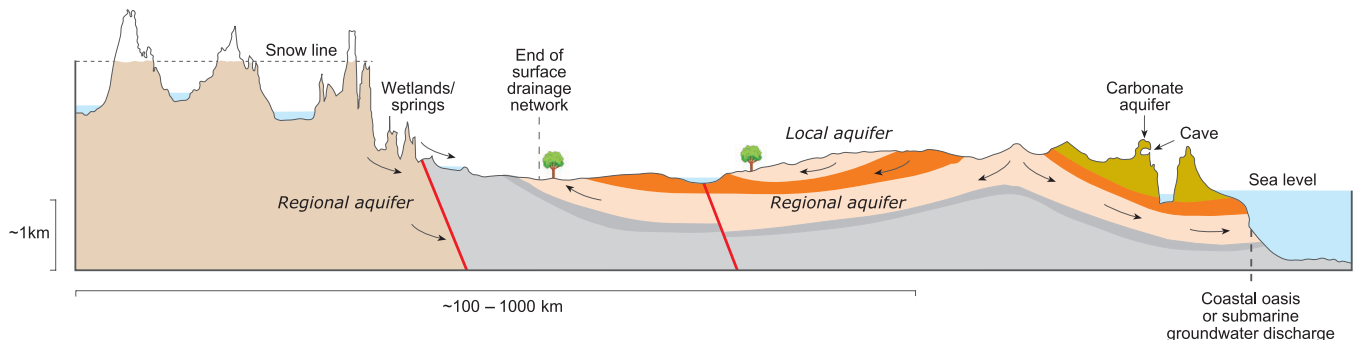


FIGURE 2 Typical Southwest Asia hydroscape, highlighting key parts of the hydrological cycle and location of key terrestrial archives within this system

be calibrated using instrumental climate observations to develop quantitative reconstructions of specific climate variables, including the Palmer Drought Severity Index (B. I. Cook, Anchukaitis, Touchan, Meko, & Cook, 2016; B. I. Cook, Ault, & Smerdon, 2015), precipitation (Grissino-Mayer, 1996), streamflow (Woodhouse, 2001), and temperature (Büntgen et al., 2016). This calibration in turn allows for quantitative comparisons between tree-ring-based palaeoclimate reconstructions and the instrumental climate record (Griffin & Anchukaitis, 2014) and climate model simulations of the past (e.g., B. I. Cook et al., 2015; E. R. Cook, Seager, et al., 2015). Furthermore, because trees are fairly ubiquitous across broad geographical regions, individual sampling sites can be combined into networks to generate gridded spatial reconstructions. Such reconstructions can be especially informative for understanding climate dynamics because many modes of natural climate variability have distinct spatial fingerprints that manifest in these reconstructions (Herweijer, Seager, Cook, & Emile-Geay, 2007).

One of the biggest disadvantages to tree rings as palaeoclimate proxies is the relatively short lifespan of individual trees, typically hundreds to, occasionally, several thousands of years. This limits the length of most well replicated tree-ring reconstructions to the last 2,000 years. An additional side effect of the relatively short lifespan is the difficulty of preserving centennial scale (or longer) variability in tree-ring-based reconstructions, a phenomenon often referred to as the “segment length curse” (E. R. Cook, Briffa, Meko, Graybill, & Funkhouser, 1995). However, considerable strides have been made to ameliorate this problem, beginning with use of “regional curve standardization” (Briffa et al., 1992) and followed up by the development and use of “signal-free” detrending methods (Melvin & Briffa, 2008). In doing so it is possible to preserve multicentennial variations in climate in large tree-ring datasets composed of overlapping tree-ring series from living trees and remnant wood samples.

Over SW Asia specifically, the availability of tree-ring chronologies is limited by the extreme aridity in much of the region (precluding growth of trees) and the long history of human settlement and occupation (which can make it difficult to find old, undisturbed trees). This is highlighted in Figure 1 (tree ring data from Touchan et al., 2014), which shows that tree-ring chronologies in SW Asia are largely found in Turkey, Syria, Jordan, and Lebanon, with recent work also identifying chronologies in Cyprus and Iran (Griggs, Pearson, Manning, & Lorentzen, 2014; Nadi, Bazrafshan, Pourtahmasi, & Bräuning, 2017). Among these, the longest chronologies (>200 years) are confined to Turkey, Jordan, and Lebanon.

2.2 | Caves

Caves are abundant throughout the region and a number of speleothem-based palaeoclimate reconstructions have been developed for Turkey, Lebanon, Israel, Iraq, and Oman (Table 1). The chronologies of all stalagmite records are primarily based on uranium-series dates, which are sometimes supported by annual layer counts (e.g., Fleitmann et al., 2004; Flohr et al., 2017). Typical chronological uncertainties of stalagmite records vary between 0.5 and 2% of the absolute age, depending on the uranium content and purity of the calcite.

Oxygen and carbon stable isotope ratios ($\delta^{18}\text{O}$ and $\delta^{13}\text{C}$, respectively, using the common delta notation) of stalagmite calcite are the most frequently used hydroclimate proxies. Stalagmite $\delta^{18}\text{O}$ values are primarily influenced by the $\delta^{18}\text{O}$ of cave drip water, usually a function of surface precipitation, where $\delta^{18}\text{O}$ of precipitation is influenced by multiple climate parameters such as temperature, and the origin and amount of rainfall (see below). The interpretation of $\delta^{18}\text{O}$ stalagmite calcite values is therefore not straightforward, and may be further complicated by other factors, as shown by recent work in other semiarid environments, that demonstrate that speleothems respond to complex recharge, as well as in-cave, processes (e.g., Baker et al., 2018; Cuthbert, Baker, et al., 2014; Cuthbert, Rau, et al., 2014; Markowska et al., 2016). The majority of stalagmite records from the region have been interpreted as reflecting changes in the amount of precipitation (e.g., Bar-Matthews, Ayalon, Gilmore, Matthews, & Hawkesworth, 2003; Cheng et al., 2015; Fleitmann et al., 2003; Flohr et al., 2017) and changes in the source of moisture (e.g., Fleitmann et al., 2007; Ünal-İmer et al., 2015).

Like $\delta^{18}\text{O}$, stalagmite $\delta^{13}\text{C}$ values are influenced by several environmental factors that include changes in surface vegetation (including vegetation density, proportion of C_3 to C_4 photosynthetic plants), soil microbial activity, recharge conditions (open vs. closed system recharge), and kinetic fractionation processes during calcite precipitation in the cave (drip rates and cave air $p\text{CO}_2$). Almost all factors influencing $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ values in stalagmites are therefore influenced, directly or indirectly, by temperature and precipitation.

Trace elements are an additional climate proxy increasingly measured at high resolution from speleothems, where magnesium, barium, phosphorous, and uranium concentrations appear to be additional proxies for the amount of precipitation (Flohr et al., 2017). However, their full potential as an additional hydrological proxy in stalagmites from across the region is not yet fully exploited.

2.3 | Lakes and wetlands

There is a relatively long history of palaeoenvironment research from lake archives from across the region (Table 1 and see summaries in Roberts et al., 2018; Roberts & Reed, 2009). A number of proxies have been regularly used including pollen

TABLE 1 Key Lake and Cave sites from SW Asia (Figure 1)

Site		Elevation (masl)	Duration (ka)	Selected references
Lakes and wetlands				
Iznik	L1	85	36.2–present day	Roeser et al. (2012); Ülgen et al. (2012)
Van	L2	1,648	600–present day	Çağatay et al. (2014); Cukur et al. (2014); Kuzucuoğlu et al. (2010); Kwiecien et al. (2014); Litt and Anselmetti (2014); Stockhecke et al. (2014); Wick, Lemcke, and Sturm (2003)
Eski Acıgöl	L3	1,270	17–present day	M. D. Jones, Roberts, and Leng (2007); N. Roberts et al. (2001); Turner, Roberts, and Jones (2008); Woldring and Bottema (2001)
Nar	L4	1,363	13.8–present day	Dean et al. (2013, 2015); England, Eastwood, Roberts, Turner, and Haldon (2008); M. D. Jones, Roberts, Leng, and Turkes (2006); N. Roberts et al. (2016); Woodbridge and Roberts (2011)
Neor	L5	2,500	13–present day	Ponel et al. (2013); Sharifi et al. (2015)
Urmia	L6	1,267	200–present day	Djamali et al. (2008); Stevens, Djamali, Andrieu-Ponel, and de Beaulieu (2012)
Zeribar	L7	1,300	42.6–present day	Stevens, Wright, and Ito (2001); Wasylikowa and Witkowski (2008); Wasylikowa et al. (2006)
Yammouneh	L8	1,360	400–present day	Develle et al. (2011) Develle, Herreros, Vidal, Surssock, and Gasse (2010); Gasse et al. (2015)
Mirabad	L9	800	9.3–present day	Griffiths, Schwalb, and Stevens (2001); Stevens, Ito, Schwalb, and Wright (2006)
Dead Sea	L10	418	220–present day	Litt, Ohlwein, Neumann, Hense, and Stein (2012); Migowski, Stein, Prasad, Negendank, and Agnon (2006); Neugebauer et al. (2014); Torfstein, Goldstein, Stein, and Enzel (2013)
Parishan	L11	823	3.9–present day	Djamali et al. (2016); M. D. Jones et al. (2015)
Tayma	L12	801	10–present day	Engel et al. (2012)
Awafi	L13	8	8.1–3	Parker et al. (2004, 2006, 2016)
Speleothems				
Sofular	C1	440	50–present day	Badertscher et al. (2014); Fleitmann et al. (2009); Göktürk et al. (2011)
Karaca	C2	1,536	77–6	Rowe et al. (2012)
Akcakale	C3	1,530	0.5–present day	Jex et al. (2011)
Dim	C4	232	90–10	Ünal-İmer et al. (2015); Ünal-İmer, Shulmeister, Zhao, Uysal, and Feng (2016)
Gejkar	C5	650	2.5–present day	Flohr et al. (2017)
Jeita	C6	100	20.3–0.4	Cheng et al. (2015); Verheyden, Nader, Cheng, Edwards, and Swennen (2008)
Soreq	C7	400	185–1	Bar-Matthews and Ayalon (2011); Bar-Matthews, Ayalon, Kaufman, and Wasserburg (1999); A. Matthews, Ayalon, and Bar-Matthews (2000); Orland et al. (2009)
Hoti	C8	800	330–present day	Burns, Matter, Frank, and Mangini (1998); Fleitmann et al. (2003, 2004); Neff et al. (2001)
Qunf	C9	650	10.3–0.4	Fleitmann et al. (2007)

(see Section 4), oxygen isotopes (Roberts et al., 2008), and diatoms (Vossel, Roeser, Litt, & Reed, 2018), with sediment chemistry, at increasingly high resolution given improvements in nondestructive scanner technologies, also becoming more common place (Sharifi et al., 2015). All these proxies have been used as records of changing water availability, with some debate surrounding these interpretations (Jones & Roberts, 2008). Radiocarbon dating or U-series methods (where old carbon impacts the radiocarbon chronology) are commonly used to establish the age of lake archives (e.g., Dean, Jones, et al., 2015) with individual age-estimate errors typically in the order of tens to hundreds of years. Recent advances in age-depth modelling using Bayesian techniques (Blaauw & Christen, 2011) have great potential in reducing these errors and calibration uncertainties. Tephrochronology, where applicable, is another way to provide further dating control (Eastwood, Tibby, Roberts, Birks, & Lamb, 2002). At some sites lake sediments are annually laminated (varved) opening up the possibility of annual, or seasonal, records of past change (Zolitschka, Francus, Ojala, & Schimmelmman, 2015).

There has been much recent interest in identifying former lakes and wetlands in landscapes where they are either currently absent or greatly diminished, and in the interpretation of their sedimentary archives (Pigati, Rech, Quade, & Bright, 2014) particularly in the Arabian Peninsula (Engel et al., 2017; Enzel et al., 2015; Enzel, Quade, & Kushnir, 2017) and Jordan (Catlett et al., 2017; Jones et al., 2016). These studies follow a long history of research into lake levels in the more humid parts of the region, such as Konya (Roberts, 1983), Lake Van (Kuzucuoğlu et al., 2010), and the Dead Sea/Lake Lisan (Torfstein et al., 2013).

Irrespective of the uncertainty surrounding the nature of some water bodies, there is clear evidence for a greater presence of open water in the past. This is usually explained by some combination of enhanced precipitation, enhanced groundwater recharge, decreased open water evaporation, and enhanced local and/or regional groundwater discharge. In addition to the effects of changing precipitation and evaporation, the increase in human population over the Holocene will also have resulted in increased consumption of water, potentially impacting lakes and wetlands (Jones et al., 2015).

Evidence that is currently available points to localized wetland development in the past, particularly in more arid parts of the region (Jones, Maher, Richter, et al., 2016). If recharge was regionally enhanced, then this would also have created a regionally raised water table, capable of forming water bodies or wetlands wherever the land was below that level. If a rise in regional unconfined water-tables is not supported by empirical evidence elsewhere in the region, then the more localized development of wetlands (in the absence of surface drainage networks) is likely to be related to fault-mediated discharges from confined aquifers or break in slope supply (i.e., springs). This poses an interesting problem for linking the timing of wetter land surfaces with wetter climates, since depending on the length of the groundwater systems involved (Box 1), the lags between the two could be thousands to tens of thousands of years (Cuthbert et al., 2017; Cuthbert et al., 2019). The Deep Sandstone Aquifer complex in Jordan, for example, consists of palaeowater that, while difficult to date, is largely a remnant of Late Pleistocene to Early Holocene climate (Abu-Jaber & El-Naser, 2016). Spring calcretes, such as those described in the Late Pleistocene of Wadi Sabra (Bertrams et al., 2012) and the Neolithic of Beidha (Rambeau et al., 2011), therefore, may mark paleosprings reflecting either local recharge at the times of deposition or an older, nonlocal, recharge event that took considerable time to propagate to the surface.

3 | SOME KEY TIME PERIODS OF CLIMATIC INTEREST

Here, we briefly summarize some of the key periods of interest in terms of regional environmental and climatic change, and particularly those that have been linked to substantial key societal changes in the region. We focus on key debates surrounding changing environments at different frequencies.

3.1 | Millennial scale changes

There is continued interest in the transition from the last glacial into the Holocene and the potential role this shift played in the transition to agriculture by Neolithic people (Box 2). Continuous records that span the full transition from 20,000 years BP into the early Holocene remain relatively scarce, but those available point to a gradient of conditions across the region. Lake levels of Lake İznik in northwest (NW) Turkey (Roeser et al., 2012) and Lake Van in eastern Turkey (Tomonaga et al., 2017)

BOX 2

THE YOUNGER DRYAS

The Younger Dryas, or more correctly the regional temporal and climatic manifestation of the late glacial stadial, has long been discussed as a potential trigger for the transition from hunting and gathering to agriculture through the last glacial–interglacial transition in SW Asia (e.g., Bar-Yosef, 2009; Bar-Yosef & Belfer-Cohen, 2002; Hillman, Hedges, Moore, Colledge, & Pettitt, 2001; Moore & Hillman, 1992), in part due to apparent correlations between societal and climatic changes (Balter, 2010). In general, the period is considered to have been drier across the region, from speleothem (e.g., Bar-Matthews et al., 1999; Verheyden et al., 2008) and pollen evidence (e.g., Baruch & Bottema, 1999; Kadosh, Sivan, Kutiel, & Weinstein-Evron, 2004; Van Zeist & Wright, 1963), but actual records for this time period remain relatively scarce, and chronological uncertainties in the palaeoclimate records remain (Meadows, 2005). Some recent work suggests that the Younger Dryas was cool, but not as dry as previously thought (Hartman, Bar-Yosef, Brittingham, Grosman, & Munro, 2016). Broadly speaking, the Younger Dryas drying has been described as a significant stress factor that influenced both the subsistence economy as well as the settlement pattern of late Pleistocene groups in the Levant (references above).

In recent years, archaeologists have begun to revise the scenarios for the impact of the Younger Dryas (Rosen & Rivera-Collazo, 2012). In the southern Levant, the chronological correlation between the Younger Dryas and the emergence of the late Natufian cultural assemblage of the Epipalaeolithic (Figure 6) is now largely in doubt (e.g., Caracuta et al., 2016; Grosman, 2013; Maher et al., 2011) and further north evidence for plant cultivation at Abu Hureyra during the late Natufian is now disputed (Colledge & Conolly, 2010). Evidence has also begun to emerge for substantial early and late Natufian settlement outside the original core zone (e.g., Richter, 2017; Richter, Arranz-Otaegui, Yeomans, & Boaretto, 2017; Rodríguez et al., 2013) and for more continuity in sedentary settlement in the Mediterranean zone (Grosman et al., 2016). Importantly, improvements in dating archaeological sites, also suggest that the Pre-Pottery Neolithic A now began within the Younger Dryas (e.g., Blockley & Pinhasi, 2011; Wicks et al., 2016). Thus, recent work has reduced the apparent importance of the Younger Dryas as a trigger event for the onset of the first crop-based agricultural societies.

were relatively low at the last glacial maximum (LGM), for example, although pore-water salinity from the latter only provides a multimillennial resolution record, and there are high terraces around Lake Van radiocarbon dated to ~25 and ~21 ka (Kuzucuoğlu et al., 2010). The advance of glaciers (Sarikaya & Çiner, 2015) and continuous speleothem growth in Dim Cave in SW Turkey (Ünal-İmer et al., 2015; Figure 3), as well as high levels of Lake Lisan in the Dead Sea Basin (Torfstein et al., 2013; Figure 3), and continuous deposition of speleothems in the vicinity of the Dead Sea (Frumkin, Bar-Yosef, & Schwarcz, 2011), suggest relatively wetter conditions. This spatial pattern is complicated by the timing of shifts in condition, which suggest relatively wetter conditions up to the LGM (26–21 ka), but very dry conditions after it (18–15 ka), the youngest dates for highstands at Lake Konya, for example, are radiocarbon dated at ca. 21 ka (Roberts, 1983).

Lake $\delta^{18}\text{O}$ records from across the eastern Mediterranean show a transition from the late glacial to the early Holocene from positive to negative, generally interpreted to indicate wetter conditions (Roberts et al., 2008). The transition occurs relatively abruptly in some records; for example, in less than 200 years (with over half the shift in oxygen isotopes occurring in just 9 years) at Lake Nar in central Turkey (Dean, Jones, et al., 2015; Figure 3). There are also differences in the cave $\delta^{18}\text{O}$ records through the transition, with a positive shift recorded in Sofular Cave in NW Turkey and a negative shift recorded further south in Dim and Soreq caves (Figure 1). These differences probably reflect differing controls on the isotopic composition of rainfall at these sites, the balance between changing conditions in the source waters in the Black and Mediterranean Seas, respectively, and the changing amount of local rainfall (Bar-Matthews et al., 2003; Fleitmann et al., 2009). The Hoti and Qunf Cave $\delta^{18}\text{O}$ records (Figure 3) from Oman both reflect the strength of the Indian Ocean monsoon (Fleitmann et al., 2007). These latter examples underline that varying factors can influence environmental signals recorded by the same proxy ($\delta^{18}\text{O}$) in the same type of archive (caves) across the region in different ways, most notably here due to the different sources of rainfall.

A wetter early Holocene continued until ~7 ka, when a transition to drier conditions began in the eastern Mediterranean (Clarke et al., 2016; Roberts, Brayshaw, Kuzucuoğlu, Perez, & Sadori, 2011), with a similar trend of decreasing monsoonal

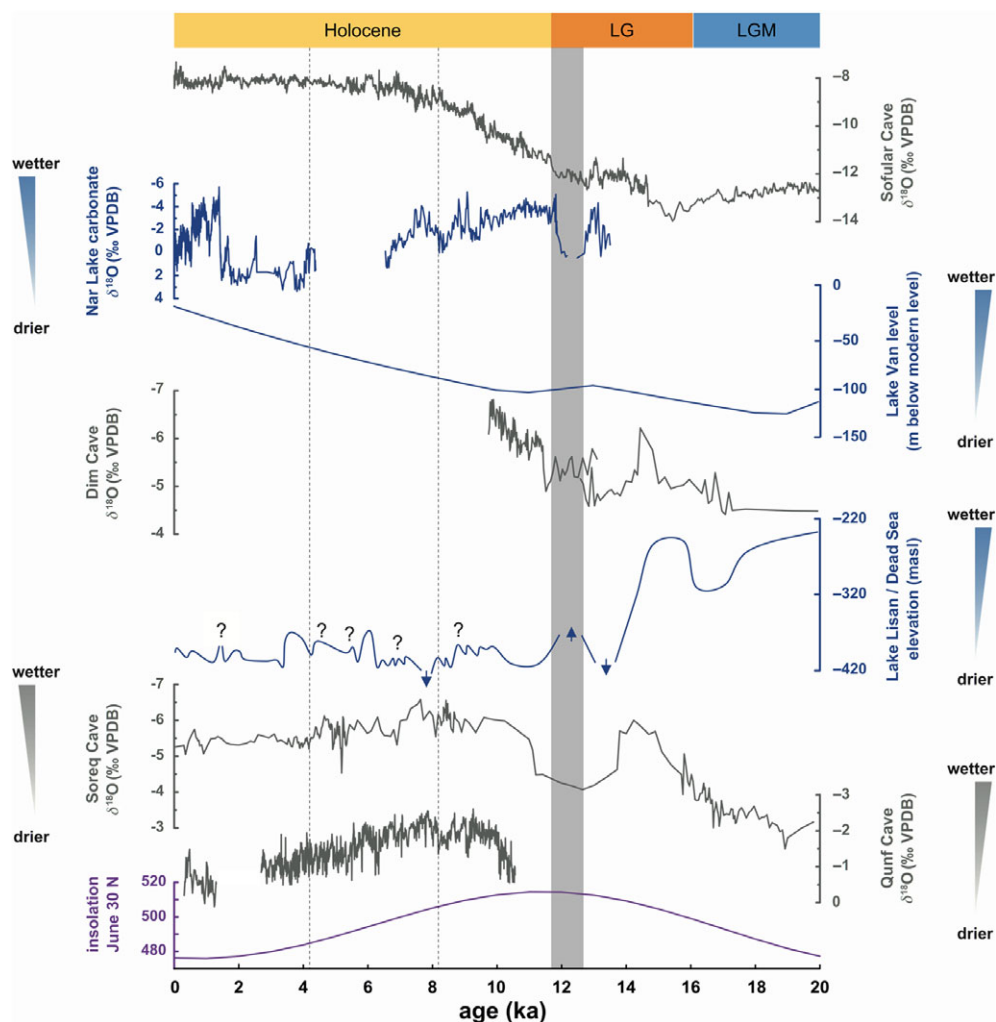


FIGURE 3 Summary of regional palaeoenvironmental change for the last 20 ka. Key time periods (LG, late glacial; LGM, last glacial maximum) and climatic events discussed in the text are highlighted (late glacial stadial, grey shading; 8.2 and 4.2 ka, dashed lines). Note that each record is plotted based on its own chronology. Insolation data are from A. Berger and Loutre (1991), see Table 1 for site references

rains documented in southern Arabian speleothems (Fleitmann et al., 2003, 2007). The end of the early Holocene “humid” period ~5.9 ka in SE Arabia (Parker et al., 2006; Fleitmann et al., 2007) corresponds with the end of the Neolithic in this region (Preston et al., 2015) and with evidence for violence and conflict, possibly over water resources, coeval with increasing aridity (Kutterer & Uerpmann, 2012; Uerpmann, Uerpmann, & Jasim, 2008).

In general, these millennial scale trends in the region follow summer insolation trends (Figure 3). A reduction in insolation through the Holocene led to a southward shift in the boreal summer Inter Tropical Convergence Zone, and weaker summer monsoonal rains in southern Arabia, while in the eastern Mediterranean region this drying was caused by a northerly shift in westerly storm tracks that reduced winter precipitation (Dean, Jones, et al., 2015).

3.2 | Abrupt events

In the Holocene, where more continuous, higher resolution records are available, the longer term millennial trends in climate are punctuated by centennial- and multi-decadal-scale periods that are drier than the millennial average. The recognition of these brief events in any archive is partly sample-resolution dependent, but even some relatively high-resolution records such as the Sofular Cave record in NW Turkey (Göktürk et al., 2011) do not document all of the events recorded elsewhere during the Holocene.

A drying around 9.3 ka, linked to an event in the Greenland ice cores (Rasmussen, Vinther, Clausen, & Andersen, 2007), has so far only been observed in the Lake Nar and Qunf Cave records (Dean, Jones, et al., 2015; Fleitmann et al., 2008) as well as potentially being a factor in the slowdown of speleothem growth at Dim Cave (Ünal-İmer et al., 2015). In the Nar, Qunf and Hoti Cave records, the 8.2 ka event, again recognized in the Greenland ice cores and across much of northern Europe (e.g., Alley & Agustsdottir, 2005; Daley et al., 2011), lasts longer than the 9.3 ka event, perhaps explaining why it is documented as a drier period in many more records from across SW Asia. It is expressed as an isotopic shift to more positive values in Soreq Cave, Israel (Bar-Matthews et al., 2003) and Hoti and Qunf caves, Oman (Fleitmann et al., 2007), an interruption of sapropel S1 in the eastern Mediterranean (Kotthoff et al., 2008), and increased sediment flux rates from SE Arabia (Parker et al., 2016). It has also been detected in the Riwasa playa lake on the plains of NW India (Dixit, Hodell, Sinha, & Petrie, 2014).

Assessing and interpreting any cultural and societal impact of the 8.2 ka event is controversial. It has been proposed that it prompted migration and settlement abandonment in Turkey and contributed to the Neolithization of the Aegean and SW Europe (Weninger et al., 2006, 2014). However, there is much evidence for continuity in settlement over the centuries around 8.2 ka on the Anatolia plateau (Baird, 2012b) and Neolithic communities are present in western Turkey several centuries before the 8.2 ka event. Recent evaluation of archaeological evidence suggests that the 8.2 ka event had no systematic regional scale impact on societies (Flohr, Fleitmann, Matthews, Matthews, & Black, 2016; Maher et al., 2011), although local impacts can be detected (Roffet-Salque et al., 2018).

A dry period ~5.2 ka is recorded by speleothems from Israel (Bar-Matthews & Ayalon, 2011) and lakes/wetlands in SE Arabia (Parker et al., 2006) and Turkey (Kuzucuoğlu et al., 2011) and has been linked by some to the end of the late Uruk period societies in Mesopotamia (H. Weiss, 2003). Another period of drought at ~4.2 ka (Kaniewski, Marriner, Cheddadi, Guiot, & Van Campo, 2018) is recorded in lakes in Turkey (Dean, Jones, et al., 2015; Eastwood, Leng, Roberts, & Davis, 2007), SE Arabia (Parker et al., 2006), the Dead Sea (Litt et al., 2012), and in the Gulf of Oman (Cullen et al., 2000). This drying period, which has been identified in South Asia as a weakening of the monsoon (Berkelhammer et al., 2012; Dixit, Hodell, & Petrie, 2014; Prasad et al., 2014), has been linked as contributing to the “collapse” of the Akkadian civilization in Mesopotamia, the possible disintegration of urban communities in the southern Levant (Staubwasser & Weiss, 2006; Weiss, 2016), the decline in known settlement in SE Arabia at the end of the early Bronze Age (Preston et al., 2015) and significant changes to settlement systems and irrigation technologies in SE Iran (Fouache, Francfort, Cosandey, & Adle, 2015). This event has come under recent focus with the subdivision of the Holocene (Cohen, Finney, Gibbard, & Fan, 2013; Walker et al., 2018) and we discuss more of the detail relating to the impact on societies of the 4.2 ka event in Section 5.

The final significant dry period was centered at ~3.2–3.1 ka and has been described from Turkey (Roberts et al., 2001; Wright, Fairbairn, Faith, & Matsumura, 2015), Lake Zeribar in Iran (Stevens et al., 2001) and Jeita Cave in Lebanon (Verheyden et al., 2008). Similar to the ~5.2 and 4.2 ka events, high resolution analysis suggests that this drier period comprised several drought episodes interspersed within decades of wetter climate (Kuzucuoğlu, 2009). Hittite texts referred to drought (Kuzucuoğlu, 2015) and their civilization declined at the end of the Bronze Age, with their capital Hattuşa destroyed ~3.18 ka (Weiss, 1982). There is also evidence of crop failures in Syria (Kaniewski et al., 2010) and evidence of large-scale migration/displacement of “Sea Peoples” throughout the eastern Mediterranean including Anatolia and the Levantine littoral (Van De Mieroop, 2008) at this time. However, as with previous climatic events, which as noted above are not seen in all regional proxy records, not all societies were impacted in the same way. Physical anthropological studies at Tell Dothan, in

the West Bank, suggest continuation of a sedentary lifestyle of agronomists through this period (Gregoricka & Sheridan, 2017).

The ~9.3, 8.2, 4.2, and 3.2 ka events are approximately coeval with Bond events, periods of increased ice rafting in the North Atlantic (Bond et al., 1997). The 5.2 ka event seen in SW Asia, however, does not appear to have an equivalent in the North Atlantic. Much of the precipitation that falls in the eastern Mediterranean originates from the North Atlantic, so it is unsurprising that the majority of these dry “events” coincide with cool sea surface temperature events in the North Atlantic which led to a reduction in cyclogenesis (Bartov, Goldstein, Stein, & Enzel, 2003). Nonetheless, the fact that some of these dry “events” last longer, in the Nar and Qunf records, for example, than the more discrete “events” of the North Atlantic, suggests there may have been other underlying causes, perhaps related to changes in solar activity via the winter Siberian High (Rohling & Palike, 2005). It is also important to consider the impact of the Indian summer monsoon (ISM), and its periods of fluctuation during the Holocene (Jones et al., 2006). As noted above, the 4.2 ka event has been linked to weakening monsoon in NW India, and also in the Himalayas and Central India. Biophysical nonlinear feedback mechanisms have been suggested as a possible factor prolonging phases of increased aridity in SW Asia (Parker et al., 2016).

3.3 | The last 2,000 years

There are a number of high-resolution (annual or near-annual) records from, or attributed to, the region covering the last 2,000 years (Figure 4). In the eastern Mediterranean, the medieval climate anomaly (MCA) was generally wetter and the little ice age (LIA) drier (Luterbacher et al., 2012; Roberts et al., 2012). A recent multiproxy speleothem record from northern Iraq suggests increasingly dry conditions over the last 1,000 years in that part of SW Asia (Flohr et al., 2017).

One of the most recent and comprehensive reconstructions of hydroclimate over SW Asia and adjacent regions through the last 2 ka is the Old World Drought Atlas (OWDA; E. R. Cook, Seager, et al., 2015). The OWDA is an annually resolved, tree-ring-based reconstruction of the summer season (June–July–August) Palmer Drought Severity Index (PDSI) at half degree spatial resolution across Europe and the Mediterranean. PDSI is a normalized indicator of soil moisture availability (negative values indicating drought), integrating changes in moisture supply (precipitation) and demand (evapotranspiration) on time-scales of typically 12–18 months. Spatial coverage of the OWDA extends into SW Asia, though the proxy coverage there is poor relative to other areas, limiting robust analyses over this region to 0.9 ka and later (Cook et al., 2016).

One notable feature of SW Asia in the OWDA is the tendency towards anti-phased hydroclimate variability between the southern Levant (including Israel, Jordan, southern Syria, Lebanon) and a region encompassing Greece, Anatolia, and the Black Sea (Figure 5; Cook et al., 2016). This feature is persistent across centuries, with strong and significant antiphase coherency from interannual to multidecadal timescales. The spatial fingerprint of this dipole bears a strong resemblance to the pattern that would be expected from variability in the North Atlantic Oscillation (Cook et al., 2016). Such a dipole has been described previously (Xoplaki, Gonzalez-Rouco, Luterbacher, & Wanner, 2004), and its persistent existence in the OWDA suggests that it is a robust feature of natural hydroclimate variability in the region.

The OWDA has also shed some light on the early 21st century AD drought in the Levant region that may have contributed, in part, to the recent social unrest in Syria (Gleick, 2014; Kelley et al., 2015). Over the Levant, this protracted period of drought (1998–2012 in the OWDA) is the most intense 15-year drought of the last 900 years, and also includes the single most severe individual drought year in the record; 2000 AD (Cook et al., 2016). The OWDA, therefore, provides some independent support for analyses that suggest this recent drought was exceptional relative to natural variability, a sign that the expected drying from anthropogenic climate change in the region may already be beginning to occur (Kelley et al., 2015).

4 | BIOGEOGRAPHY AND VEGETATION CHANGE

SW Asia encompasses four major biogeographical regions: (a) the Saharo-Sindian region in the south, comprising desert vegetation and pseudo-savannas in the Arabian Peninsula and southern Mesopotamia and tropical arid vegetation in the southern Arabian Peninsula and southern Iran; (b) the Irano-Turanian region in the center and east, comprising *Artemisia* steppes in internal Iranian plateaus and Mesopotamia with open woodlands and steppe forests of deciduous trees and junipers in the Irano-Anatolian mountains; (c) the Mediterranean region in the west comprising typical Mediterranean vegetation in western and southern Anatolia and the Levant; and (iv) the Euro-Siberian region in the north and northwest including mesic deciduous and mixed conifer-deciduous forest in the South Caspian and Black Sea regions (White & Léonard, 1990). Halophytic and hygro-halophytic vegetation also occurs locally in endorheic depressions (Zohary, 1973). Continentality, winter temperatures and precipitation seasonality are the most important parameters determining the boundaries between these regions (Djamali, Brewer, Breckle, & Jackson, 2012).

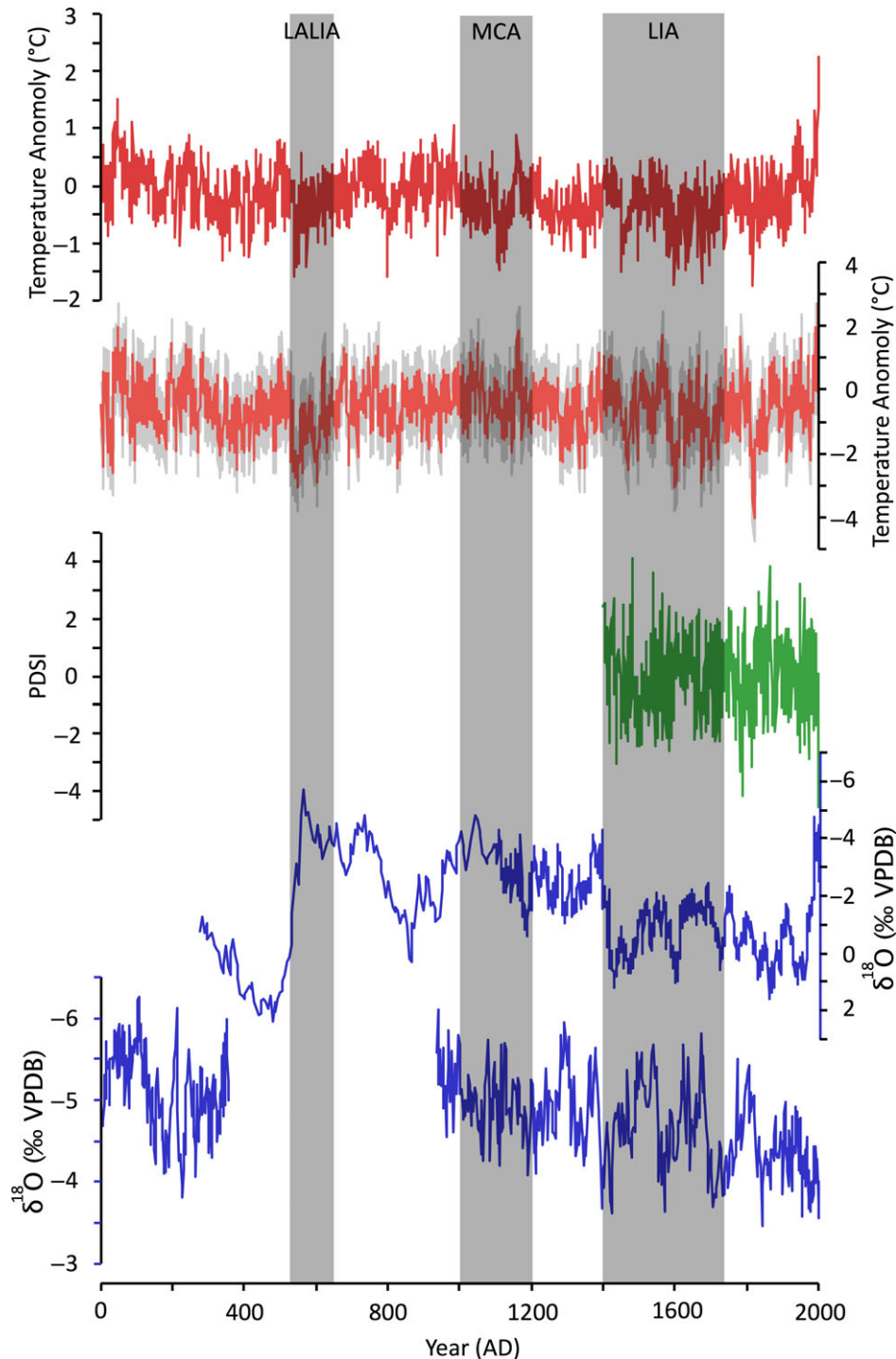


FIGURE 4 High resolution records of the last 2000 years. From the top: European temperature anomalies from the PAGES2k consortium (Ahmed et al., 2013), European summer temperature anomalies (Büntgen et al., 2011), Palmer Drought Severity Index (PDSI) reconstruction for southwest Asia (Cook et al., 2016) and oxygen isotope records from Lake Nar (Jones et al., 2006), and Gejker Cave (Flohr et al., 2017). Grey shading picks out the Late Antiquity Little Ice Age (LALIA), Medieval Climate Anomaly (MCA), and Little Ice Age (LIA) as presented in Büntgen et al. (2016)

4.1 | Last glacial maximum

The relatively sparse pollen data that extend back to the LGM indicate the dominance of an extremely dry steppe in most of the continental interior of SW Asia (Djamali et al., 2008; Litt, Pickarski, Heumann, Stockhecke, & Tzedakis, 2014; van Zeist & Bottema, 1977). Modern analogues for this cold, dry steppic vegetation can be found in the subalpine zone of the Irano-Turanian mountains (Djamali et al., 2011). Forest elements of the Euxino-Hyrcanian region along the southern Caspian Sea and Black Sea were less severely affected by Quaternary glaciations making them refugia for numerous species which were mostly eliminated from northern European regions (Leroy & Arpe, 2007). In contrast to most regions of SW Asia, the

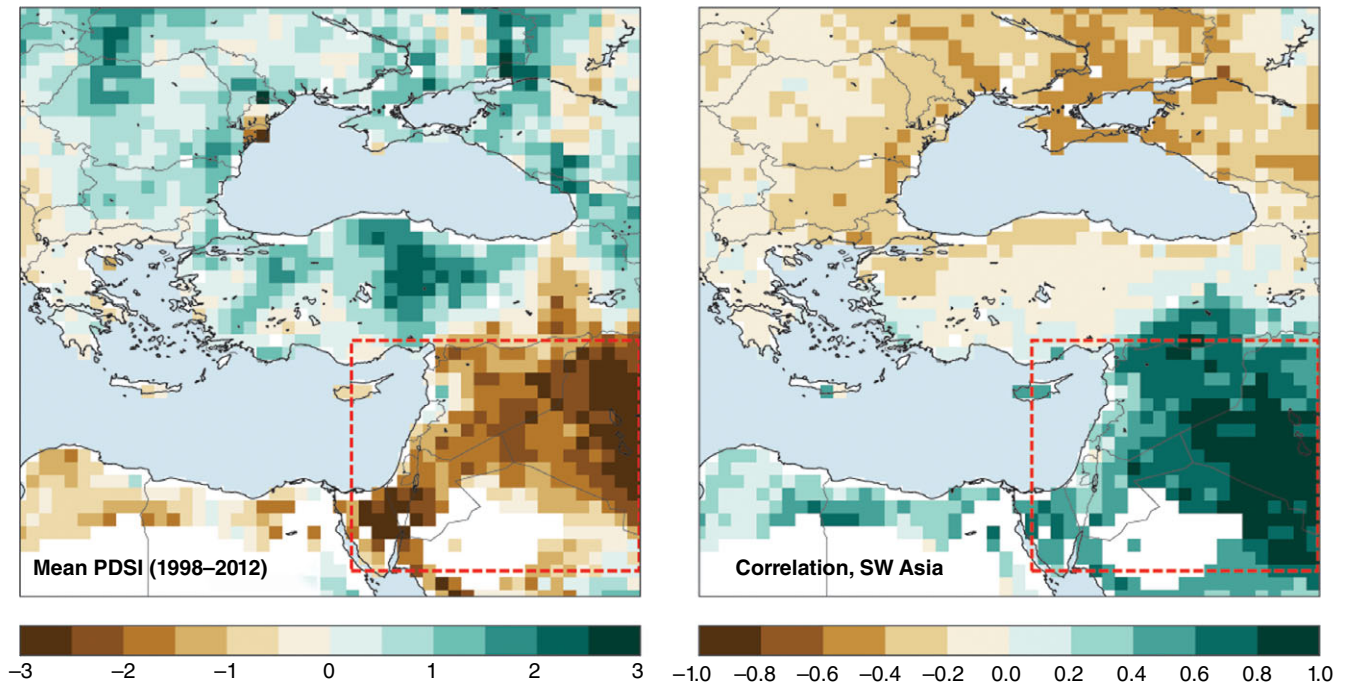


FIGURE 5 Spatial patterns in the Old World Drought Atlas (Cook et al., 2016) of mean Palmer Drought Severity Index (PDSI) values and correlation with the southwest Asia zone (red box) through the last 600 years (Figure 4)

ecotones between the Euxino-Hyrcanian and Irano-Turanian region appear to have been only moderately affected by LGM climates (Miebach, Niestrath, Roeser, & Litt, 2016). Importantly, lowland parts of the Levant also appear to have been only moderately affected by the cold-dry LGM climate; these regions provided crucial refugia for animal and plant resources and human populations during this time period (Asouti et al., 2015; Asouti & Austin, 2005; Roberts et al., 2018).

4.2 | Late glacial vegetation

Increasing number of late glacial vegetation records show moderate afforestation during the late glacial interstadial followed by a re-expansion of steppe during the late glacial stadial (Box 1; Aubert et al., 2017). However, in contrast to the LGM which is almost everywhere characterized by *Artemisia-Chenopodiaceae* steppe reflecting a cold-dry climate, late glacial vegetation dynamics display contrasting patterns in different regions of SW Asia. In the mountain region of NW Iran and eastern Anatolia it is characterized by the dominance of a dry steppe with a slight expansion of boreal trees (e.g., *Betula*) and desert shrubs (*Ephedra*) while in the Levant and the biome transitional zone of northern Turkey, more trees were present in the landscape (Miebach et al., 2016).

4.3 | Early Holocene delayed vegetation advance and the precipitation paradox

Postglacial afforestation by deciduous trees in temperate Europe began at the onset of the Holocene (Berglund, Barnekow, Hammarlund, Sandgren, & Snowball, 1996), however, a different pattern is evident in the much of the Mediterranean region. In the northern and western Mediterranean basin afforestation also occurred at the onset of the Holocene (e.g., Pons & Reille, 1988; Watts, 1985), indicating that the peninsulas of southern Europe and northern borderlands of Greece were important primary refugia. However, pollen data from the south-central Mediterranean (Sicily and Malta) show afforestation delayed until about 7.3 ka (Gambin et al., 2016; Tinner et al., 2009). This Early Holocene delay is also especially marked in the upland interiors of the Irano-Anatolian plateaus (Djamali et al., 2010) where pollen data show a 3,000–5,000-year lag between the onset of climatic amelioration at the beginning of the Holocene and the expansion of deciduous oak woodland.

The delayed response of woodland in inner Anatolia and the Zagros-Anti-Taurus mountains to climatic amelioration at the start of the Holocene has been much discussed in the literature and several hypotheses have been advanced to explain it. Ecological dynamics including autecology (growth rates of individual woodland trees), rates of dispersal, competition, starting positions and locations of primary and secondary refugia, suitable edaphic conditions and physical geographical barriers may all have affected the rate of migration and expansion of forests into regions of sparse tree cover (Roberts et al., 2011; van Zeist & Bottema, 1991).

Early interpretations of pollen data suggested that climatic aridity was responsible (Roberts & Wright, 1993), but more recent stable isotope data show that there was probably increased moisture availability during the early Holocene (Roberts et al., 2008). Other hypotheses invoked suggest that the “oak steppe-parklands” in the Zagros-Anti-Taurus arc and mountains of central Anatolia and NW Iran may have been anthropogenically maintained through the use of both natural and human-induced fire (Roberts, 2002). Micro-charcoal records show the start of the Holocene was a period of frequent/intense wildfires, which would have tended to maintain open (e.g., grassland) vegetation at the expense of closed woodland (Turner, Roberts, Eastwood, Jenkins, & Rosen, 2010). In addition, it is clear the nuts/fruits of terebinth/almond woodland were an important resource for early sedentary communities in central Anatolia. These trees are less visible in the pollen record, but are found in anthracological records from archaeological sites (Asouti & Kabukcu, 2014), and may be evidence of woodland response to increased precipitation and temperature in central Anatolia from the beginning of the Holocene. Landscapes may also have been maintained by people to aid their growth (Baird et al., 2018), as well as for more general vegetation management for grazing, fuel, fodder, and timber (Asouti & Kabukcu, 2014). Although recent data (Roberts et al., 2018) have shown a long-term trend of increasing population for the Neolithic of the early Holocene, it can be postulated that the density of archaeological sites and overall population levels for this time period in the Zagros-Anti-Taurus arc and mountains/plateau of central Anatolia and NW Iran were insufficient to effect vegetation on the regional scale and with the synchronicity that the pollen data appear to infer.

An alternative hypothesis considers that climate seasonality during the early Holocene was greater due to enhanced solar radiation which fueled more intensive ISM circulation, resulting in more rainfall that spilled over into the Arabian Peninsula and North Africa, but led to an extended dry season in the continental interior of the Zagros-Anti-Taurus arc and plateau of central Anatolia and NW Iran, bioclimatically favoring grasslands over deciduous woodlands (Djamali et al., 2010). The important role of seasonality of precipitation has been shown to control biome distribution and biogeographical regionalization in SW Eurasia (Djamali et al., 2012). Furthermore, radiocarbon ages of the expansion of deciduous oak and juniper at pollen sites in the Zagros-Anti-Taurus arc and mountains and plateau of central Anatolia and NW Iran are coeval with the timings for the southward retreat of the Intertropical Convergence Zone (ITCZ) and concomitant southeastward retreat of the ISM. Thus postglacial re-expansion of deciduous oak woodlands was consequently delayed until weakening of the ISM at ~6.3 ka (Djamali et al., 2010).

5 | THE ARCHAEOLOGICAL DATASETS

Given the previous discussion it is worth reflecting on human response to the inferred rises in precipitation and temperature in the early Holocene. Evidence from the Anatolian plateau suggests a rapid human response with sedentary behaviors appearing immediately following the beginning of the Holocene in a major contrast to the very low visibility of settlement evidence on the Anatolian plateau in the Younger Dryas compared to the Bølling-Allerød/late glacial interstadial (Baird et al., 2013, 2018). Interestingly, this response seems significant based on the spread of terebinth/almond woodland juxtaposed with rich grasslands and wetlands supporting large game-like aurochs, boar, and equid in some density, as opposed to sedentism related to increased cereal and legume density as in the late glacial interstadial in the southern Levant (Baird, 2012b). In this section we review some of the wider issues regarding the archaeological datasets available from the region which are important for discussion of human–climate–environment relationships.

5.1 | Data availability

The archaeological and historical datasets from SW Asia and the surrounding regions that can be used to investigate human–environment interaction and the impact of climate and environmental change are, like the palaeoenvironmental records, not resolved at consistent chronological or spatial scales (e.g., Lawrence et al., 2016; Maher et al., 2011; McCormick et al., 2012). Archaeological surveys have varying extent, intensity and methods, sometimes leaving considerable gaps in our knowledge of individual periods and areas with differential distribution of excavated sites and surveyed regions (e.g., de Miroschedji, 2003; Hole, 1987; Petrie, Weeks, Potts, & Roustaei, 2006; Singh et al., 2009; Wilkinson et al., 1994).

Turkey, despite its substantial land area, has witnessed much less intensive archaeological investigation than most areas to its south, notably the Fertile Crescent areas of the Levant and Mesopotamia. This is a conjunction of the size of the land mass, alongside regulations and predilections that have focused work on long-term excavation of major sites, especially of Bronze Age and Classical periods, some excavated for over 100 years (e.g., Hattusas and Ephesus), along with major rescue projects related to various dams on the Tigris and Euphrates rivers. Thus a few areas are well-researched, for example, the Konya Plain, stretches of the Tigris and Euphrates, and limited areas around major sites like Troy and Miletos, while the northern half of the country remains relatively poorly understood. As a result, some questions, for example, the apparent rarity of

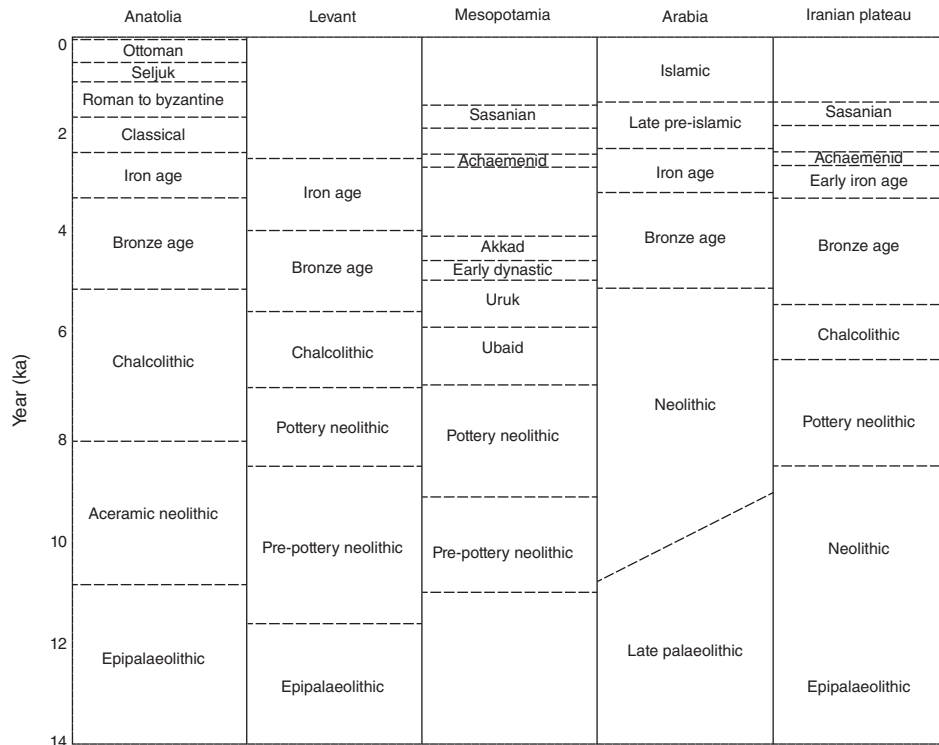


FIGURE 6 Summary of regional archaeological chronologies. Derived from, for Anatolia (Allcock & Roberts, 2014), for the Levant (Finkelstein & Piasetzky, 2007; Maher et al., 2011; Regev, Finkelstein, Adams, & Boaretto, 2014), for Mesopotamia (Matthews, 2013; Nishiaki & Le Miere, 2005), for the Arabian Peninsula (Magee, 2014), and for the Iranian Plateau (Potts, 2013). The dating of some of these periods is complicated and debated, and varies within some of the regions defined here. Some periods were too short-lived to appear on this summary figure. More details can be found in the references provided

Epipalaeolithic and Neolithic sites, are very hard to resolve as genuine phenomena or a function of previous archaeological research focus (Düring, 2010; Düring, Glatz, & Şerifoğlu, 2012). However, the data that are available do still allow some insights into human–environment interactions (Allcock, 2017).

In the case of the prehistory of the Levant, there is considerable disparity between Israel, which has been intensively surveyed and has a high number of recorded archaeological sites, and neighboring regions, where far fewer sites are known. For example, although Lebanon has an analogous geographical and environmental context to Israel, very few Epipalaeolithic, Pre-Pottery Neolithic A and B (see Figure 6 for explanation of archaeological time periods) sites have been documented there, and even fewer have been excavated. Recent fieldwork in both northeast and southeast Jordan has demonstrated that this region was actually much more densely populated during the late Epipalaeolithic and the end of the Neolithic than previously thought (Akkermans, Huigens, & Bruning, 2014; Bertrams et al., 2012; Rambeau et al., 2011; Richter, 2017; Rollefson, Rowan, & Wasse, 2014; Rowan et al., 2015). Recent fieldwork in Saudi Arabia is extending the archaeological record of Neolithic occupation associated with palaeolake deposits and watercourses that were created/activated in the early Holocene (Breeze et al., 2017; Jennings et al., 2015; Matter et al., 2016). Consideration of sea-level changes since the LGM and related marine transgression in the Persian Gulf indicate that previous potential areas of refugia for people during periods of climate-induced stress are now submerged (Cutler, 2013; Rose, Černý, & Bayoumi, 2013).

In southern Mesopotamia, there has been a long tradition of archaeological excavations and survey that provide insight into changing settlement systems, including the rise of urbanism, and the concomitant fluctuations in population size and distribution (e.g., Adams, 1969; Adams & Nissen, 1972). Similar data have also long been available for parts of Iran (e.g., Alden, 2013; Hole, 1987; Smith et al., 1972), and there are growing datasets for northern Mesopotamia and the adjacent uplands (e.g., Lawrence et al., 2016; Wilkinson et al., 1994), and parts of SE Arabia (al-Jahwari & Kennet, 2008, 2010; Magee, 2014) and also south Asia (e.g., Kumar, 2009; Posschl, 1999).

5.2 | Water use

Since the development of agriculture (Box 2) there has been a more pressing need for people to deal with soil-moisture or hydrological droughts, in ways other than moving across the landscape, and these choices and adaptations have been the key to the resulting degree of societal success. Access to water in some areas is reasonably straightforward, such as the parts of northern Mesopotamia and Iran that receive sufficient direct winter rainfall to support dry farming, and the various piedmont areas

that receive broadly predictable water supplies via run-off onto alluvial fans (Petrie & Thomas, 2015; Prickett, 1986; Sherratt, 1980).

Some societies adapted to and developed specific niche environments. Cities in southern Mesopotamia, for example, grew and developed within or on the edges of an anastomosed and deltaic riverine system characterized by abundant salt- and freshwater marshes (Pournelle, Algaze, Crawford, McMahon, & Postgate, 2012; Wilkinson, Ur, & Hritz, 2013). These environments allowed early southern Mesopotamian societies not only access to varied and abundant resources and productive strategies, but also provided a buffer against significant climate fluctuations, such as at 4.2 ka (see discussion above), which are argued to have dramatically affected rain-fed agricultural systems in northern Mesopotamia (Weiss, 2015). In Bronze Age, SE Iran and SE Arabia, integrated oasis agro-pastoral systems developed from the domestication of the date palm, which created an anthropogenic ecosystem of sub-canopy microclimates suitable for the cultivation of a variety of crops in hot and arid environments (Tengberg, 2012).

In other areas, more careful management of water is required and this is evident from at least the middle Pre-Pottery Neolithic B (PPNB, ca. 9 ka) onwards (Richter, 2016). Early Neolithic wells have been documented at Mylouthkia and Shillourokambos on Cyprus (Peltenburg, 2012), while the middle PPNB site of Wadi Abu Tulayha in south eastern Jordan's Jafr Basin produced evidence for a cistern and a dam (Fujii, 2007, 2008). Expansion in the use of water harvesting systems at specific locations, including terraces, qanats, dams and cisterns, may reflect a variety of factors influencing the people living there. Climatic and environmental factors (such as deforestation) may have played a key role in triggering expanded water harvesting. The Pottery Neolithic water harvesting and soil conservation installations at Dhra' (south of the Dead Sea) seem to be an early example of such efforts (Kuijt, Finlayson, & MacKay, 2015), and may be related to the 8.2 ka dry event. Social and economic factors such as increased wealth or larger populations may also be factors in this. The motive for the early Bronze Age expansion of water harvesting and terrace agriculture in Jawa, northeast Jordan, for example, is difficult to ascertain. However, it is noteworthy that the age of the water harvesting systems at the site (Meister et al., 2017) corresponds with the aforementioned 5.2 ka drying event. On the other hand, water harvesting by the Nabateans of Petra from around 2–1.2 ka (Beckers, Schutt, Tsukamoto, & Frechen, 2013) seems to correspond to a wetter episode in the southeastern Mediterranean at the time (Dermody et al., 2012). Moreover, hydrological studies of the terrace systems suggests that they may have been introduced there as flood control measures rather than for water harvesting or as an agricultural installation (Al Qudah, Abdelal, Hamarneh, & Abu-Jaber, 2016).

5.3 | Water demand

The impact a hydroclimatic anomaly will have on a given society depends on the reliance that group has on the resource, their adaption capacity, and how supply can, or cannot continue to meet demand. Recent challenges to SW Asian societies in terms of drought mitigation have been exacerbated by substantial regional population growth and migration (Kelley et al., 2015). It is important to therefore consider resource demand in the past, to the extent that it is possible from the archaeological dataset, as well as the severity of past reduction in rainfall, in assessing the potential risk of drought to past societies.

Estimating population sizes from archaeological datasets is notoriously difficult due to the uncertainties in the representation of sample datasets as being real reflections of larger number of people within sites and across landscapes that we may not be able to detect. While localized estimations of populations may be possible on a site level at well-preserved and extensively excavated sites, in order to extend these to the larger landscape beyond a site, we need improved and coherent methods to assess demographic trends (see Section 6). Even at well-excavated sites, chronological resolution often makes it very difficult to determine whether buildings within the same archaeological phase were occupied simultaneously or slightly apart in time (Birch-Chapman, Jenkins, Coward, & Maltby, 2017). There are also challenges posed by populations that are mobile across the landscapes that they inhabit, who left an archaeological record that is challenging to interpret in terms of demography. Archaeologists have often extrapolated population size from one small excavated area to an entire site by calculating available living space within the excavated area and multiplying it by the total size of the site (Hassan, 1978). This approach, however, is largely untested in real terms, since the scale and character of architecture in the unexcavated parts of the site is unknown, and because there is even less control over chronology beyond the excavated area (Birch-Chapman et al., 2017). Additionally, although site size is often combined with number of archaeological sites in a given area (Sumner, 1994), it is unclear whether such signatures necessarily reflect population growth or whether they are indicative of population aggregation. In arid areas such as Arabia, assessing population sizes in the early-mid Holocene is complicated by the common practice of residential mobility at various geographical scales (Cavulli & Scaruffi, 2013; Crassard & Drechsler, 2013; Lézine et al., 2010; McCorriston, Harrower, Martin, & Oches, 2012; Zazzo et al., 2016). Moreover, for the subsequent Bronze Age agro-pastoral communities of eastern Arabia, the burial archaeology record is at times much more prominent than that of contemporaneous settlement (Cleuziou & Tosi, 2007; Højlund, 2007), providing an alternative perspective on population size, distribution and, where

skeletal remains are preserved, local demographic parameters (e.g., Baustian & Martin, 2010; Martin, 2007; McSweeney, Méry, & al Tikriti, 2010).

5.4 | Resilience or collapse?

There has been considerable interest in linking climate events identified in climate proxy records, and periods and phases of cultural transformation identified in archaeological sequences (Büntgen et al., 2016). In this regard it is interesting to compare the human responses to well-documented climate change events in inner Anatolia (see also Section 3). Both the 8.2 and 4.2 ka events are seen in Anatolian plateau proxies (as discussed above). The 4.2 ka event seems particularly significant for people in terms of either a major decline in frequency and scale of sites in areas like the Çarşamba fan (Baird, 2001) and other parts of Anatolia (Bachhuber, 2015), or major settlement restructuring with population concentration and growth of a few large sites. In the same areas the 8.2 ka event is not represented by any major settlement discontinuities (Baird, 2012a), although there may be subtle responses such as those also seen in north Syria. Rather than collapse, the 8.2 ka event on the Anatolian plateau has been linked with population displacement (Clare & Weninger, 2014) and the spread of the Neolithic into western Anatolia, although the latter has been recently questioned (Berger et al., 2016; Kılınc et al., 2017).

The contrasts in the same areas between these two climatic events may be due to their relative severity, but this was likely combined with the resilience of Neolithic communities at lower population levels and with significant flexibility in subsistence practices. These may have included ability to vary mobility, degrees of reliance on agriculture relative to foraging and pastoralism, and landscape exploitation practices that may have had significant incidental or deliberate conservation features. By 4.2 ka human populations had achieved much greater scale and density than that of previous millennia (Bachhuber, 2015; Baird, 2001), highly dependent on large-scale agriculture in prime but sensitive semiarid locations like the Konya basin alluvial fans, and were more dependent on elaborate land management schemes. People here had probably started to impact the landscape at a level much greater than in earlier millennia and operated in highly competitive political contexts (Bachhuber, 2015).

The western parts of South Asia are affected by both winter Mediterranean and summer monsoonal rainfall systems that produces a climatically and ecologically diverse landscape that is subject to interannual and more long-term variability. In this context, the available palaeoclimate records have variable proximity to archaeological sites and regions of interest, which means that it is illogical to draw simple correlations between evidence for climate change and cultural changes observed in the archaeological record (e.g., Petrie, 2017; Petrie et al., 2017; R. P. Wright, 2010). Furthermore, beyond spatial coverage, one of the biggest challenges with palaeoclimate records from South Asia has been the limited chronological control of the data. Most of the early climate records have poor chronological control (Madella & Fuller, 2006), and while this has been improved by more recent studies (Berkelhammer et al., 2012; Dixit et al., 2018; Dixit, Hodell, & Petrie, 2014; Dixit, Hodell, Sinha, & Petrie, 2014), their spatial distribution remains limited and contested (Finné, Holmgren, Sundqvist, Weiberg, & Lindblom, 2011). It will only be with spatially proximate proxy records from different zones within the region that it will be possible to characterize climatic variability accurately across the region as a whole, and link to local archaeological evidence.

6 | MODELLING CLIMATE, PROXIES, AND HUMAN RESPONSES

Modelling can help fill some of the gaps in the spatial and temporal coverage of climatic, hydrological and archaeological data identified above, and allow uncertainties in interpretations of these data to be better understood.

6.1 | Modelling climate

Concerted modelling efforts of past climates such as the Climate and Paleoclimate Modelling Intercomparison Projects (CMIPs, PMIPs), has usually focused only on key intervals such as the LGM (~21 ka) and the mid-Holocene (~6 ka). In SW Asia these time intervals are often transition periods in proxy records, leading to some difficulties in data model comparison (Reuter, Buening, & Yoshimura, 2017). The most comprehensive modelling study of the Mediterranean region for the Holocene yet attempted has been performed by Brayshaw, Rambeau, and Smith (2011) who analyzed regional Mediterranean climate simulations, carried out at 2,000-year intervals from 0 ka to 12 ka. At 6 ka the model suggested that the annual average surface air temperature in the region was relatively similar to modern levels, however, cooler winters and warmer summers caused by insolation change meant that the seasonal cycle was amplified by 2–3°C. A multimodel ensemble from PMIP2 (Braconnot et al., 2007), suggested that Turkey and SW Asia were 2–5°C cooler than today at the LGM, and 1–2°C warmer at 6 ka. Brayshaw et al. (2011) did not find substantial summer precipitation in the eastern Mediterranean at any time 0–12 ka, however, during the wet season (October–May) simulations representing 8–12 ka showed an increase in precipitation over the

eastern Mediterranean (Turkey, Syria, Lebanon, Jordan). They suggested that these patterns were due to changes in lower tropospheric circulation which meant a stronger south-westerly flow in winter. Perez-Sanz, Li, Gonzalez-Samperiz, and Harrison (2014) noted that CMIP5 models generally show an increase in 6 ka precipitation in the Mediterranean region, but also noted the difficulties in modelling precipitation for the region.

The challenge of model disagreement is further highlighted by modelling of LGM precipitation across the region. Robinson et al. (2006) suggested that the eastern Mediterranean and the Levant had reduced winter precipitation relative to modern levels; but Arpe, Leroy, and Mikolajewicz (2011) showed that with a high resolution version of the ECHAM5 climate model, a shift in circulation at the LGM could lead to enhanced precipitation over the Levant. This shift in circulation did not occur in lower spatial resolution versions of ECHAM5. Models generally agree that the seasonal patterns of precipitation are similar between the LGM and the preindustrial period (ca. 1750 AD) with most of the precipitation in the winter season and summers being very dry. Robinson et al. (2006) noted that LGM evaporation was in excess of LGM precipitation in this region, despite the drop in temperature. They also note that in the LGM a significant amount of the modelled winter precipitation falls as snow.

The addition of isotope hydrology to General Circulation Models of climate potentially improves the comparison of climate model output to proxy records of change. Risi, Bony, Vimeux, and Jouzel (2010) modelled global changes in the oxygen isotope values of precipitation ($\delta^{18}\text{O}_p$) between the LGM and the preindustrial period. In SW Asia they found no strong change in $\delta^{18}\text{O}_p$ between the two periods, despite the LGM being 3–9°C cooler. Reuter et al. (2017) found that modelled $\delta^{18}\text{O}_p$ at 6 ka was more negative than present day, due to a combination of effects including changes in local surface temperature and precipitation amount.

6.2 | Proxy system models

Palaeoclimatic archives and proxies filter the climate signal, as discussed above, such that comparison of climate model output with geological records requires an understanding of the systematics of the archive, such as that provided by Proxy System Models (PSMs; Dee et al., 2015; Evans, Tolwinski-Ward, Thompson, & Anchukaitis, 2013). These models have different levels of complexity from regression-based models, such as of PDSI from tree rings (Cook et al., 2016), through to transfer function approaches from pollen (Eastwood et al., 2007) or diatoms (Woodbridge & Roberts, 2011), and more mechanistic models of lakes and wetlands (Jones & Imbers, 2010; Rohling, 2016). The more mechanistic PSMs allow for the investigation of the impact of changing climate seasonality, including the impact of changing amounts of snow relative to rainfall, on these archive systems (Dean et al., 2018).

Given our discussions above, a key challenge in PSMs for SW Asia is correctly modelling the hydrology of a given archive, a crucial part of the *Environment Model* in the PSM framework of Evans et al. (2013). Dryland environments present particular challenges for hydrological models (Wheater, Sorooshian, & Sharma, 2007), but there has been some success in using hydraulic models, which can better represent the water routing at the event scale (e.g., Jarhani, Larsen, Callow, McVicar, & Johansen, 2015; Massuel et al., 2011). The very large spatial heterogeneity in rainfall, soil properties, and vegetation that combine to produce runoff in dryland environments is difficult to capture (Jothityangkoon, Sivapalan, & Farmer, 2001), meaning the processes driving modern dryland hydrology remain poorly understood. There is a lack of understanding of hydrological systems in the present day, from systems with heavy anthropogenic modification, to aid benchmarking such models. On top of this, groundwater spring resources may have notably lagged responses to climate (see discussion above) and, therefore, these proxies of water availability may be spatially and temporally incoherent, and have little fit to climate model outputs. This is a substantial knowledge gap, and leads to challenges in constraining the water supply for past populations.

6.3 | Modelling human water demand

A number of recent studies have begun to use computational techniques to address questions of settlement scale and resource demand in the archaeological record. At the most basic level, the presence of humans in a region requires a supply of accessible, potable water. This has a bearing on the most arid regions of SW Asia, which appear to have only been periodically inhabitable over the last 20,000 years (Groucutt & Petraglia, 2012). The known Upper Palaeolithic and Epipalaeolithic record in Arabia is sparse (Maher, 2010), and the region was not recolonized until the late Pleistocene/Early Holocene (Crassard, Petraglia, Drake, et al., 2013; Crassard, Petraglia, Parker, et al., 2013; Hilbert et al., 2014). Consequently, Breeze et al. (2015) have demonstrated that modelled palaeohydrological features are a good predictor of the location of archaeological sites in the hyperarid interior of the Arabian Peninsula. Bretzke, Drechsler, and Conard (2012) also found that modelled water availability is correlated with the distribution of sites in the Palaeolithic of the Syrian Desert, both in terms of where settlements were located, and the intensity of occupation through time.

On a larger scale, the aggregate demand for water in the region can be assumed to track its net population. The difficulty of deriving absolute estimates of population size in archaeological contexts was discussed previously, but in recent years archaeologists have increasingly used the summed radiocarbon or “dates as data” method (Chaput & Gajewski, 2016; Rick, 1987) to model long-term demographic trends. Although not without its critics (Bamforth & Grund, 2012; Contreras & Meadows, 2014; Torfing, 2015), the application of this technique has produced significant insights into the relationship between population and climate (Bevan et al., 2017; Shennan et al., 2013) and the results correlate well with other demographic proxies (Downey, Bocaage, Kerig, Edinborough, & Shennan, 2014; Lechterbeck et al., 2014; Woodbridge et al., 2012).

Several authors have recently applied the summed radiocarbon technique to corpuses of dates from SW Asia, focused primarily on the Pleistocene–Holocene transition (Borrell, Junno, & Barcelo, 2015; Flohr et al., 2016; Roberts et al., 2018). The overall trend evident in each study is one of the exponential population growth, with a marked “boom” beginning around 14.5 ka. This agrees with earlier population growth estimates based on palaeoanthropological (Guerrero, Naji, & Bocquet-Appel, 2008) and site frequency data (Goring-Morris & Belfer-Cohen, 2010). The evidence for demographic responses to specific climate events is more equivocal. Borrell et al. (2015) found a pronounced break in settlement in the northern Levant—the “near abandonment of the region”—corresponding to early Holocene warming, but this was not replicated in the larger dataset of Roberts et al. (2018). The latter authors did see a synchronicity between climate events and population events, but stressed the overall continuity of settlement in the region. Similarly, Flohr et al. (2016) tested the human response to the 9.2 and 8.2 ka climate events and found that the population was resilient to both. All three studies found evidence for significant regional variation in population dynamics and demographic response to climate, highlighting the importance of factoring local environmental conditions into this class of model.

For the later prehistoric and historic periods, demographic trends can be modelled more concretely using quantitative historical data, an approach sometimes dubbed “cliodynamics” (Turchin, 2008). Reba, Reitsma, and Seto (2016) recently produced a database of estimated urban population growth between 5.7 ka and the present. This database is of particular interest because not only was SW Asia the location of the world's earliest urban societies, the advent and growth of these societies is frequently linked to growing demand and centralized management of water resources (Wilkinson & Rayne, 2010). Integrating this data into linked models with the climate record, and with the prehistoric summed radiocarbon sequence, is therefore a promising avenue for future research.

Modelling has also been used to reconstruct water demand on subregional or site-local scales. This reveals considerable variation in people's reliance on hydrological resources depending on the environmental and cultural context. In the Late Pleistocene of steppe eastern Jordan, for example, Byrd, Garrard, and Brandy (2016) modelled mobile foragers as needing to be only within one or two days' walk of freshwater, and posit that they only temporarily gathered at perennial water sources during the dry season. By contrast, Whitehead, Smith, and Wade (2011) and Whitehead et al. (2008) constructed a model of the hydrological resources of the nearby early Bronze Age site of Jawa, concluding that its system of diversions and storage ponds was constructed to support a permanent, sedentary population of 6,000 or more people and their livestock. Even in the same environmental zone, therefore, the human response to hydroclimatic events can be expected to differ radically between social contexts.

7 | SUMMARY

Our review of this topic has highlighted the continued challenges involved in answering the key research questions regarding human vulnerability to climate change over the last 20,000 years. The paper illustrates a considerable, and growing, body of literature from across the palaeoscience disciplines dealing with these issues in SW Asia. Development of new analytical technologies will continue to provide new ways of examining archaeological and palaeoenvironmental archives and further add to our knowledge base. For example, new excavations and palaeoclimate reconstructions in Anatolia, Iraqi Kurdistan, southern Mesopotamia, Iran, Arabia, and South Asia are providing opportunities, not only to improve the spatial resolution of data available, but to implement the use of the full-spectrum of state of the art bio-archaeological approaches. Archaeobotany and archaeozoology have been widely used for some time, but isotopic analysis of human and animal remains is being increasingly utilized to investigate questions of mobility, diet and the impact of climate change on water availability and use (e.g., Chase, Ajithprasad, Rajesh, Patel, & Sharma, 2014; Chase, Meiggs, Ajithprasad, & Slater, 2014, 2018; P. J. Jones, 2018; Kenoyer, Price, & Burton, 2013; Kutterer & Uerpmann, 2012; Riehl, Pustovoytov, Weippert, Klett, & Hole, 2014; Valentini et al., 2015). Residue analysis of ceramics from new excavations, and also museum collections, is also now being attempted (Craig et al., 2013; Gregg, 2010; Gregg, Banning, Gibbs, & Slater, 2009).

Coming to a conclusion regarding ongoing work is difficult, so here we summarize two key themes that pervade this paper and are likely foci of future work. While collaboration across the palaeoscience disciplines working in SW Asia has often been the norm, our review here highlights the need for these communities to look further, to hydrologists, soils scientists and

modelers, for example, to take such work forward. We need to fully integrate data and approaches, new and existing, in interdisciplinary ways towards answering focused research questions (Jones, 2013).

7.1 | Scale

Although modelling, quantitatively or conceptually, is likely to aid future interpretation of new and existing datasets, a major challenge is integrating widely divergent scales of archaeological and palaeohydrological data. For example, how do we best compare regional summed radiocarbon population proxies to site-specific water management strategies, with palaeoenvironmental data conditioned by regional climatic and local hydrological and ecological states?

To aid resolving such complexity more care needs to be taken in understanding the likely spatial relevance of a given proxy record, and the public availability of instrumental meteorological and climate-model data give one way make this possible (e.g., Jex, Phipps, Baker, & Bradley, 2013; Yiğitbaşıoğlu et al., 2015), alongside improved monitoring of archive systems (e.g., Dean, Eastwood, et al., 2015; Djamali et al., 2009). Some of the nonconformity between scales can also be addressed by palaeoenvironmental work in direct association with individual archaeological sites, that is, from the local hydroscape, and only then linking to, often more continuous, palaeoclimate records from further away (Jones, Maher, Richter, et al., 2016). Modelling, particularly of local, hydrology is often a “missing link” in the palaeonarratives reviewed here. Appreciation of a given locale's place in the regional hydroscape, be it a palaeoenvironmental archive or archaeological site, allows more robust links to climate to be made where possible, and demands the resolution of factors acting across multiple scales.

7.2 | Seasonality

Palaeoenvironmental records rarely have the resolution to pick apart conditions in multiple seasons of a given year, but many reflect conditions of a particular season (Dean et al., 2018). Given the very different seasonal conditions across SW Asia, even with a first-order conceptual hydrological understanding, the balance between the now predominant winter wet season and summer dry season is likely to be important for water availability, and the climatic systems that ultimately control these conditions range from the Atlantic Ocean to east Asia, and from Iceland to south of the equator. Over the last 20,000 years, the timing (Stevens et al., 2001) and type (Robinson et al., 2006) of precipitation across the region may have changed, as would have the degree of summer evaporation (Djamali et al., 2010). The balance of seasonal precipitation and evaporation patterns is a primary control on recharge, and will therefore impact water availability and how a given palaeoenvironmental archive records shifting patterns. Changing amounts of snowfall, and snowmelt, would also alter this filtering of climate by hydrological systems, and are an important focus for future work, especially in understanding glacial–interglacial changes in proxy records from the region.

7.3 | Close

The water resources available to a given group of people at any time is a function of climate moderated by landscape and technology, such as vegetation type and quantity, soil cover and stability, geology, and topography, which influence surface- and ground-water systems, and the people themselves. People have been making the most of these resources, and adapting as they change, across SW Asia for over 20,000 years. Whatever the complications in detailing the finer points of past societies' relationships with changing water availability it is clear that dealing with changing elements of that relationship is far from a new issue for people, in SW Asia arguably for longer than anywhere else. Our review here highlights the need for a multidisciplinary, multiscale approach to furthering work in this area. A 20,000-year perspective shows that societies are often vulnerable to changing climate and have sometimes struggled to adapt. The implications for resilience in our own, present day and projected future contexts are salutary.

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CONFLICT OF INTEREST

The authors have declared no conflicts of interest for this article.

Mediterranean wetlands: Archaeology, ecology, and sustainability

Deciphering long-term records of natural variability and human impact as recorded in lake sediments: A palaeolimnological puzzle

ORCID

Matthew D. Jones  <https://orcid.org/0000-0001-8116-5568>

Mark O. Cuthbert  <https://orcid.org/0000-0001-6721-022X>

Jonathan R. Dean  <https://orcid.org/0000-0001-8071-000X>

Joshua Larsen  <https://orcid.org/0000-0002-0650-7369>

Cameron A. Petrie  <https://orcid.org/0000-0002-2926-7230>

Joe Roe  <https://orcid.org/0000-0002-1011-1244>

Lloyd Weeks  <https://orcid.org/0000-0003-4736-9633>

REFERENCES

- Abu-Jaber, N., & El-Naser, H. (2016). Geology and hydrochemistry of the deep sandstone aquifers of Jordan. *Environmental Earth Sciences*, 75(10), 875. <https://doi.org/10.1007/S12665-016-5680-8>
- Adams, R. M. (1969). The study of ancient Mesopotamian settlement patterns and the problem of urban origins. *Sumer*, 25(1), 111–124.
- Adams, R. M., & Nissen, H. J. (1972). *The Uruk countryside: The natural setting of urban societies*. Chicago, IL: University of Chicago Press.
- Ahmed, M., Anchukaitis, K. J., Asrat, A., Borgaonkar, H. P., Braida, M., Buckley, B. M., ... Zorita, E. (2013). Continental-scale temperature variability during the past two millennia. *Nature Geoscience*, 6(5), 339–346. <https://doi.org/10.1038/NNGEO1797>
- Akkermans, P. M. M. G., Huigens, H. O., & Bruning, M. L. (2014). A landscape of preservation: Late prehistoric settlement and sequence in the Jebel Qurma region, north-eastern Jordan. *Levant*, 46(2), 186–205. <https://doi.org/10.1179/0075891414Z.00000000041>
- Al Qudah, K., Abdelal, Q., Hamarneh, C., & Abu-Jaber, N. (2016). Taming the torrents: The hydrological impacts of ancient terracing practices in Jordan. *Journal of Hydrology*, 542, 913–922. <https://doi.org/10.1016/j.jhydrol.2016.09.061>
- Alden, J. (2013). The Kur River Basin in the Proto-Elamite era—surface survey, settlement patterns, and the appearance of full-time transhumant pastoral nomadism. In C. Petrie (Ed.), *Ancient Iran and its neighbours: Local developments and long-range interactions in the fourth millennium BC* (pp. 207–232). Oxford, England: Oxbow Books.
- al-Jahwari, N. S., & Kennet, D. (2010). Umm an-Nar settlement in the Wadi Andam (Sultanate of Oman). *Proceedings of the Seminar for Arabian Studies*, 40, 201–212.
- al-Jahwari, N. S., & Kennet, D. (2008). *A Field Methodology for the Quantification of Ancient Settlement in an Arabian Context*. Paper presented at the proceedings of the seminar for Arabian studies.
- Allcock, S. L. (2017). Long-term socio-environmental dynamics and adaptive cycles in Cappadocia, Turkey during the Holocene. *Quaternary International*, 446, 66–82. <https://doi.org/10.1016/j.quaint.2017.06.065>
- Allcock, S. L., & Roberts, N. (2014). Changes in regional settlement patterns in Cappadocia (central Turkey) since the Neolithic: A combined site survey perspective. *Anatolian Studies*, 64, 33–57. <https://doi.org/10.1017/S0066154614000040>
- Alley, R. B., & Agustsdottir, A. M. (2005). The 8k event: Cause and consequences of a major Holocene abrupt climate change. *Quaternary Science Reviews*, 24(10–11), 1123–1149. <https://doi.org/10.1016/j.quascirev.2004.12.004>
- Arpe, K., Leroy, S. A. G., & Mikolajewicz, U. (2011). A comparison of climate simulations for the last glacial maximum with three different versions of the ECHAM model and implications for summer-green tree refugia. *Climate of the Past*, 7(1), 91–114. <https://doi.org/10.5194/cp-7-91-2011>
- Asouti, E., & Austin, P. (2005). Reconstructing woodland vegetation and its exploitation by past societies, based on the analysis and interpretation of archaeological wood charcoal macro-remains. *Environmental Archaeology*, 10(1), 1–18. <https://doi.org/10.1179/env.2005.10.1.1>
- Asouti, E., & Kabukcu, C. (2014). Holocene semi-arid oak woodlands in the Irano-Anatolian region of southwest Asia: Natural or anthropogenic? *Quaternary Science Reviews*, 90, 158–182. <https://doi.org/10.1016/j.quascirev.2014.03.001>
- Asouti, E., Kabukcu, C., White, C. E., Kuijt, I., Finlayson, B., & Makarewicz, C. (2015). Early Holocene woodland vegetation and human impacts in the arid zone of the southern Levant. *The Holocene*, 25(10), 1565–1580. <https://doi.org/10.1177/0959683615580199>
- Aubert, C., Brisset, E., Djarnali, M., Sharifi, A., Ponel, P., Gambin, B., ... Gandouin, E. (2017). Late glacial and early Holocene hydroclimate variability in northwest Iran (Talesh Mountains) inferred from chironomid and pollen analysis. *Journal of Paleolimnology*, 58(2), 151–167. <https://doi.org/10.1007/s10933-017-9969-8>
- Bachhuber, C. (2015). *Citadel and cemetery in early bronze age anatolia*. Sheffield, England: Equinox Publishing.
- Badertscher, S., Borsato, A., Frisia, S., Cheng, H., Edwards, R. L., Tuysuz, O., & Fleitmann, D. (2014). Speleothems as sensitive recorders of volcanic eruptions—The Bronze Age Minoan eruption recorded in a stalagmite from Turkey. *Earth and Planetary Science Letters*, 392, 58–66. <https://doi.org/10.1016/j.epsl.2014.01.041>
- Bae, C. J., Douka, K., & Petraglia, M. D. (2017). On the origin of modern humans: Asian perspectives. *Science*, 358(6368), eaai9067. <https://doi.org/10.1126/science.aai9067>
- Baird, D. (2001). Settlement and landscape in the Konya Plain, south central Turkey, from the epipalaeolithic to the Medieval period. In H. Karpuz, A. Bas, & R. Duran (Eds.), *Uluslararası Çatalhöyük'ten Günümüze Çumra Kongresi* (pp. 269–276). Çumra, Turkey: Çumra Belediyesi.
- Baird, D. (2012a). The Late Epipaleolithic, Neolithic, and Chalcolithic of the Anatolian Plateau, 13,000–4000 BC. In D. T. Potts (Ed.), *A companion to the Archaeology of the Ancient Near East* (pp. 431–465). Chichester, England: John Wiley and Sons.
- Baird, D. (2012b). Pınarbaşı: From Epipalaeolithic campsite to sedentary village in central Anatolia. In M. Ozgogan, N. Basgelen, & P. adn Juniholm (Eds.), *The Neolithic in Turkey* (Vol. 3, pp. 181–218). Istanbul, Turkey: Arkeoloji ve Sanat Yayınları.
- Baird, D., Asouti, E., Astruc, L., Baysal, A., Baysal, E., Carruthers, D., ... Pirie, A. (2013). Juniper smoke, skulls and wolves' tails. The Epipalaeolithic of the Anatolian plateau in its south-west Asian context; insights from Pınarbaşı. *Levant*, 45(2), 175–209. <https://doi.org/10.1179/0075891413Z.00000000024>
- Baird, D., Fairbairn, A., Jenkins, E., Martin, L., Middleton, C., Pearson, J., ... Elliott, S. (2018). Agricultural origins on the Anatolian plateau. *Proceedings of the National Academy of Sciences*, 115(14), E3077–E3086. <https://doi.org/10.1073/pnas.1800163115>
- Baker, A., Duan, W., Cuthbert, M., Treble, P., Banner, J., & Hankin, S. (2018). Climatic influences on the offset between $\delta^{18}\text{O}$ of cave drip waters and precipitation inferred from global monitoring data.

- Balter, M. (2010). Archaeology. The tangled roots of agriculture. *Science*, 327(5964), 404–406. <https://doi.org/10.1126/science.327.5964.404>
- Bamforth, D. B., & Grund, B. (2012). Radiocarbon calibration curves, summed probability distributions, and early Paleoindian population trends in North America. *Journal of Archaeological Science*, 39(6), 1768–1774. <https://doi.org/10.1016/j.jas.2012.01.017>
- Barker, G. (2009). *The agricultural revolution in prehistory: Why did foragers become farmers?* Oxford, England: Oxford University Press.
- Bar-Matthews, M., & Ayalon, A. (2011). Mid-Holocene climate variations revealed by high-resolution speleothem records from Soreq Cave, Israel and their correlation with cultural changes. *The Holocene*, 21(1), 163–171. <https://doi.org/10.1177/0959683610384165>
- Bar-Matthews, M., Ayalon, A., Gilmour, M., Matthews, A., & Hawkesworth, C. J. (2003). Sea–land oxygen isotopic relationships from planktonic foraminifera and speleothems in the Eastern Mediterranean region and their implication for paleorainfall during interglacial intervals. *Geochimica et Cosmochimica Acta*, 67(17), 3181–3199. [https://doi.org/10.1016/S0016-7037\(02\)01031-1](https://doi.org/10.1016/S0016-7037(02)01031-1)
- Bar-Matthews, M., Ayalon, A., Kaufman, A., & Wasserburg, G. J. (1999). The Eastern Mediterranean paleoclimate as a reflection of regional events: Soreq cave, Israel. *Earth and Planetary Science Letters*, 166(1–2), 85–95. [https://doi.org/10.1016/S0012-821x\(98\)00275-1](https://doi.org/10.1016/S0012-821x(98)00275-1)
- Bartov, Y., Goldstein, S. L., Stein, M., & Enzel, Y. (2003). Catastrophic arid episodes in the Eastern Mediterranean linked with the North Atlantic Heinrich events. *Geology*, 31(5), 439–442. [https://doi.org/10.1130/0091-7613\(2003\)031<0439:caeite>2.0.co;2](https://doi.org/10.1130/0091-7613(2003)031<0439:caeite>2.0.co;2)
- Baruch, U., & Bottema, S. (1999). A new pollen diagram from Lake Hula: vegetational, climatic and anthropogenic implications. In H. Kawanabe, G. W. Coulter, & A. C. Roosevelt (Eds.), *Ancient lakes. their cultural and biological diversity* (pp. 75–86). Gent, Belgium: Kenobi Productions.
- Bar-Yosef, O. (2009). Social changes triggered by the Younger Dryas and the early Holocene climatic fluctuations in the Near East. In C. T. Fisher, J. Brett Hill, & G. M. Feinman (Eds.), *The archaeology of environmental change: Socionatural legacies of degradation and resilience* (pp. 192–208). Tucson, AZ: University of Arizona Press.
- Bar-Yosef, O., & Belfer-Cohen, A. (2002). Facing environmental crisis. Societal and cultural changes at the transition from the Younger Dryas to the Holocene in the Levant. In *Workshop on Transition from foraging to farming in southwest Asia, Dawn of farming in the near East*, 1998 (Vol. 6, pp. 55–66), Groningen, The Netherlands.
- Baustian, K., & Martin, D. (2010). Patterns of mortality in a Bronze Age tomb from Tell Abraq. In *Death and burial in Arabia and beyond: Multidisciplinary perspectives* p. 55, 59. Oxford, England: Archaeopress.
- Beckers, B., Schutt, B., Tsukamoto, S., & Frechen, M. (2013). Age determination of Petra's engineered landscape—Optically stimulated luminescence (OSL) and radiocarbon ages of runoff terrace systems in the Eastern Highlands of Jordan. *Journal of Archaeological Science*, 40(1), 333–348. <https://doi.org/10.1016/j.jas.2012.06.041>
- Benjamin, J., Rovere, A., Fontana, A., Furlani, S., Vacchi, M., Inglis, R. H., ... Gehrels, R. (2017). Late quaternary sea-level changes and early human societies in the central and eastern Mediterranean Basin: An interdisciplinary review. *Quaternary International*, 449, 29–57. <https://doi.org/10.1016/j.quaint.2017.06.025>
- Berger, A., & Loutre, M.-F. (1991). Insolation values for the climate of the last 10 million years. *Quaternary Science Reviews*, 10(4), 297–317.
- Berger, J. F., Lespez, L., Kuzucuoglu, C., Glais, A., Hourani, F., Barra, A., & Guilaine, J. (2016). Interactions between climate change and human activities during the early to mid-Holocene in the eastern Mediterranean basins. *Climate of the Past*, 12(9), 1847–1877. <https://doi.org/10.5194/cp-12-1847-2016>
- Berglund, B. E., Barnekow, L., Hammarlund, D., Sandgren, P., & Snowball, I. F. (1996). Holocene forest dynamics and climate changes in the Abisko Area, Northern Sweden: The Sonesson model of vegetation history reconsidered and confirmed. *Ecological Bulletins*, 45, 15–30.
- Berkelhammer, M., Sinha, A., Stott, L., Cheng, H., Pausata, F. S., & Yoshimura, K. (2012). An abrupt shift in the Indian monsoon 4000 years ago. In L. Giosan, D. Q. Fuller, K. Nicoll, R. K. Flad, & P. D. Clift (Eds.), *Climates, landscapes, and civilizations* (pp. 75–88). Washington, DC: American Geophysical Union.
- Bertrams, M., Protze, J., Lohrer, R., Schyle, D., Richter, J., Hilgers, A., ... Lehmkuhl, F. (2012). Multiple environmental change at the time of the modern human passage through the Middle East: First results from geoarchaeological investigations on Upper Pleistocene sediments in the Wadi Sabra (Jordan). *Quaternary International*, 274, 55–72. <https://doi.org/10.1016/j.quaint.2012.02.047>
- Bevan, A., Colledge, S., Fuller, D., Fyfe, R., Shennan, S., & Stevens, C. (2017). Holocene fluctuations in human population demonstrate repeated links to food production and climate. *Proceedings of the National Academy of Sciences of the United States of America*, 114(49), E10524–E10531. <https://doi.org/10.1073/pnas.1709190114>
- Birch-Chapman, S., Jenkins, E., Coward, F., & Maltby, M. (2017). Estimating population size, density and dynamics of Pre-Pottery Neolithic villages in the central and southern Levant: An analysis of Beidha, southern Jordan. *Levant*, 49(1), 1–23. <https://doi.org/10.1080/00758914.2017.1287813>
- Blaauw, M., & Christen, J. A. (2011). Flexible paleoclimate age-depth models using an autoregressive gamma process. *Bayesian Analysis*, 6(3), 457–474. <https://doi.org/10.1214/11-BA618>
- Blockley, S. P. E., & Pinhasi, R. (2011). A revised chronology for the adoption of agriculture in the Southern Levant and the role of Lateglacial climatic change. *Quaternary Science Reviews*, 30(1–2), 98–108. <https://doi.org/10.1016/j.quascirev.2010.09.021>
- Bond, G., Showers, W., Cheseby, M., Lotti, R., Almasi, P., deMenocal, P., ... Bonani, G. (1997). A pervasive millennial-scale cycle in North Atlantic Holocene and glacial climates. *Science*, 278(5341), 1257–1266. <https://doi.org/10.1126/science.278.5341.1257>
- Borrell, F., Junno, A., & Barcelo, J. A. (2015). Synchronous environmental and cultural change in the emergence of agricultural economies 10,000 years ago in the Levant. *PLoS One*, 10(8), e0134810. <https://doi.org/10.1371/journal.pone.0134810>
- Braconnot, P., Otto-Bliesner, B., Harrison, S., Joussaume, S., Peterchmitt, J. Y., Abe-Ouchi, A., ... Zhao, Y. (2007). Results of PMIP2 coupled simulations of the mid-Holocene and Last Glacial Maximum. Part 1: Experiments and large-scale features. *Climate of the Past*, 3(2), 261–277. <https://doi.org/10.5194/cp-3-261-2007>
- Brayshaw, D. J., Rambeau, C. M. C., & Smith, S. J. (2011). Changes in Mediterranean climate during the Holocene: Insights from global and regional climate modeling. *The Holocene*, 21(1), 15–31. <https://doi.org/10.1177/0959683610377528>
- Breeze, P. S., Drake, N. A., Groucutt, H. S., Parton, A., Jennings, R. P., White, T. S., ... Petraglia, M. D. (2015). Remote sensing and GIS techniques for reconstructing Arabian palaeohydrology and identifying archaeological sites. *Quaternary International*, 382, 98–119. <https://doi.org/10.1016/j.quaint.2015.01.022>
- Breeze, P. S., Groucutt, H. S., Drake, N. A., Louys, J., Scerri, E. M. L., Armitage, S. J., ... Petraglia, M. D. (2017). Prehistory and palaeoenvironments of the western Nefud Desert, Saudi Arabia. *Archaeological Research in Asia*, 10, 1–16. <https://doi.org/10.1016/j.ara.2017.02.002>
- Bretzke, K., Drechsler, P., & Conard, N. J. (2012). Water availability and landuse during the Upper and Epipaleolithic in southwestern Syria. *Journal of Archaeological Science*, 39(7), 2272–2279. <https://doi.org/10.1016/j.jas.2012.02.033>
- Briffa, K. R., Jones, P. D., Bartholin, T. S., Eckstein, D., Schweingruber, F. H., Karlen, W., ... Eronen, M. (1992). Fennoscandian Summers from Ad-500—Temperature changes on short and long timescales. *Climate Dynamics*, 7(3), 111–119. <https://doi.org/10.1007/Bf00211153>
- Büntgen, U., Myglan, V. S., Ljungqvist, F. C., McCormick, M., Di Cosmo, N., Sigl, M., ... Kirdyanov, A. V. (2016). Cooling and societal change during the Late Antique Little Ice Age from 536 to around 660 AD. *Nature Geoscience*, 9, 231. <https://doi.org/10.1038/ngeo2652>
- Büntgen, U., Tegel, W., Nicolussi, K., McCormick, M., Frank, D., Trouet, V., ... Esper, J. (2011). 2500 years of European climate variability and human susceptibility. *Science*, 331(6017), 578–582. <https://doi.org/10.1126/science.1197175>
- Burns, S. J., Matter, A., Frank, N., & Mangini, A. (1998). Speleothem-based paleoclimate record from northern Oman. *Geology*, 26(6), 499–502. [https://doi.org/10.1130/0091-7613\(1998\)026](https://doi.org/10.1130/0091-7613(1998)026)
- Byrd, B. F., Garrard, A. N., & Brandy, P. (2016). Modeling foraging ranges and spatial organization of Late Pleistocene hunter-gatherers in the southern Levant—A least-cost GIS approach. *Quaternary International*, 396, 62–78. <https://doi.org/10.1016/j.quaint.2015.07.048>

- Çağatay, M. N., Ögretmen, N., Damcı, E., Stockhecke, M., Sancar, Ü., Eriş, K. K., & Özeren, S. (2014). Lake level and climate records of the last 90ka from the Northern Basin of Lake Van, eastern Turkey. *Quaternary Science Reviews*, *104*, 97–116. <https://doi.org/10.1016/j.quascirev.2014.09.027>
- Caracuta, V., Weinstein-Evron, M., Yeshurun, R., Kaufman, D., Tsatskin, A., & Boaretto, E. (2016). Charred wood remains in the natufian sequence of el-Wad terrace (Israel): New insights into the climatic, environmental and cultural changes at the end of the Pleistocene. *Quaternary Science Reviews*, *131*, 20–32. <https://doi.org/10.1016/j.quascirev.2015.10.034>
- Catlett, G. A., Rech, J. A., Pigati, J. S., Al Kuisi, M., Li, S., & Honke, J. S. (2017). Activation of a small ephemeral lake in southern Jordan during the last full glacial period and its paleoclimatic implications. *Quaternary Research*, *88*(1), 98–109. <https://doi.org/10.1017/qua.2017.29>
- Cavulli, F., & Scaruffi, S. (2013). Thoughts on nomadism in Middle Holocene Oman. *Arabian Archaeology and Epigraphy*, *24*(1), 15–27. <https://doi.org/10.1111/aae.12023>
- Chaput, M. A., & Gajewski, K. (2016). Radiocarbon dates as estimates of ancient human population size. *Anthropocene*, *15*, 3–12. <https://doi.org/10.1016/j.ancene.2015.10.002>
- Chase, B., Ajithprasad, P., Rajesh, S. V., Patel, A., & Sharma, B. (2014). Materializing Harappan identities: Unity and diversity in the borderlands of the Indus Civilization. *Journal of Anthropological Archaeology*, *35*, 63–78. <https://doi.org/10.1016/j.jaa.2014.04.005>
- Chase, B., Meiggs, D., Ajithprasad, P., & Slater, P. A. (2014). Pastoral land-use of the Indus Civilization in Gujarat: Faunal analyses and biogenic isotopes at Bagasra. *Journal of Archaeological Science*, *50*, 1–15. <https://doi.org/10.1016/j.jas.2014.06.013>
- Chase, B., Meiggs, D., Ajithprasad, P., & Slater, P. A. (2018). What is left behind: Advancing interpretation of pastoral land-use in Harappan Gujarat using herbivore dung to examine biosphere strontium isotope ($87\text{Sr}/86\text{Sr}$) variation. *Journal of Archaeological Science*, *92*, 1–12. <https://doi.org/10.1016/j.jas.2018.01.007>
- Cheng, H., Sinha, A., Verheyden, S., Nader, F. H., Li, X. L., Zhang, P. Z., ... Edwards, R. L. (2015). The climate variability in northern Levant over the past 20,000 years. *Geophysical Research Letters*, *42*(20), 8641–8650. <https://doi.org/10.1002/2015GL065397>
- Clare, L., & Weninger, B. (2014). The dispersal of Neolithic lifeways: Absolute chronology and rapid climate change. *The Neolithic in Turkey*, *6*, 10500–15200.
- Clarke, J., Brooks, N., Banning, E. B., Bar-Matthews, M., Campbell, S., Clare, L., ... Zerboni, A. (2016). Climatic changes and social transformations in the Near East and North Africa during the 'long' 4th millennium BC: A comparative study of environmental and archaeological evidence. *Quaternary Science Reviews*, *136*, 96–121. <https://doi.org/10.1016/j.quascirev.2015.10.003>
- Cleuziou, S., & Tosi, M. (2007). *In the shadow of the ancestors: The prehistoric foundations of the early Arabian civilization in Oman*. Muscat, Oman: Ministry of Heritage & Culture Sultanat of Oman.
- Cohen, K. M., Finney, S. C., Gibbard, P. L., & Fan, J. X. (2013). The ICS International Chronostratigraphic Chart. *Episodes*, *36*(3), 199–204.
- Colledge, S., & Conolly, J. (2010). Reassessing the evidence for the cultivation of wild crops during the Younger Dryas at Tell Abu Hureyra, Syria. *Environmental Archaeology*, *15*(2), 124–138. <https://doi.org/10.1179/146141010X12640787648504>
- Contreras, D. A., & Meadows, J. (2014). Summed radiocarbon calibrations as a population proxy: A critical evaluation using a realistic simulation approach. *Journal of Archaeological Science*, *52*, 591–608. <https://doi.org/10.1016/j.jas.2014.05.030>
- Cook, B. I., Anchukaitis, K. J., Touchan, R., Meko, D. M., & Cook, E. R. (2016). Spatiotemporal drought variability in the Mediterranean over the last 900 years. *Journal of Geophysical Research: Atmospheres*, *121*(5), 2060–2074. <https://doi.org/10.1002/2015JD023929>
- Cook, B. I., Ault, T. R., & Smerdon, J. E. (2015). Unprecedented 21st century drought risk in the American Southwest and Central Plains. *Science Advances*, *1*(1), e1400082. <https://doi.org/10.1126/sciadv.1400082>
- Cook, E. R., Briffa, K. R., Meko, D. M., Graybill, D. A., & Funkhouser, G. (1995). The 'segment length curse' in long tree-ring chronology development for palaeoclimatic studies. *The Holocene*, *5*(2), 229–237. <https://doi.org/10.1177/095968369500500211>
- Cook, E. R., Seager, R., Kushnir, Y., Briffa, K. R., Büntgen, U., Frank, D., ... Zang, C. (2015). Old World megadroughts and pluvials during the Common Era. *Science Advances*, *1*(10), e1500561. <https://doi.org/10.1126/sciadv.1500561>
- Craig, O. E., Saul, H., Lucquin, A., Nishida, Y., Taché, K., Clarke, L., ... Jordan, P. (2013). Earliest evidence for the use of pottery. *Nature*, *496*, 351. <https://doi.org/10.1038/nature12109>
- Crassard, R., & Drechsler, P. (2013). Towards new paradigms: Multiple pathways for the Arabian Neolithic. *Arabian Archaeology and Epigraphy*, *24*(1), 3–8. <https://doi.org/10.1111/aae.12021>
- Crassard, R., Petraglia, M. D., Drake, N. A., Breeze, P., Gratuze, B., Alsharekh, A., ... Schiettecatte, J. (2013). Middle Palaeolithic and Neolithic occupations around Mundafan Palaeolake, Saudi Arabia: Implications for climate change and human dispersals. *PLoS One*, *8*(7), e69665. <https://doi.org/10.1371/journal.pone.0069665>
- Crassard, R., Petraglia, M. D., Parker, A. G., Parton, A., Roberts, R. G., Jacobs, Z., ... Shipton, C. (2013). Beyond the Levant: First evidence of a Pre-Pottery Neolithic incursion into the Nefud Desert, Saudi Arabia. *PLoS One*, *8*(7), e68061. <https://doi.org/10.1371/journal.pone.0068061>
- Cukur, D., Krastel, S., Schmincke, H. U., Sumita, M., Tomonaga, Y., & Namık Çağatay, M. (2014). Water level changes in Lake Van, Turkey, during the past ca. 600 ka: Climatic, volcanic and tectonic controls. *Journal of Paleolimnology*, *52*(3), 201–214. <https://doi.org/10.1007/s10933-014-9788-0>
- Cullen, H. M., deMenocal, P. B., Hemming, S., Hemming, G., Brown, F. H., Guilderson, T., & Sirocko, F. (2000). Climate change and the collapse of the Akkadian empire: Evidence from the deep sea. *Geology*, *28*(4), 379–382. <https://doi.org/10.1130/0091-7613>
- Cuthbert, M. O., Baker, A., Jex, C. N., Graham, P. W., Treble, P. C., Andersen, M. S., & Ian Acworth, R. (2014). Drip water isotopes in semi-arid karst: Implications for speleothem paleoclimatology. *Earth and Planetary Science Letters*, *395*, 194–204. <https://doi.org/10.1016/j.epsl.2014.03.034>
- Cuthbert, M. O., Gleeson, T., Moosdorf, N., Befus, K. M., Schneider, A., Hartmann, J., & Lehner, B. (2019). Global patterns and dynamics of climate–groundwater interactions. *Nature Climate Change*, *9*(2), 137–141. <https://doi.org/10.1038/s41558-018-0386-4>
- Cuthbert, M. O., Gleeson, T., Reynolds, S. C., Bennett, M. R., Newton, A. C., McCormack, C. J., & Ashley, G. M. (2017). Modelling the role of groundwater hydrorefugia in East African hominin evolution and dispersal. *Nature Communications*, *8*, 15696. <https://doi.org/10.1038/ncomms15696>
- Cuthbert, M. O., Rau, G. C., Andersen, M. S., Roshan, H., Rutledge, H., Marjo, C. E., ... Baker, A. (2014). Evaporative cooling of speleothem drip water. *Scientific Reports*, *4*, 5162. <https://doi.org/10.1038/srep05162>
- Cuttler, R. (2013). Considering marine transgression as a mechanism for enforced migration and the littoral Gulf 'Ubaid phenomenon. *Arabian Archaeology and Epigraphy*, *24*(1), 37–43. <https://doi.org/10.1111/aae.12022>
- Daley, T. J., Thomas, E. R., Holmes, J. A., Street-Perrott, F. A., Chapman, M. R., Tindall, J. C., ... Roberts, C. N. (2011). The 8200 yr BP cold event in stable isotope records from the North Atlantic region. *Global and Planetary Change*, *79*(3–4), 288–302. <https://doi.org/10.1016/j.gloplacha.2011.03.006>
- de Miroschedji, P. (2003). Susa and the highlands: Major trends in the history of Elamite civilization. In N. F. Miller, & K. Abdi (Eds.), *Yeki bud, yeki nabud: Essays on the archaeology of Iran in honor of William M. Sumner* (pp. 17–38). Los Angeles, CA: The Cotsen Institute of Archaeology Press.
- Dean, J. R., Eastwood, W. J., Roberts, N., Jones, M. D., Yiğitbaşoğlu, H., Allcock, S. L., ... Leng, M. J. (2015). Tracking the hydro-climatic signal from lake to sediment: A field study from central Turkey. *Journal of Hydrology*, *529*, 608–621. <https://doi.org/10.1016/j.jhydrol.2014.11.004>
- Dean, J. R., Jones, M. D., Leng, M. J., Metcalfe, S. E., Sloane, H. J., Eastwood, W. J., & Roberts, C. N. (2018). Seasonality of Holocene hydroclimate in the Eastern Mediterranean reconstructed using the oxygen isotope composition of carbonates and diatoms from Lake Nar, central Turkey. *The Holocene*, *28*(2), 267–276. <https://doi.org/10.1177/0959683617721326>

- Dean, J. R., Jones, M. D., Leng, M. J., Noble, S. R., Metcalfe, S. E., Sloane, H. J., ... Roberts, C. N. (2015). Eastern Mediterranean hydroclimate over the late glacial and Holocene, reconstructed from the sediments of Nar lake, central Turkey, using stable isotopes and carbonate mineralogy. *Quaternary Science Reviews*, *124*, 162–174. <https://doi.org/10.1016/j.quascirev.2015.07.023>
- Dean, J. R., Jones, M. D., Leng, M. J., Sloane, H. J., Roberts, C. N., Woodbridge, J., ... Yiğitbaşıoğlu, H. (2013). Palaeo-seasonality of the last two millennia reconstructed from the oxygen isotope composition of carbonates and diatom silica from Nar Gölü, central Turkey. *Quaternary Science Reviews*, *66*, 35–44. <https://doi.org/10.1016/j.quascirev.2012.07.014>
- Dee, S., Emile-Geay, J., Evans, M. N., Allam, A., Steig, E. J., & Thompson, D. M. (2015). PRYSM: An open-source framework for PROXY System Modeling, with applications to oxygen-isotope systems. *Journal of Advances in Modeling Earth Systems*, *7*(3), 1220–1247. <https://doi.org/10.1002/2015MS000447>
- Dermoddy, B. J., de Boer, H. J., Bierkens, M. F. P., Weber, S. L., Wassen, M. J., & Dekker, S. C. (2012). A seesaw in Mediterranean precipitation during the Roman Period linked to millennial-scale changes in the North Atlantic. *Climate of the Past*, *8*(2), 637–651. <https://doi.org/10.5194/cp-8-637-2012>
- Develle, A. L., Gasse, F., Vidal, L., Williamson, D., Demory, F., Van Campo, E., ... Thouveny, N. (2011). A 250ka sedimentary record from a small karstic lake in the Northern Levant (Yammouneh, Lebanon): Paleoclimatic implications. *Palaeogeography, Palaeoclimatology, Palaeoecology*, *305*(1), 10–27. <https://doi.org/10.1016/j.palaeo.2011.02.008>
- Develle, A. L., Herreros, J., Vidal, L., Surssock, A., & Gasse, F. (2010). Controlling factors on a paleo-lake oxygen isotope record (Yammouneh, Lebanon) since the Last Glacial maximum. *Quaternary Science Reviews*, *29*(7–8), 865–886. <https://doi.org/10.1016/j.quascirev.2009.12.005>
- Dixit, Y., Hodell, D. A., Giesche, A., Tandon, S. K., Gázquez, F., Saini, H. S., ... Petrie, C. A. (2018). Intensified summer monsoon and the urbanization of Indus Civilization in northwest India. *Scientific Reports*, *8*(1), 4225. <https://doi.org/10.1038/s41598-018-22504-5>
- Dixit, Y., Hodell, D. A., & Petrie, C. A. (2014). Abrupt weakening of the summer monsoon in northwest India ~4100 yr ago. *Geology*, *42*(4), 339–342. <https://doi.org/10.1130/G35236.1>
- Dixit, Y., Hodell, D. A., Sinha, R., & Petrie, C. A. (2014). Abrupt weakening of the Indian summer monsoon at 8.2 kyr BP. *Earth and Planetary Science Letters*, *391*, 16–23. <https://doi.org/10.1016/j.epsl.2014.01.026>
- Djamali, M., Akhiani, H., Andrieu-Ponel, V., Braconnot, P., Brewer, S., de Beaulieu, J. L., ... Stevens, L. (2010). Indian Summer Monsoon variations could have affected the early-Holocene woodland expansion in the Near East. *The Holocene*, *20*(5), 813–820. <https://doi.org/10.1177/0959683610362813>
- Djamali, M., Biglari, F., Abdi, K., Andrieu-Ponel, V., de Beaulieu, J.-L., Mashkour, M., & Ponel, P. (2011). Pollen analysis of coprolites from a late Pleistocene–Holocene cave deposit (Wezmeh Cave, west Iran): Insights into the late Pleistocene and late Holocene vegetation and flora of the central Zagros Mountains. *Journal of Archaeological Science*, *38*(12), 3394–3401. <https://doi.org/10.1016/j.jas.2011.08.001>
- Djamali, M., Brewer, S., Breckle, S. W., & Jackson, S. T. (2012). Climatic determinism in phytogeographic regionalization: A test from the Irano-Turanian region, SW and Central Asia. *Flora—Morphology, Distribution, Functional Ecology of Plants*, *207*(4), 237–249. <https://doi.org/10.1016/j.flora.2012.01.009>
- Djamali, M., de Beaulieu, J. L., Campagne, P., Andrieu-Ponel, V., Ponel, P., Leroy, S. A. G., & Akhiani, H. (2009). Modern pollen rain–vegetation relationships along a forest-steppe transect in the Golestan National Park, NE Iran. *Review of Palaeobotany and Palynology*, *153*(3), 272–281. <https://doi.org/10.1016/j.revpalbo.2008.08.005>
- Djamali, M., de Beaulieu, J. L., Shah-Hosseini, M., Andrieu-Ponel, V., Ponel, P., Amini, A., ... Brewer, S. (2008). A late Pleistocene long pollen record from Lake Urmia, NW Iran. *Quaternary Research*, *69*(3), 413–420. <https://doi.org/10.1016/j.yqres.2008.03.004>
- Djamali, M., Jones, M. D., Migliore, J., Balatti, S., Fader, M., Contreras, D., ... Weeks, L. (2016). Olive cultivation in the heart of the Persian Achaemenid Empire: New insights into agricultural practices and environmental changes reflected in a late Holocene pollen record from Lake Parishan, SW Iran. *Vegetation History and Archaeobotany*, *25*(3), 255–269. <https://doi.org/10.1007/s00334-015-0545-8>
- Douglass, A. E. (1941). Crossdating in dendrochronology. *Journal of Forestry*, *39*(10), 825–831.
- Downey, S. S., Bocaege, E., Kerig, T., Edinborough, K., & Shennan, S. (2014). The Neolithic demographic transition in Europe: Correlation with juvenility index supports interpretation of the summed calibrated radiocarbon date probability distribution (SCDPD) as a valid demographic proxy. *PLoS One*, *9*(8), e105730. <https://doi.org/10.1371/journal.pone.0105730>
- Düring, B. S. (2010). *The prehistory of Asia Minor: From complex hunter-gatherers to early urban societies*. Cambridge, England: Cambridge University Press.
- Düring, B. S., Glatz, C., & Şerifoğlu, T. E. (2012). The Cide Archaeological Project 2011: Third preliminary report. *Anatolia Antiqua*, *20*(1), 167–175.
- Eastwood, W. J., Leng, M. J., Roberts, N., & Davis, B. (2007). Holocene climate change in the eastern Mediterranean region: A comparison of stable isotope and pollen data from Lake Gölhisar, southwest Turkey. *Journal of Quaternary Science*, *22*(4), 327–341. <https://doi.org/10.1002/jqs.1062>
- Eastwood, W. J., Tibby, J., Roberts, N., Birks, H. J. B., & Lamb, H. F. (2002). The environmental impact of the Minoan eruption of Santorini (Thera): Statistical analysis of palaeoecological data from Golhisar, southwest Turkey. *The Holocene*, *12*(4), 431–444. <https://doi.org/10.1191/0959683602h1557rp>
- Engel, M., Brückner, H., Pint, A., Wellbrock, K., Ginau, A., Voss, P., ... Frenzel, P. (2012). The early Holocene humid period in NW Saudi Arabia—Sediments, microfossils and palaeo-hydrological modelling. *Quaternary International*, *266*, 131–141. <https://doi.org/10.1016/j.quaint.2011.04.028>
- Engel, M., Matter, A., Parker, A. G., Parton, A., Petraglia, M. D., Preston, G. W., & Preusser, F. (2017). Lakes or wetlands? A comment on ‘The middle Holocene climatic records from Arabia: Reassessing lacustrine environments, shift of ITCZ in Arabian Sea, and impacts of the southwest Indian and African monsoons’ by Enzel et al. *Global and Planetary Change*, *148*, 258–267. <https://doi.org/10.1016/j.gloplacha.2016.11.001>
- England, A., Eastwood, W. J., Roberts, C. N., Turner, R., & Haldon, J. F. (2008). Historical landscape change in Cappadocia (central Turkey): A palaeoecological investigation of annually laminated sediments from Nar lake. *The Holocene*, *18*(8), 1229–1245. <https://doi.org/10.1177/0959683608096598>
- Enzel, Y., Arnit, R., Dayan, U., Crouvi, O., Kahana, R., Ziv, B., & Sharon, D. (2008). The climatic and physiographic controls of the eastern Mediterranean over the late Pleistocene climates in the southern Levant and its neighboring deserts. *Global and Planetary Change*, *60*(3–4), 165–192. <https://doi.org/10.1016/j.gloplacha.2007.02.003>
- Enzel, Y., Kushnir, Y., & Quade, J. (2015). The middle Holocene climatic records from Arabia: Reassessing lacustrine environments, shift of ITCZ in Arabian Sea, and impacts of the southwest Indian and African monsoons. *Global and Planetary Change*, *129*, 69–91. <https://doi.org/10.1016/j.gloplacha.2015.03.004>
- Enzel, Y., Quade, J., & Kushnir, Y. (2017). Response to Engel et al. (in press): Lakes or wetlands? A comment on "The middle Holocene climatic records from Arabia: Reassessing lacustrine environments, shift of ITCZ in Arabian Sea, and impacts of the southwest Indian and African monsoons" by Enzel et al. (2015). *Global and Planetary Change*, *148*, 268–271. <https://doi.org/10.1016/j.gloplacha.2016.11.003>
- Evans, M. N., Tolwinski-Ward, S. E., Thompson, D. M., & Anchukaitis, K. J. (2013). Applications of proxy system modeling in high resolution paleoclimatology. *Quaternary Science Reviews*, *76*, 16–28. <https://doi.org/10.1016/j.quascirev.2013.05.024>
- Finkelstein, I., & Piasezky, E. (2007). Radiocarbon dating and Philistine chronology with an addendum on el-Ahwat. *Egypt and the Levant*, *17*, 73–82.
- Finné, M., Holmgren, K., Sundqvist, H. S., Weiberg, E., & Lindblom, M. (2011). Climate in the eastern Mediterranean, and adjacent regions, during the past 6000 years—A review. *Journal of Archaeological Science*, *38*(12), 3153–3173. <https://doi.org/10.1016/j.jas.2011.05.007>
- Fleitmann, D., Burns, S. J., Mangini, A., Mudelsee, M., Kramers, J., Villa, I., ... Matter, A. (2007). Holocene ITCZ and Indian monsoon dynamics recorded in stalagmites from Oman and Yemen (Socotra). *Quaternary Science Reviews*, *26*(1–2), 170–188. <https://doi.org/10.1016/j.quascirev.2006.04.012>
- Fleitmann, D., Burns, S. J., Mudelsee, M., Neff, U., Kramers, J., Mangini, A., & Matter, A. (2003). Holocene forcing of the Indian monsoon recorded in a stalagmite from southern Oman. *Science*, *300*(5626), 1737–1739. <https://doi.org/10.1126/science.1083130>

- Fleitmann, D., Burns, S. J., Neff, U., Mudelsee, M., Mangini, A., & Matter, A. (2004). Palaeoclimatic interpretation of high-resolution oxygen isotope profiles derived from annually laminated speleothems from Southern Oman. *Quaternary Science Reviews*, 23(7), 935–945. <https://doi.org/10.1016/j.quascirev.2003.06.019>
- Fleitmann, D., Cheng, H., Badertscher, S., Edwards, R. L., Mudelsee, M., Gökürk, O. M., ... Tüysüz, O. (2009). Timing and climatic impact of Greenland interstadials recorded in stalagmites from northern Turkey. *Geophysical Research Letters*, 36(19), L19707. <https://doi.org/10.1029/2009GL040050>
- Fleitmann, D., Mudelsee, M., Burns, S. J., Bradley, R. S., Kramers, J., & Matter, A. (2008). Evidence for a widespread climatic anomaly at around 9.2 ka before present. *Paleoceanography*, 23(1). <https://doi.org/10.1029/2007PA001519>
- Flohr, P., Fleitmann, D., Matthews, R., Matthews, W., & Black, S. (2016). Evidence of resilience to past climate change in southwest Asia: Early farming communities and the 9.2 and 8.2 ka events. *Quaternary Science Reviews*, 136, 23–39. <https://doi.org/10.1016/j.quascirev.2015.06.022>
- Flohr, P., Fleitmann, D., Zorita, E., Sadekov, A., Cheng, H., Bosomworth, M., ... Matthews, R. (2017). Late Holocene droughts in the fertile Crescent recorded in a speleothem from northern Iraq. *Geophysical Research Letters*, 44(3), 1528–1536. <https://doi.org/10.1002/2016GL071786>
- Fouache, E., Francfort, H.-P., Cosandey, C., & Adle, C. (2015). La crise d'aridité climatique de la fin du 3ème millénaire av. J.-C., à la lumière des contextes géomorphologique de 3 sites d'Iran Oriental (Bam, Tepe Damghani, Jiroft). In R. Rante (Ed.), *Greater Khorasan: History, geography, archaeology and material culture* (Vol. 29, p. 27). Berlin, Germany: De Gruyter.
- Frumkin, A., Bar-Yosef, O., & Schwarcz, H. P. (2011). Possible paleohydrologic and paleoclimatic effects on hominin migration and occupation of the Levantine Middle Paleolithic. *Journal of Human Evolution*, 60(4), 437–451. <https://doi.org/10.1016/j.jhevol.2010.03.010>
- Fujii, S. (2007). PPNB barrage systems at Wadi Abu Tulayha and Wadi ar-Ruwayshid ash-Sharqi: A preliminary report of the 2006 spring field season of the Jafr Basin prehistoric project, phase 2. *Annual of the Department of Antiquities of Jordan*, 51, 403–427.
- Fujii, S. (2008). Wadi Abu Tulayha: A preliminary report on the 2006 summer fields season of the Jafr Basin prehistoric project, phase 2. *Annual of the Department of Antiquities of Jordan*, 52, 446–478.
- Gambin, B., Andrieu-Ponel, V., Medail, F., Marriner, N., Peyron, O., Montade, V., ... Djarnali, M. (2016). 7300 years of vegetation history and climate for NW Malta: A Holocene perspective. *Climate of the Past*, 12(2), 273–297. <https://doi.org/10.5194/cp-12-273-2016>
- Gasse, F., Vidal, L., Van Campo, E., Demory, F., Develle, A.-L., Tachikawa, K., ... Thouveny, N. (2015). Hydroclimatic changes in northern Levant over the past 400,000 years. *Quaternary Science Reviews*, 111, 1–8. <https://doi.org/10.1016/j.quascirev.2014.12.019>
- Gleick, P. H. (2014). Water, drought, climate change, and conflict in Syria. *Weather Climate and Society*, 6(3), 331–340. <https://doi.org/10.1175/Wcas-D-13-00059.1>
- Gökürk, O. M., Fleitmann, D., Badertscher, S., Cheng, H., Edwards, R. L., Leuenberger, M., ... Kramers, J. (2011). Climate on the southern Black Sea coast during the Holocene: Implications from the Sofular Cave record. *Quaternary Science Reviews*, 30(19), 2433–2445. <https://doi.org/10.1016/j.quascirev.2011.05.007>
- Goldberg, S. L., Lau, H. C. P., Mitrovica, J. X., & Lamy, K. (2016). The timing of the Black Sea flood event: Insights from modeling of glacial isostatic adjustment. *Earth and Planetary Science Letters*, 452, 178–184. <https://doi.org/10.1016/j.epsl.2016.06.016>
- Goring-Morris, N., & Belfer-Cohen, A. (2010). "Great Expectations," or the inevitable collapse of the early Neolithic in the Near East. In M. S. Bandy & J. R. Fox (Eds.), *Becoming villagers: Comparing early village societies* (pp. 62–78). Tucson, AZ: University of Arizona Press.
- Gregg, M. W. (2010). *Organic residue analysis and the first uses of pottery in the ancient Middle East*. Oxford, England: British Archaeological Reports Oxford.
- Gregg, M. W., Banning, E. B., Gibbs, K., & Slater, G. F. (2009). Subsistence practices and pottery use in Neolithic Jordan: Molecular and isotopic evidence. *Journal of Archaeological Science*, 36(4), 937–946. <https://doi.org/10.1016/j.jas.2008.09.009>
- Gregoricka, L. A., & Sheridan, S. G. (2017). Continuity or conquest? A multi-isotope approach to investigating identity in the Early Iron Age of the Southern Levant. *American Journal of Physical Anthropology*, 162(1), 73–89. <https://doi.org/10.1002/ajpa.23086>
- Griffin, D., & Anchukaitis, K. J. (2014). How unusual is the 2012–2014 California drought? *Geophysical Research Letters*, 41(24), 9017–9023. <https://doi.org/10.1002/2014GL062433>
- Griffiths, H. I., Schwab, A., & Stevens, L. R. (2001). Environmental change in southwestern Iran: The Holocene ostracod fauna of Lake Mirabad. *The Holocene*, 11(6), 757–764. <https://doi.org/10.1191/09596830195771>
- Griggs, C., Pearson, C., Manning, S. W., & Lorentzen, B. (2014). A 250-year annual precipitation reconstruction and drought assessment for Cyprus from *Pinus brutia* Ten. tree-rings. *International Journal of Climatology*, 34(8), 2702–2714. <https://doi.org/10.1002/joc.3869>
- Grissino-Mayer, H. D. (1996). A 2129-year reconstruction of precipitation for northwestern New Mexico, USA. In J. S. Dean, D. M. Meko, & T. W. Swetnam (Eds.), *Tree rings, environment and humanity* (pp. 191–204). Tucson, AZ: Radiocarbon.
- Grosman, L. (2013). The Natufian chronological scheme—New insights and their implications, In O. Bar-Yosef, & F. R. Valla (Eds.), *Natufian Foragers in the Levant: Terminal Pleistocene Social Changes in Western Asia* (Vol. 19, pp. 622–637). Oxford, England: Berghahn Books.
- Grosman, L., Munro, N. D., Abadi, I., Boaretto, E., Shaham, D., Belfer-Cohen, A., & Bar-Yosef, O. (2016). Nahal Ein Gev II, a Late Natufian community at the Sea of Galilee. *PLoS One*, 11(1), e0146647. <https://doi.org/10.1371/journal.pone.0146647>
- Groucutt, H. S., & Petraglia, M. D. (2012). The prehistory of the Arabian peninsula: Deserts, dispersals, and demography. *Evolutionary Anthropology*, 21(3), 113–125. <https://doi.org/10.1002/evan.21308>
- Guerrero, E., Naji, S., & Bocquet-Appel, J.-P. (2008). The signal of the Neolithic demographic transition in the Levant. In *The Neolithic demographic transition and its consequences* (pp. 57–80). Dordrecht, Netherlands: Springer.
- Hartman, G., Bar-Yosef, O., Brittingham, A., Grosman, L., & Munro, N. D. (2016). Hunted gazelles evidence cooling, but not drying, during the Younger Dryas in the southern Levant. *Proceedings of the National Academy of Sciences*, 113(15), 3997–4002. <https://doi.org/10.1073/pnas.1519862113>
- Hassan, F. A. (1978). Demographic Archaeology. In M. B. Schiffer (Ed.), *Advances in Archaeological Method and Theory* (Vol. 1, pp. 49–103). London, England: Academic Press.
- Hershkovitz, I., Weber, G. W., Quam, R., Duval, M., Grün, R., Kinsley, L., ... Weinstein-Evron, M. (2018). The earliest modern humans outside Africa. *Science*, 359(6374), 456–459. <https://doi.org/10.1126/science.aap8369>
- Herweijer, C., Seager, R., Cook, E. R., & Emile-Geay, J. (2007). North American droughts of the last millennium from a gridded network of tree-ring data. *Journal of Climate*, 20(7), 1353–1376. <https://doi.org/10.1175/jcli4042.1>
- Heyvaert, V. M. A., & Baeteman, C. (2007). Holocene sedimentary evolution and palaeocoastlines of the Lower Khuzestan plain (southwest Iran). *Marine Geology*, 242(1–3), 83–108. <https://doi.org/10.1016/j.margeo.2007.01.008>
- Hilbert, Y. H., White, T. S., Parton, A., Clark-Balzan, L., Crassard, R., Groucutt, H. S., ... Petraglia, M. D. (2014). Epipalaeolithic occupation and palaeoenvironments of the southern Nefud desert, Saudi Arabia, during the terminal Pleistocene and early Holocene. *Journal of Archaeological Science*, 50, 460–474. <https://doi.org/10.1016/j.jas.2014.07.023>
- Hillman, G., Hedges, R., Moore, A., Colledge, S., & Pettiit, P. (2001). New evidence of Late glacial cereal cultivation at Abu Hureyra on the Euphrates. *The Holocene*, 11(4), 383–393. <https://doi.org/10.1191/095968301678302823>
- Højlund, F. (2007). *The burial mounds of Bahrain*. Moesgaard, Denmark: Jutland Archaeological Society.
- Hole, F. (1987). *The archaeology of western Iran: Settlement and society from prehistory to the Islamic conquest* (Vol. 27). Washington, DC: Smithsonian Institution Press.

- Jarihani, A. A., Larsen, J. R., Callow, J. N., McVicar, T. R., & Johansen, K. (2015). Where does all the water go? Partitioning water transmission losses in a data-sparse, multi-channel and low-gradient dryland river system using modelling and remote sensing. *Journal of Hydrology*, 529, 1511–1529. <https://doi.org/10.1016/j.jhydrol.2015.08.030>
- Jennings, R. P., Shipton, C., Al-Omari, A., Alsharekh, A. M., Crassard, R., Groucutt, H., & Petraglia, M. D. (2015). Rock art landscapes beside the Jubbah palaeolake, Saudi Arabia. *Antiquity*, 87(337), 666–683. <https://doi.org/10.1017/S0003598X00049383>
- Jex, C. N., Baker, A., Eden, J. M., Eastwood, W. J., Fairchild, I. J., Leng, M. J., ... Sloane, H. J. (2011). A 500 yr speleothem-derived reconstruction of late autumn–winter precipitation, northeast Turkey. *Quaternary Research*, 75(3), 399–405. <https://doi.org/10.1016/j.yqres.2011.01.005>
- Jex, C. N., Phipps, S. J., Baker, A., & Bradley, C. (2013). Reducing uncertainty in the climatic interpretations of speleothem $\delta^{18}\text{O}$. *Geophysical Research Letters*, 40(10), 2259–2264. <https://doi.org/10.1002/grl.50467>
- Jones, M. D. (2013). What do we mean by wet? Geoarchaeology and the reconstruction of water availability. *Quaternary International*, 308-309, 76–79. <https://doi.org/10.1016/j.quaint.2012.12.015>
- Jones, M. D., Djamali, M., Holmes, J., Weeks, L., Leng, M. J., Lashkari, A., ... Metcalfe, S. E. (2015). Human impact on the hydroenvironment of Lake Parishan, SW Iran, through the late-Holocene. *The Holocene*, 25(10), 1651–1661. <https://doi.org/10.1177/0959683615594242>
- Jones, M. D., & Imbers, J. (2010). Modeling Mediterranean lake isotope variability. *Global and Planetary Change*, 71(3-4), 193–200. <https://doi.org/10.1016/j.gloplacha.2009.10.001>
- Jones, M. D., Maher, L. A., Macdonald, D. A., Ryan, C., Rambeau, C., Black, S., & Richter, T. (2016). The environmental setting of Epipalaeolithic aggregation site Kharaneh IV. *Quaternary International*, 396, 95–104. <https://doi.org/10.1016/j.quaint.2015.08.092>
- Jones, M. D., Maher, L. A., Richter, T., Macdonald, D. A., & Martin, L. (2016). Human–environment interactions through the epipalaeolithic of eastern Jordan. In D. Contreras (Ed), *Archaeology of human–environment interactions*. London, England: Routledge.
- Jones, M. D., & Roberts, C. N. (2008). Interpreting lake isotope records of Holocene environmental change in the Eastern Mediterranean. *Quaternary International*, 181, 32–38. <https://doi.org/10.1016/j.quaint.2007.01.012>
- Jones, M. D., Roberts, C. N., & Leng, M. J. (2007). Quantifying climatic change through the last glacial-interglacial transition based on lake isotope palaeohydrology from central Turkey. *Quaternary Research*, 67(3), 463–473. <https://doi.org/10.1016/j.yqres.2007.01.004>
- Jones, M. D., Roberts, C. N., Leng, M. J., & Turkes, M. (2006). A high-resolution late Holocene lake isotope record from Turkey and links to North Atlantic and monsoon climate. *Geology*, 34(5), 361–364. <https://doi.org/10.1130/G22407.1>
- Jones, P. J. (2018). *Climate change, water stress and agriculture in the Indus Civilisation, 3000–1500 BC*. Cambridge, England: University of Cambridge.
- Jothityangkoon, C., Sivapalan, M., & Farmer, D. L. (2001). Process controls of water balance variability in a large semi-arid catchment: Downward approach to hydrological model development. *Journal of Hydrology*, 254(1), 174–198. [https://doi.org/10.1016/S0022-1694\(01\)00496-6](https://doi.org/10.1016/S0022-1694(01)00496-6)
- Kadosh, D., Sivan, D., Kutiel, H., & Weinstein-Evron, M. (2004). A late quaternary paleoenvironmental sequence from Dor, Carmel coastal plain, Israel. *Palynology*, 28(1), 143–157.
- Kaniewski, D., Marriner, N., Cheddadi, R., Guiot, J., & Van Campo, E. (2018). The 4.2 ka BP event in the Levant. *Climate of the Past*, 14(10), 1529–1542. <https://doi.org/10.5194/cp-14-1529-2018>
- Kaniewski, D., Paulissen, E., Van Campo, E., Weiss, H., Otto, T., Bretschneider, J., & Van Lerberghe, K. (2010). Late second-early first millennium BC abrupt climate changes in coastal Syria and their possible significance for the history of the Eastern Mediterranean. *Quaternary Research*, 74(2), 207–215. <https://doi.org/10.1016/j.yqres.2010.07.010>
- Kaniewski, D., Van Campo, E., & Weiss, H. (2012). Drought is a recurring challenge in the Middle East. *Proceedings of the National Academy of Sciences*, 109(10), 3862–3867. <https://doi.org/10.1073/pnas.1116304109>
- Kelley, C. P., Mohtadi, S., Cane, M. A., Seager, R., & Kushnir, Y. (2015). Climate change in the Fertile Crescent and implications of the recent Syrian drought. *Proceedings of the National Academy of Sciences of the United States of America*, 112(11), 3241–3246. <https://doi.org/10.1073/pnas.1421533112>
- Kenoyer, J. M., Price, T. D., & Burton, J. H. (2013). A new approach to tracking connections between the Indus Valley and Mesopotamia: Initial results of strontium isotope analyses from Harappa and Ur. *Journal of Archaeological Science*, 40(5), 2286–2297. <https://doi.org/10.1016/j.jas.2012.12.040>
- Kılınc, G. M., Koptekin, D., Atakuman, Ç., Sümer, A. P., Dönertaş, H. M., Yaka, R., ... Somel, M. (2017). Archaeogenomic analysis of the first steps of Neolithization in Anatolia and the Aegean. *Proceedings of the Royal Society B: Biological Sciences*, 284(1867), 20172064. <https://doi.org/10.1098/rspb.2017.2064>
- Kotthoff, U., Pross, J., Müller, U. C., Peyron, O., Schmiedl, G., Schulz, H., & Bordon, A. (2008). Climate dynamics in the borderlands of the Aegean Sea during formation of sapropel S1 deduced from a marine pollen record. *Quaternary Science Reviews*, 27(7–8), 832–845. <https://doi.org/10.1016/j.quascirev.2007.12.001>
- Kuijt, I., Finlayson, B., & MacKay, J. (2015). Pottery Neolithic landscape modification at Dhra'. *Antiquity*, 81(311), 106–118. <https://doi.org/10.1017/S0003598X00094874>
- Kumar, M. (2009). Harappan settlements in the Ghaggar–Yamuna divide. *Linguistics, Archaeology and the Human Past*, 7, 1–75.
- Kutterer, A., & Uerpmann, H.-P. (2012). Social implications and potential causes of violence at Neolithic Al-Buhais 18. In D. A. Potts & P. Hellyer (Eds.), *Fifty years of emirates archaeology* (pp. 54–62). UAE: Ministry of Culture, Youth and Community Development.
- Kuzucuoğlu, C. (2009). Climate and environment in times of cultural changes from the 4th to the 1st mill. BC in the Near and Middle East. *Scienze dell'Antichità*, 15(15), 193–216.
- Kuzucuoğlu, C. (2015). The rise and fall of the Hittite State in central Anatolia: How, when, where, did climate intervene? In D. Beyer & H. A. Tibet (Eds.), *Archaeology of southern Cappadocia, Proceedings of the 3rd Archaeology Meeting of the IFEA* (pp. 17–41). Istanbul, Turkey: Anatolica Varia.
- Kuzucuoğlu, C., Christol, A., Mouralis, D., Doğu, A.-F., Akköprü, E., Fort, M., ... Guillou, H. (2010). Formation of the Upper Pleistocene terraces of Lake Van (Turkey). *Journal of Quaternary Science*, 25(7), 1124–1137. <https://doi.org/10.1002/jqs.1431>
- Kuzucuoğlu, C., Dörfler, W., Kunesch, S., & Goupille, F. (2011). Mid- to late-Holocene climate change in central Turkey: The Tecer Lake record. *The Holocene*, 21(1), 173–188. <https://doi.org/10.1177/0959683610384163>
- Kwicien, O., Stockhecke, M., Pickarski, N., Heumann, G., Litt, T., Sturm, M., ... Haug, G. H. (2014). Dynamics of the last four glacial terminations recorded in Lake Van, Turkey. *Quaternary Science Reviews*, 104, 42–52. <https://doi.org/10.1016/j.quascirev.2014.07.001>
- Lawrence, D., Philip, G., Hunt, H., Snape-Kennedy, L., & Wilkinson, T. J. (2016). Long term population, city size and climate trends in the fertile crescent: A first approximation. *PLoS One*, 11(3), e0152563. <https://doi.org/10.1371/journal.pone.0152563>
- Lechterbeck, J., Edinborough, K., Kerig, T., Fyfe, R., Roberts, N., & Shennan, S. (2014). Is Neolithic land use correlated with demography? An evaluation of pollen-derived land cover and radiocarbon-inferred demographic change from Central Europe. *The Holocene*, 24(10), 1297–1307. <https://doi.org/10.1177/0959683614540952>
- Leroy, S. A. G., & Arpe, K. (2007). Glacial refugia for summer-green trees in Europe and south-west Asia as proposed by ECHAM3 time-slice atmospheric model simulations. *Journal of Biogeography*, 34(12), 2115–2128. <https://doi.org/10.1111/j.1365-2699.2007.01754.x>
- Leroy, S. A. G., Lahijani, H. A. K., Djamali, M., Naqinezhad, A., Moghadam, M. V., Arpe, K., ... Naderi Beni, M. (2011). Late Little Ice Age palaeoenvironmental records from the Anzali and Amirkola Lagoons (south Caspian Sea): Vegetation and sea level changes. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 302(3), 415–434. <https://doi.org/10.1016/j.palaeo.2011.02.002>

- Lézine, A.-M., Robert, C., Cleuziou, S., Inizan, M.-L., Braemer, F., Saliège, J.-F., ... Steimer-Herbet, T. (2010). Climate change and human occupation in the southern Arabian lowlands during the last deglaciation and the Holocene. *Global and Planetary Change*, 72(4), 412–428. <https://doi.org/10.1016/j.gloplacha.2010.01.016>
- Lionello, P., Malanotte-Rizzoli, P., Boscolo, R., Alpert, P., Artale, V., Li, L., ... Xoplaki, E. (2006). The Mediterranean climate: An overview of the main characteristics and issues. In P. Lionello, P. Malanotte-Rizzoli, & R. Boscolo (Eds.), *Developments in Earth and environmental sciences* (Vol. 4, pp. 1–26). Elsevier.
- Litt, T., & Anselmetti, F. S. (2014). Lake Van deep drilling project PALEOVAN. *Quaternary Science Reviews*, 104, 1–7. <https://doi.org/10.1016/j.quascirev.2014.09.026>
- Litt, T., Ohlwein, C., Neumann, F. H., Hense, A., & Stein, M. (2012). Holocene climate variability in the Levant from the Dead Sea pollen record. *Quaternary Science Reviews*, 49, 95–105. <https://doi.org/10.1016/j.quascirev.2012.06.012>
- Litt, T., Pickarski, N., Heumann, G., Stockhecke, M., & Tzedakis, P. C. (2014). A 600,000 year long continental pollen record from Lake Van, eastern Anatolia (Turkey). *Quaternary Science Reviews*, 104, 30–41. <https://doi.org/10.1016/j.quascirev.2014.03.017>
- Luterbacher, J., García-Herrera, R., Akcer-On, S., Allan, R., Alvarez-Castro, M.-C., Benito, G., ... Zorita, E. (2012). A review of 2000 years of Paleoclimatic evidence in the Mediterranean. In P. Lionello (Ed.), *The climate of the Mediterranean region* (pp. 87–185). Oxford, England: Elsevier.
- Madella, M., & Fuller, D. Q. (2006). Palaeoecology and the Harappan civilisation of South Asia: A reconsideration. *Quaternary Science Reviews*, 25(11), 1283–1301. <https://doi.org/10.1016/j.quascirev.2005.10.012>
- Magee, P. (2014). *The archaeology of prehistoric Arabia: Adaptation and social formation from the Neolithic to the Iron Age*. Cambridge, England: Cambridge University Press.
- Maher, L. A. (2010). The Late Pleistocene of Arabia in relation to the Levant. In P. Petraglia, & J. I. Rose (Ed.), *The evolution of human populations in Arabia* (pp. 187–202). Dordrecht, Netherlands: Springer Netherlands.
- Maher, L. A., Banning, E. B., & Chazan, M. (2011). Oasis or Mirage? Assessing the role of abrupt climate change in the prehistory of the southern Levant. *Cambridge Archaeological Journal*, 21(01), 1–30. <https://doi.org/10.1017/S0959774311000011>
- Markowska, M., Baker, A., Andersen, M. S., Jex, C. N., Cuthbert, M. O., Rau, G. C., ... Edwards, N. (2016). Semi-arid zone caves: Evaporation and hydrological controls on $\delta^{18}\text{O}$ drip water composition and implications for speleothem paleoclimate reconstructions. *Quaternary Science Reviews*, 131, 285–301. <https://doi.org/10.1016/j.quascirev.2015.10.024>
- Martin, D. L. (2007). Bioarchaeology in the United Arab Emirates. *Arabian Archaeology and Epigraphy*, 18(2), 124–131. <https://doi.org/10.1111/j.1600-0471.2007.00285.x>
- Massuel, S., Cappelaere, B., Favreau, G., Leduc, C., Lebel, T., & Vischel, T. (2011). Integrated surface water–groundwater modelling in the context of increasing water reserves of a regional Sahelian aquifer. *Hydrological Sciences Journal*, 56(7), 1242–1264. <https://doi.org/10.1080/02626667.2011.609171>
- Matter, A., Mahjoub, A., Neubert, E., Preusser, F., Schwab, A., Szidat, S., & Wulf, G. (2016). Reactivation of the Pleistocene trans-Arabian Wadi ad Dawasir fluvial system (Saudi Arabia) during the Holocene humid phase. *Geomorphology*, 270, 88–101. <https://doi.org/10.1016/j.geomorph.2016.07.013>
- Matthews, A., Ayalon, A., & Bar-Matthews, M. (2000). D/H ratios of fluid inclusions of Soreq cave (Israel) speleothems as a guide to the eastern Mediterranean Meteoric Line relationships in the last 120 ky. *Chemical Geology*, 166(3), 183–191. [https://doi.org/10.1016/S0009-2541\(99\)00192-8](https://doi.org/10.1016/S0009-2541(99)00192-8)
- Matthews, R. (2013). *The archaeology of Mesopotamia: Theories and approaches*. London, England: Routledge.
- McCormick, M., Büntgen, U., Cane, M. A., Cook, E. R., Harper, K., Huybers, P., ... Tegel, W. (2012). Climate change during and after the Roman Empire: Reconstructing the past from scientific and historical evidence. *Journal of Interdisciplinary History*, 43(2), 169–220. https://doi.org/10.1162/JINH_a_00379
- McCorriston, J., Harrower, M., Martin, L., & Oches, E. (2012). Cattle Cults of the Arabian Neolithic and early territorial societies. *American Anthropologist*, 114(1), 45–63. <https://doi.org/10.1111/j.1548-1433.2011.01396.x>
- McSweeney, K., Méry, S., & al Tikriti, W. Y. (2010). Life and death in an Early Bronze Age community from Hili, Al Ain, UAE. In *Death and burial in Arabia and beyond: Multidisciplinary perspectives* (pp. 45–54). Oxford, England: Archaeopress.
- Meadows, J. (2005). The younger dryas episode and the radiocarbon chronologies of the Lake Huleh and Ghab Valley pollen diagrams, Israel and Syria. *The Holocene*, 15(4), 631–636. <https://doi.org/10.1191/0959683605h1838fa>
- Meister, J., Krause, J., Müller-Neuhof, B., Portillo, M., Reimann, T., & Schütt, B. (2017). Desert agricultural systems at EBA Jawa (Jordan): Integrating archaeological and paleoenvironmental records. *Quaternary International*, 434, 33–50. <https://doi.org/10.1016/j.quaint.2015.12.086>
- Melvin, T. M., & Briffa, K. R. (2008). A “signal-free” approach to dendroclimatic standardisation. *Dendrochronologia*, 26(2), 71–86.
- Miebach, A., Niestrath, P., Roesser, P., & Litt, T. (2016). Impacts of climate and humans on the vegetation in northwestern Turkey: Palynological insights from Lake Iznik since the Last Glacial. *Climate of the Past*, 12(2), 575–593. <https://doi.org/10.5194/cp-12-575-2016>
- Migowski, C., Stein, M., Prasad, S., Negendank, J. F. W., & Agnon, A. (2006). Holocene climate variability and cultural evolution in the Near East from the Dead Sea sedimentary record. *Quaternary Research*, 66(3), 421–431. <https://doi.org/10.1016/j.yqres.2006.06.010>
- Moore, A. M. T., & Hillman, G. C. (1992). The Pleistocene to Holocene transition and human economy in southwest Asia—The impact of the Younger Dryas. *American Antiquity*, 57(3), 482–494. <https://doi.org/10.2307/280936>
- Nadi, M., Bazrafshan, J., Pourtahmasi, K., & Bräuning, A. (2017). Tree-ring based reconstruction of the joint deficit index in Javan-Roud Region, Kermanshah (Iran). *International Journal of Climatology*, 37(1), 420–429. <https://doi.org/10.1002/joc.4715>
- Neff, U., Burns, S. J., Mangini, A., Mudelsee, M., Fleitmann, D., & Matter, A. (2001). Strong coherence between solar variability and the monsoon in Oman between 9 and 6 kyr ago. *Nature*, 411, 290. <https://doi.org/10.1038/35077048>
- Neugebauer, I., Brauer, A., Schwab, M. J., Waldmann, N. D., Enzel, Y., Kitagawa, H., ... Party, D. S. (2014). Lithology of the long sediment record recovered by the ICDP Dead Sea Deep Drilling Project (DSDDP). *Quaternary Science Reviews*, 102, 149–165. <https://doi.org/10.1016/j.quascirev.2014.08.013>
- Nishiaki, Y., & Le Miere, M. (2005). The oldest pottery Neolithic of Upper Mesopotamia: New evidence from Tell Seker al-Aheimar, the Khabor, northeast Syria. *Paleorient*, 31(2), 55–68.
- Orland, I. J., Bar-Matthews, M., Kita, N. T., Ayalon, A., Matthews, A., & Valley, J. W. (2009). Climate deterioration in the eastern Mediterranean as revealed by ion microprobe analysis of a speleothem that grew from 2.2 to 0.9 ka in Soreq Cave, Israel. *Quaternary Research*, 71(1), 27–35. <https://doi.org/10.1016/j.yqres.2008.08.005>
- Pachauri, R. K., Allen, M. R., Barros, V. R., Broome, J., Cramer, W., Christ, R., ... Dasgupta, P. (2014). Climate change 2014: Synthesis report. Contribution of Working Groups I, II and III to the fifth assessment report of the Intergovernmental Panel on Climate Change. IPCC.
- Parker, A. G., Eckersley, L., Smith, M. M., Goudie, A. S., Stokes, S., Ward, S., ... Hodson, M. J. (2004). Holocene vegetation dynamics in the northeastern Rub' al-Khali desert, Arabian Peninsula: A phytolith, pollen and carbon isotope study. *Journal of Quaternary Science*, 19(7), 665–676. <https://doi.org/10.1002/jqs.880>
- Parker, A. G., Goudie, A. S., Stokes, S., White, K., Hodson, M. J., Manning, M., & Kennet, D. (2006). A record of Holocene climate change from lake geochemical analyses in southeastern Arabia. *Quaternary Research*, 66(3), 465–476. <https://doi.org/10.1016/j.yqres.2006.07.001>
- Parker, A. G., Preston, G. W., Parton, A., Walkington, H., Jardine, P. E., Leng, M. J., & Hodson, M. J. (2016). Low-latitude Holocene hydroclimate derived from lake sediment flux and geochemistry. *Journal of Quaternary Science*, 31(4), 286–299. <https://doi.org/10.1002/jqs.2859>
- Peltenburg, E. (2012). East Mediterranean water wells of the 9th–7th millennium BC. In F. Klimscha, R. Eichmann, C. Schuler, & H. adn Fahlbusch (Eds.), *Wasserwirtschaftliche Innovationen im archäologischen Kontext Von den prähistorischen Anfängen bis zu den Metropolen der Antike Herausgegeben* (pp. 69–82). Rahden, Germany: Marie Leidorf.

- Perez-Sanz, A., Li, G., Gonzalez-Samperiz, P., & Harrison, S. P. (2014). Evaluation of modern and mid-Holocene seasonal precipitation of the Mediterranean and northern Africa in the CMIP5 simulations. *Climate of the Past*, 10(2), 551–568. <https://doi.org/10.5194/cp-10-551-2014>
- Petrie, C. (2017). Crisis, what crisis? Adaptation, resilience and transformation in the Indus Civilisation. In T. Cunningham, & J. Driessen (Eds.), *Crisis to collapse: The archaeology of social breakdown* (pp. 59–80). Louvain, Belgium: UCL University Press.
- Petrie, C., Weeks, L., Potts, D., & Roustaei, K. (2006). Perspectives on the cultural sequence of Mamasani. The Mamasani Archaeological Project Stage One: A report on the first two seasons of the ICAR–University of Sydney Joint Expedition to the Mamasani District, Fars Province, Iran, Iranian Center for Archaeological Research, Tehran, pp. 169–196.
- Petrie, C. A., Singh, R. N., Bates, J., Dixit, Y., French, C. A. I., Hodell, D. A., ... Singh, D. P. (2017). Adaptation to variable environments, resilience to climate change: Investigating land, water and settlement in Indus northwest India. *Current Anthropology*, 58(1), 1–30. <https://doi.org/10.1086/690112>
- Petrie, C. A., & Thomas, K. D. (2015). The topographic and environmental context of the earliest village sites in western South Asia. *Antiquity*, 86(334), 1055–1067. <https://doi.org/10.1017/S0003598X00048249>
- Pigati, J. S., Rech, J. A., Quade, J., & Bright, J. (2014). Desert wetlands in the geologic record. *Earth-Science Reviews*, 132, 67–81. <https://doi.org/10.1016/j.earscirev.2014.02.001>
- Ponel, P., Andrieu-Ponel, V., Djmal, M., Lahijani, H., Leydet, M., & Mashkour, M. (2013). Fossil beetles as possible evidence for transhumance during the middle and late Holocene in the high mountains of Talysh (Talesh) in NW Iran? *Environmental Archaeology*, 18(3), 201–210. <https://doi.org/10.1179/1749631413Y.0000000007>
- Pons, A., & Reille, M. (1988). The holocene- and upper pleistocene pollen record from Padul (Granada, Spain): A new study. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 66(3), 243–263. [https://doi.org/10.1016/0031-0182\(88\)90202-7](https://doi.org/10.1016/0031-0182(88)90202-7)
- Possehl, G. L. (1999). *Indus Age: The beginnings* (Vol. 1). Philadelphia, PA: University of Pennsylvania Press.
- Potts, D. T. (2013). *The Oxford handbook of ancient Iran*. New York, NY: Oxford University Press.
- Pournelle, J. R., Algaze, G., Crawford, H., McMahon, A., & Postgate, N. (2012). Travels in Edin: Deltaic resilience and early urbanism in greater Mesopotamia. In A. McMahon, & H. Crawford (Eds.), *Preludes to urbanism: Studies in the Late Chalcolithic of Mesopotamia in honour of Joan Oates*. Cambridge, England: McDonald Institute for Archaeological Research.
- Prasad, S., Anoop, A., Riedel, N., Sarkar, S., Menzel, P., Basavaiah, N., ... Stebich, M. (2014). Prolonged monsoon droughts and links to Indo-Pacific warm pool: A Holocene record from Lonar Lake, central India. *Earth and Planetary Science Letters*, 391, 171–182. <https://doi.org/10.1016/j.epsl.2014.01.043>
- Preston, G. W., Thomas, D. S. G., Goudie, A. S., Atkinson, O. A. C., Leng, M. J., Hodson, M. J., ... Parker, A. G. (2015). A multi-proxy analysis of the Holocene humid phase from the United Arab Emirates and its implications for southeast Arabia's Neolithic populations. *Quaternary International*, 382, 277–292. <https://doi.org/10.1016/j.quaint.2015.01.054>
- Prickett, M. E. (1986). *Man, land, and water: Settlement distribution and the development of irrigation agriculture in the upper Rud-i Gushk drainage, southeastern Iran* (Vol. I-III). Ann Arbor, MI: University of Michigan.
- Rambeau, C., Finlayson, B., Smith, S., Black, S., Inglis, R., & Robinson, S. (2011). Palaeoenvironmental reconstruction at Beidha, southern Jordan (c. 18,000–8,500 BP): Implications for human occupation during the Natufian and Pre-Pottery Neolithic. In S. Mithen, & E. Black (Eds.), *Water, life and civilisation: Climate, environment and society in the Jordan Valley* (pp. 245–268). Cambridge, England: Cambridge University Press.
- Rasmussen, S. O., Vinther, B. M., Clausen, H. B., & Andersen, K. K. (2007). Early Holocene climate oscillations recorded in three Greenland ice cores. *Quaternary Science Reviews*, 26(15), 1907–1914. <https://doi.org/10.1016/j.quascirev.2007.06.015>
- Reba, M., Reitsma, F., & Seto, K. C. (2016). Spatializing 6,000 years of global urbanization from 3700 BC to AD 2000. *Scientific Data*, 3, 160034. <https://doi.org/10.1038/sdata.2016.34>
- Regev, J., Finkelstein, I., Adams, M. J., & Boaretto, E. (2014). Wiggle-matched 14C chronology of Early Bronze Megiddo and the synchronisation of Egyptian and Levantine chronologies. *Egypt and the Levant*, 24, 241–264.
- Reuter, J., Buening, N., & Yoshimura, K. (2017). Evaluating hydrological influences on mid-latitude $\delta^{18}\text{O}_p$ in the Middle East. *Climate Dynamics*, 50, 3153–3170. <https://doi.org/10.1007/s00382-017-3798-3>
- Richter, T. (2016). The role of water in the last gatherer–hunter and earliest farming communities of southwest Asia. *Water of Life*, 62–75.
- Richter, T. (2017). Natufian and early Neolithic in the Black Desert quaternary environments. In *Climate change and humans in the Levant* (pp. 715–722). Cambridge, England: Cambridge University Press.
- Richter, T., Arranz-Otaegui, A., Yeomans, L., & Boaretto, E. (2017). High resolution AMS dates from Shubayqa 1, northeast Jordan reveal complex origins of Late Epipalaeolithic Natufian in the Levant. *Scientific Reports*, 7(1), 17025. <https://doi.org/10.1038/s41598-017-17096-5>
- Rick, J. W. (1987). Dates as data: An examination of the Peruvian preceramic radiocarbon record. *American Antiquity*, 52(1), 55–73. <https://doi.org/10.2307/281060>
- Riehl, S., Pustovoytov, K. E., Weippert, H., Klett, S., & Hole, F. (2014). Drought stress variability in ancient Near Eastern agricultural systems evidenced by $\delta^{13}\text{C}$ in barley grain. *Proceedings of the National Academy of Sciences*, 111(34), 12348–12353. <https://doi.org/10.1073/pnas.1409516111>
- Risi, C., Bony, S., Vimeux, F., & Jouzel, J. (2010). Water-stable isotopes in the LMDZ4 general circulation model: Model evaluation for present-day and past climates and applications to climatic interpretations of tropical isotopic records. *Journal of Geophysical Research: Atmospheres*, 115(D12), D12118. <https://doi.org/10.1029/2009JD013255>
- Roberts, C. N., & Reed, J. M. (2009). Mediterranean lakes, wetlands and Holocene environmental change. In J. C. Woodward (Ed.), *The physical geography of the Mediterranean* (pp. 255–286). Oxford, England: Oxford University Press.
- Roberts, C. N., & Wright, H. E. (1993). Vegetational, lake-level, and climatic history of the Near East and Southwest Asia. In H. E. Wright, J. E. Kutzbach, T. Webb, W. F. Ruddiman, F. A. Street-Perrott, & P. J. Bartlein (Eds.), *Global climates since the Last Glacial maximum*. Minneapolis, MN: University of Minnesota Press.
- Roberts, N. (1983). Age, paleoenvironments, and climatic significance of Late Pleistocene Konya Lake, Turkey. *Quaternary Research*, 19(2), 154–171. [https://doi.org/10.1016/0033-5894\(83\)90002-9](https://doi.org/10.1016/0033-5894(83)90002-9)
- Roberts, N. (2002). Did prehistoric landscape management retard the post-glacial spread of woodland in southwest Asia? *Antiquity*, 76(294), 1002–1010. <https://doi.org/10.1017/S0003598X0009181X>
- Roberts, N., Allcock, S. L., Arnaud, F., Dean, J. R., Eastwood, W. J., Jones, M. D., ... Yiğitbaşıoğlu, H. (2016). A tale of two lakes: A multi-proxy comparison of Late Glacial and Holocene environmental change in Cappadocia, Turkey. *Journal of Quaternary Science*, 31(4), 348–362. <https://doi.org/10.1002/jqs.2852>
- Roberts, N., Brayshaw, D., Kuzucuoglu, C., Perez, R., & Sadori, L. (2011). The mid-Holocene climatic transition in the Mediterranean: Causes and consequences. *The Holocene*, 21(1), 3–13. <https://doi.org/10.1177/0959683610388058>
- Roberts, N., Jones, M. D., Benkaddour, A., Eastwood, W. J., Filippi, M. L., Frogley, M. R., ... Zanchetta, G. (2008). Stable isotope records of Late Quaternary climate and hydrology from Mediterranean lakes: The ISOMED synthesis. *Quaternary Science Reviews*, 27(25–26), 2426–2441. <https://doi.org/10.1016/j.quascirev.2008.09.005>
- Roberts, N., Moreno, A., Valero-Garcés, B. L., Corella, J. P., Jones, M., Allcock, S., ... Türkeş, M. (2012). Palaeolimnological evidence for an east–west climate seesaw in the Mediterranean since AD 900. *Global and Planetary Change*, 84–85, 23–34. <https://doi.org/10.1016/j.gloplacha.2011.11.002>
- Roberts, N., Reed, J. M., Leng, M. J., Kuzucuoglu, C., Fontugne, M., Bertaux, J., ... Karabiyikoglu, M. (2001). The tempo of Holocene climatic change in the eastern Mediterranean region: New high-resolution crater-lake sediment data from central Turkey. *The Holocene*, 11(6), 721–736. <https://doi.org/10.1191/09596830195744>

- Roberts, N., Woodbridge, J., Bevan, A., Palmisano, A., Shennan, S., & Asouti, E. (2018). Human responses and non-responses to climatic variations during the last Glacial–Interglacial transition in the eastern Mediterranean. *Quaternary Science Reviews*, *184*, 47–67. <https://doi.org/10.1016/j.quascirev.2017.09.011>
- Robinson, S. A., Black, S., Sellwood, B. W., & Valdes, P. J. (2006). A review of palaeoclimates and palaeoenvironments in the Levant and Eastern Mediterranean from 25,000 to 5000 years BP: Setting the environmental background for the evolution of human civilisation. *Quaternary Science Reviews*, *25*(13–14), 1517–1541. <https://doi.org/10.1016/j.quascirev.2006.02.006>
- Rodríguez, A., Haidar-Boustani, M., Urquijo, J. E. G., Ibáñez, J. J., Al-Maqdissi, M., Terradas, X., & Zapata, L. (2013). The Early Natufian site of Jeftelek (Homs Gap, Syria). In O. Bar-Yosef & F. R. Valla (Eds.), *Natufian foragers in the Levant: Terminal Pleistocene social changes in western Asia* (pp. 61–72). Ann Arbor, MI: Berghahn Books.
- Roeser, P. A., Franz, S. O., Litt, T., Ülgen, U. B., Hilgers, A., Wulf, S., ... Melles, M. (2012). Lithostratigraphic and geochronological framework for the paleoenvironmental reconstruction of the last ~36 ka cal BP from a sediment record from Lake Iznik (NW Turkey). *Quaternary International*, *274*, 73–87. <https://doi.org/10.1016/j.quaint.2012.06.006>
- Roffet-Salque, M., Marciniak, A., Valdes, P. J., Pawłowska, K., Pyzel, J., Czeraniak, L., ... Evershed, R. P. (2018). Evidence for the impact of the 8.2-kyBP climate event on Near Eastern early farmers. *Proceedings of the National Academy of Sciences*, *115*, 8705–8709. <https://doi.org/10.1073/pnas.1803607115>
- Rohling, E. J. (2016). Of lakes and fields: A framework for reconciling palaeoclimatic drought inferences with archaeological impacts. *Journal of Archaeological Science*, *73*, 17–24. <https://doi.org/10.1016/j.jas.2016.07.005>
- Rohling, E. J., & Paliike, H. (2005). Centennial-scale climate cooling with a sudden cold event around 8,200 years ago. *Nature*, *434*(7036), 975–979. <https://doi.org/10.1038/nature03421>
- Rollefson, G., Rowan, Y., & Wasse, A. (2014). The Late Neolithic colonization of the Eastern Badiya of Jordan. *Levant*, *46*(2), 285–301. <https://doi.org/10.1179/0075891414Z.00000000046>
- Rose, J. I., Černý, V., & Bayoumi, R. (2013). Tabula rasa or refugia? Using genetic data to assess the peopling of Arabia. *Arabian Archaeology and Epigraphy*, *24*(1), 95–101. <https://doi.org/10.1111/aae.12017>
- Rosen, A. M., & Rivera-Collazo, I. (2012). Climate change, adaptive cycles, and the persistence of foraging economies during the late Pleistocene/Holocene transition in the Levant. *Proceedings of the National Academy of Sciences*, *109*(10), 3640–3645. <https://doi.org/10.1073/pnas.1113931109>
- Rowan, Y. M., Rollefson, G. O., Wasse, A., Abu-Azizeh, W., Hill, A. C., & Kersel, M. M. (2015). The “land of conjecture”: New late prehistoric discoveries at Maitland’s Mesa and Wisad Pools, Jordan. *Journal of Field Archaeology*, *40*(2), 176–189. <https://doi.org/10.1179/0093469015Z.000000000117>
- Rowe, P. J., Mason, J. E., Andrews, J. E., Marca, A. D., Thomas, L., van Calsteren, P., ... Al-Omari, S. (2012). Speleothem isotopic evidence of winter rainfall variability in northeast Turkey between 77 and 6 ka. *Quaternary Science Reviews*, *45*, 60–72. <https://doi.org/10.1016/j.quascirev.2012.04.013>
- Sarikaya, M. A., & Çiner, A. (2015). Late Pleistocene glaciations and paleoclimate of Turkey. *Bulletin of the Mineral Research and Exploration*, *151*(151), 107–127.
- Sharifi, A., Pourmand, A., Canuel, E. A., Ferer-Tyler, E., Peterson, L. C., Aichner, B., ... Swart, P. K. (2015). Abrupt climate variability since the last deglaciation based on a high-resolution, multi-proxy peat record from NW Iran: The hand that rocked the Cradle of Civilization? *Quaternary Science Reviews*, *123*, 215–230. <https://doi.org/10.1016/j.quascirev.2015.07.006>
- Shennan, S., Downey, S. S., Timpson, A., Edinborough, K., Colledge, S., Kerig, T., ... Thomas, M. G. (2013). Regional population collapse followed initial agriculture booms in mid-Holocene Europe. *Nature Communications*, *4*, 2486. <https://doi.org/10.1038/ncomms3486>
- Sherratt, A. (1980). Water, soil and seasonality in early cereal cultivation. *World Archaeology*, *11*(3), 313–330. <https://doi.org/10.1080/00438243.1980.9979770>
- Singh, R. N., Petrie, C. A., French, C. A., Goudie, A. S., Gupta, S., Tewari, R., ... Yadav, S. (2009). Settlements in context: Reconnaissance in western Uttar Pradesh and Haryana. *Man and Environment*, *33*(2), 71–87.
- Smith, P. E. L., Bivar, A. D. H., Fehérvári, G., Burney, C., Ghirshman, R., Weiss, H., ... Young, T. C. (1972). Survey of excavations in Iran during 1970–71. *Iran*, *10*(1), 165–186. <https://doi.org/10.1080/05786967.1972.11834163>
- St. George, S., & Ault, T. R. (2014). The imprint of climate within Northern Hemisphere trees. *Quaternary Science Reviews*, *89*, 1–4. <https://doi.org/10.1016/j.quascirev.2014.01.007>
- Staubwasser, M., & Weiss, H. (2006). Holocene climate and cultural evolution in late prehistoric-early historic West Asia—Introduction. *Quaternary Research*, *66*(3), 372–387. <https://doi.org/10.1016/j.yqres.2006.09.001>
- Stevens, L. R., Djalali, M., Andrieu-Ponel, V., & de Beaulieu, J.-L. (2012). Hydroclimatic variations over the last two glacial/interglacial cycles at Lake Urmia, Iran. *Journal of Paleolimnology*, *47*(4), 645–660. <https://doi.org/10.1007/s10933-012-9588-3>
- Stevens, L. R., Ito, E., Schwab, A., & Wright, H. E. (2006). Timing of atmospheric precipitation in the Zagros Mountains inferred from a multi-proxy record from Lake Mirabad, Iran. *Quaternary Research*, *66*(3), 494–500. <https://doi.org/10.1016/j.yqres.2006.06.008>
- Stevens, L. R., Wright, H. E., & Ito, E. (2001). Proposed changes in seasonality of climate during the Lateglacial and Holocene at Lake Zeribar, Iran. *The Holocene*, *11*(6), 747–755. <https://doi.org/10.1191/09596830195762>
- Stockhecke, M., Sturm, M., Brunner, I., Schmincke, H.-U., Sumita, M., Kipfer, R., ... Anselmetti, F. S. (2014). Sedimentary evolution and environmental history of Lake Van (Turkey) over the past 600 000 years. *Sedimentology*, *61*(6), 1830–1861. <https://doi.org/10.1111/sed.12118>
- Sumner, W. (1994). The evolution of tribal society in the southern Zagros mountains, Iran. In *Chiefdoms and Early States in the Near East: The organizational dynamics of complexity* (pp. 47–65). Madison, WI: Prehistory Press.
- Tengberg, M. (2012). Beginnings and early history of date palm garden cultivation in the Middle East. *Journal of Arid Environments*, *86*, 139–147. <https://doi.org/10.1016/j.jaridenv.2011.11.022>
- Tinner, W., van Leeuwen, J. F. N., Colombaroli, D., Vescovi, E., van der Knaap, W. O., Henne, P. D., ... La Mantia, T. (2009). Holocene environmental and climatic changes at Gorgo Basso, a coastal lake in southern Sicily, Italy. *Quaternary Science Reviews*, *28*(15), 1498–1510. <https://doi.org/10.1016/j.quascirev.2009.02.001>
- Tomonaga, Y., Brennwald, M. S., Livingstone, D. M., Kwiciczen, O., Randlett, M.-È., Stockhecke, M., ... Kipfer, R. (2017). Porewater salinity reveals past lake-level changes in Lake Van, the Earth’s largest soda lake. *Scientific Reports*, *7*, 313. <https://doi.org/10.1038/s41598-017-00371-w>
- Torffing, T. (2015). Neolithic population and summed probability distribution of 14C-dates. *Journal of Archaeological Science*, *63*, 193–198. <https://doi.org/10.1016/j.jas.2015.06.004>
- Torfstein, A., Goldstein, S. L., Stein, M., & Enzel, Y. (2013). Impacts of abrupt climate changes in the Levant from Last Glacial Dead Sea levels. *Quaternary Science Reviews*, *69*, 1–7. <https://doi.org/10.1016/j.quascirev.2013.02.015>
- Touchan, R., Anchukaitis, K. J., Shishov, V. V., Sivrikaya, F., Attieh, J., Ketmen, M., ... Meko, D. M. (2014). Spatial patterns of eastern Mediterranean climate influence on tree growth. *The Holocene*, *24*(4), 381–392. <https://doi.org/10.1177/0959683613518594>
- Turchin, P. (2008). Arise ‘cliodynamics’. *Nature*, *454*(7200), 34–35. <https://doi.org/10.1038/454034a>
- Turner, R., Roberts, N., Eastwood, W. J., Jenkins, E., & Rosen, A. (2010). Fire, climate and the origins of agriculture: Micro-charcoal records of biomass burning during the last glacial–interglacial transition in southwest Asia. *Journal of Quaternary Science*, *25*(3), 371–386. <https://doi.org/10.1002/jqs.1332>
- Turner, R., Roberts, N., & Jones, M. D. (2008). Climatic pacing of Mediterranean fire histories from lake sedimentary microcharcoal. *Global and Planetary Change*, *63*(4), 317–324. <https://doi.org/10.1016/j.gloplacha.2008.07.002>
- Uerpmann, H., Uerpmann, M., & Jasim, S. (2008). The archaeology of Jebel al-Buhais. In H. Uerpmann, M. Uerpmann, & S. Jasim (Eds.), *The natural environment of Jebel al-Buhais: Past and present* (Vol. 2). Tübingen, Germany: Kerns Verlag.

- Ülgen, U. B., Franz, S. O., Biltekin, D., Çatayat, M. N., Roeser, P. A., Doner, L., & Thein, J. (2012). Climatic and environmental evolution of Lake Iznik (NW Turkey) over the last ~4700 years. *Quaternary International*, 274, 88–101. <https://doi.org/10.1016/j.quaint.2012.06.016>
- Ünal-İmer, E., Shulmeister, J., Zhao, J.-X., Tonguç Uysal, I., Feng, Y.-X., Duc Nguyen, A., & Yüce, G. (2015). An 80 kyr-long continuous speleothem record from Dim Cave, SW Turkey with paleoclimatic implications for the Eastern Mediterranean. *Scientific Reports*, 5, 13560. <https://doi.org/10.1038/srep13560>
- Ünal-İmer, E., Shulmeister, J., Zhao, J.-X., Uysal, I. T., & Feng, Y.-X. (2016). High-resolution trace element and stable/radiogenic isotope profiles of late Pleistocene to Holocene speleothems from Dim Cave, SW Turkey. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 452, 68–79. <https://doi.org/10.1016/j.palaeo.2016.04.015>
- Ur, J. (2015). Urban adaptations to climate change in northern Mesopotamia. In S. Kerner, R. Dann, & P. Bangsgaard (Eds.), *Climate and ancient societies* (pp. 69–95). Copenhagen, Denmark: Museum Tusulanum Press.
- Ur, J. (2017). The birth of cities in ancient west Asia. In A. Tsuneki, S. Yamada, & K.-i. Hisada (Eds.), *Ancient west Asian civilization: Geoenvironment and society in the Pre-Islamic Middle East* (pp. 133–147). Singapore: Springer.
- Vaks, A., Bar-Matthews, M., Ayalon, A., Schilman, B., Gilmour, M., Hawkesworth, C. J., ... Matthews, A. (2003). Paleoclimate reconstruction based on the timing of speleothem growth and oxygen and carbon isotope composition in a cave located in the rain shadow in Israel. *Quaternary Research*, 59(2), 182–193. [https://doi.org/10.1016/S0033-5894\(03\)00013-9](https://doi.org/10.1016/S0033-5894(03)00013-9)
- Valentine, B., Kamenov, G. D., Kenoyer, J. M., Shinde, V., Mushrif-Tripathy, V., Otarola-Castillo, E., & Krigbaum, J. (2015). Evidence for patterns of selective urban migration in the Greater Indus Valley (2600–1900 BC): A lead and strontium Isotope mortuary analysis. *PLoS One*, 10(4), e0123103. <https://doi.org/10.1371/journal.pone.0123103>
- Van De Mierop, M. (2008). *The Eastern Mediterranean in the age of Ramesses II*. Chichester, England: Wiley-Blackwell.
- Van Loon, A. F., & Van Lanen, H. (2012). A process-based typology of hydrological drought. *Hydrology and Earth System Sciences*, 16(7), 1915. <https://doi.org/10.5194/hess-16-1915-2012>
- Van Loon, A. F. (2015). Hydrological drought explained. *WIREs Water*, 2(4), 359–392. <https://doi.org/10.1002/wat2.1085>
- van Zeist, W., & Bottema, S. (1977). Palynological investigations in western Iran. *Palaeohistoria*, 19, 19–85.
- van Zeist, W., & Bottema, S. (1991). *Late quaternary vegetation of the Near East*: Reichert.
- Van Zeist, W., & Wright, H. E. (1963). Preliminary pollen studies at Lake Zeribar, Zagros mountains, southwestern Iran. *Science*, 140(3562), 65–67.
- Verheyden, S., Nader, F. H., Cheng, H. J., Edwards, L. R., & Swennen, R. (2008). Paleoclimate reconstruction in the Levant region from the geochemistry of a Holocene stalagmite from the Jeita cave, Lebanon. *Quaternary Research*, 70(3), 368–381. <https://doi.org/10.1016/j.yqres.2008.05.004>
- Vossel, H., Roeser, P., Litt, T., & Reed, J. M. (2018). Lake Kinneret (Israel): New insights into Holocene regional palaeoclimate variability based on high-resolution multi-proxy analysis. *Holocene*, 28(9), 1395–1410. <https://doi.org/10.1177/0959683618777071>
- Walker, M., Head, M.H., Berkehammer, M., Björck, S., Cheng, H., Cwynar, L., Fisher, D., Gkinis, V., Long, A., Lowe, J., Newnham, R., Rasmussen, S.O., and Weiss, H. (2018). Formal ratification of the subdivision of the Holocene Series/Epoch (Quaternary System/Period): two new Global Boundary Stratotype Sections and Points (GSSPs) and three new stages/subseries. *Episodes*, 41(4), 213–223. <https://doi.org/10.18814/epiiugs/2018/018016>
- Wasylikowa, K., & Witkowski, A. (Eds.). (2008). *The palaeoecology of Lake Zeribar and surrounding areas, Western Iran, during the last 48,000 years* (Vol. 8). Königstein, Germany: A.R.G. Gantner Verlag K.G.
- Wasylikowa, K., Witkowski, A., Walanus, A., Hutorowicz, A., Alexandrowicz, S. W., & Langer, J. J. (2006). Palaeolimnology of Lake Zeribar, Iran, and its climatic implications. *Quaternary Research*, 66(3), 477–493. <https://doi.org/10.1016/j.yqres.2006.06.006>
- Watts, W. A. (1985). A long pollen record from Laghi di Monticchio, southern Italy: A preliminary account. *Journal of the Geological Society*, 142(3), 491–499.
- Weiss, B. (1982). The decline of Late Bronze Age civilization as a possible response to climatic change. *Climatic Change*, 4(2), 173–198.
- Weiss, H. (2003). Ninevite 5 periods and processes. In *The origins of north Mesopotamian Civilization. Subartu IX* (pp. 593–624). Turnhout, Belgium: Brepols.
- Weiss, H. (2015). Megadrought, collapse, and resilience in late 3rd millennium BC Mesopotamia. In H. W. Arz, R. Jung, H. Meller, & R. Risch (Eds.), *2200 BC-Ein Klimasturz als Ursache für den Zerfall der Alten Welt* (pp. 35–52). Halle (Saale), Germany: Landesamt für Denkmalpflege und Archäologie Sachsen-Anhalt, Landesmuseum für Vorgeschichte.
- Weiss, H. (2016). Global megadrought, societal collapse and resilience at 4.2–3.9 ka BP across the mediterranean and west Asia. *PAGES*, 24, 62–63.
- Weninger, B., Alram-Stern, E., Bauer, E., Clare, L., Danzeglocke, U., Joris, O., ... van Andel, T. (2006). Climate forcing due to the 8200 cal yr BP event observed at Early Neolithic sites in the eastern Mediterranean. *Quaternary Research*, 66(3), 401–420. <https://doi.org/10.1016/j.yqres.2006.06.009>
- Weninger, B., Lee, C., Gerritsen, F., Horejs, B., Krauß, R., Linstädter, J., ... Rohling, E. J. (2014). Neolithisation of the Aegean and southeast Europe during the 6600–6000 calBC period of rapid climate change. *Documenta Praehistorica*, 41, 1.
- Wheater, H., Sorooshian, S., & Sharma, K. D. (Eds.). (2007). *Hydrological modelling in arid and semi-arid areas*. Cambridge, England: Cambridge University Press.
- White, F., & Léonard, J. (1990). *Phytogeographical links between Africa and southwest Asia*. Paper presented at the proceedings of the Third Plant Life of southwest Asia Symposium, Berlin, Germany.
- Whitehead, P., Smith, S., & Wade, A. (2011). Modelling water resources and climate change at the Bronze Age site of Jawa in northern Jordan: A new approach utilising stochastic simulation techniques. In S. Mithen & E. Black (Eds.), *Water, life and civilisation: Climate, environment and society in the Jordan Valley* (p. 289). Cambridge, England: Cambridge University Press.
- Whitehead, P. G., Smith, S. J., Wade, A. J., Mithen, S. J., Finlayson, B. L., Sellwood, B., & Valdes, P. (2008). Modelling of hydrology and potential population levels at Bronze Age Jawa, Northern Jordan: A Monte Carlo approach to cope with uncertainty. *Journal of Archaeological Science*, 35(3), 517–529. <https://doi.org/10.1016/j.jas.2007.05.001>
- Wick, L., Lemcke, G., & Sturm, M. (2003). Evidence of Lateglacial and Holocene climatic change and human impact in eastern Anatolia: High-resolution pollen, charcoal, isotopic and geochemical records from the laminated sediments of Lake Van, Turkey. *The Holocene*, 13(5), 665–675. <https://doi.org/10.1191/0959683603hl653rp>
- Wicks, K., Finlayson, B., Marićević, D., Smith, S., Jenkins, E., & Mithen, S. (2016). *Dating WF16: Exploring the chronology of a Pre-Pottery Neolithic A settlement in the Southern Levant*. Proceedings of the Prehistoric Society, 82, 73–123.
- Wilkinson, T. J., Ur, J., & Hritz, C. (2013). Settlement archaeology of Mesopotamia. In T. J. Wilkinson, M. Gibson, & M. Widell (Eds.), *Models of Mesopotamian landscapes. How small-scale processes contributed to the growth of early civilizations* (pp. 34–55). Oxford, England: Archaeopress.
- Wilkinson, T. J., Bintliff, J., Curvers, H. H., Halstead, P., Kohl, P. L., Liverani, M., ... Courty, M.-A. (1994). The structure and dynamics of dry-farming states in Upper Mesopotamia [and comments and reply]. *Current Anthropology*, 35(5), 483–520.
- Wilkinson, T. J., & Rayne, L. (2010). Hydraulic landscapes and imperial power in the Near East. *Water History*, 2(2), 115–144.
- Willcox, G., Buxo, R., & Herveux, L. (2009). Late Pleistocene and early Holocene climate and the beginnings of cultivation in northern Syria. *The Holocene*, 19(1), 151–158. <https://doi.org/10.1177/0959683608098961>
- Woldring, H., & Bottema, S. (2001). The vegetation history of East-Central Anatolia in relation to archaeology: The Eski Acıgöl pollen evidence compared with the Near Eastern environment. *Palaeohistoria*, 43/44, 1–34.
- Woodbridge, J., Fyfe, R. M., Roberts, N., Downey, S., Edinborough, K., & Shennan, S. J. (2012). The impact of the Neolithic agricultural transition in Britain: A comparison of pollen-based land-cover and archaeological 14C date-inferred population change. *Journal of Archaeological Science*, 51, 216–224. <https://doi.org/10.1016/j.jas.2012.10.025>

- Woodbridge, J., & Roberts, N. (2011). Late Holocene climate of the Eastern Mediterranean inferred from diatom analysis of annually-laminated lake sediments. *Quaternary Science Reviews*, 30(23–24), 3381–3392. <https://doi.org/10.1016/j.quascirev.2011.08.013>
- Woodhouse, C. A. (2001). A tree-ring reconstruction of streamflow for the Colorado Front Range 1. *Journal of the American Water Resources Association*, 37(3), 561–569.
- Wright, N. J., Fairbairn, A. S., Faith, J. T., & Matsumura, K. (2015). Woodland modification in Bronze and Iron Age central Anatolia: An anthracological signature for the Hittite state? *Journal of Archaeological Science*, 55, 219–230. <https://doi.org/10.1016/j.jas.2014.12.021>
- Wright, R. P. (2010). *The ancient Indus: Urbanism, economy, and society*. Cambridge, England: Cambridge University Press.
- Xoplaki, E., Gonzalez-Rouco, J. F., Luterbacher, J., & Wanner, H. (2004). Wet season Mediterranean precipitation variability: Influence of large-scale dynamics and trends. *Climate Dynamics*, 23(1), 63–78. <https://doi.org/10.1007/s00382-004-0422-0>
- Yiğitbaşıoğlu, H., Dean, J. R., Eastwood, W. J., Roberts, N., Jones, M. D., & Leng, M. J. (2015). A 600 year-long drought index for central Anatolia [Special Issue]. *Journal of the Black Sea/Mediterranean Environment*, 84–88.
- Zazzo, A., Munoz, O., Badel, E., Beguier, I., Genchi, F., & Marcucci, L. G. (2016). A revised radiocarbon chronology of the Aceramic Shell Midden of Ra's Al-Hamra 6 (Muscat, Sultanate of Oman): Implication for occupational sequence, marine reservoir age, and human mobility. *Radiocarbon*, 58(2), 383–395. <https://doi.org/10.1017/RDC.2016.3>
- Zohary, M. (1973). *Geobotanical foundations of the Middle East*. Stuttgart, Germany: Gustav Fischer Verlag.
- Zolitschka, B., Francus, P., Ojala, A. E. K., & Schimmelfmann, A. (2015). Varves in lake sediments—A review. *Quaternary Science Reviews*, 117, 1–41. <https://doi.org/10.1016/j.quascirev.2015.03.019>